



**ADDIS ABABA UNIVERSITY
SCHOOL OF GRADUATE STUDIES
FACULTY OF SCIENCE
DEPARTMENT OF EARTH SCIENCES**

**VELOCITY STRUCTURE OF THE UPPERMOST CRUST ALONG
THE MAIN ETHIOPIAN RIFT**



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ABSTRACT

Ethiopia is located at a plate tectonic triple-rift junction (Mohr 1980) where the East African Rift System meets the Gulf of Aden and Red Sea with spreading axes at Afar.

A seismic study over the area has obvious implications on our knowledge concerning the structure and evolution of the Earth. This study focuses on the determination of the velocity structure of the uppermost crust along the Main Ethiopian Rift, which runs from Awassa to Nazareth to a depth of 10km on a 104km seismic profile.

Seismic waves generated by explosive charges detonated in boreholes are used to study the area. The seismograph Texans deployed on the surface are used to record the seismic energy refracted back to the surface from geological interfaces.

The time of first arrivals of these refracted waves were picked, processed and inverted using the General Reciprocal Method (GRM).

Basically three refractors at an average depth of, 2000m, 4500m, and 7000m were identified with their corresponding average P wave velocities of 3950m/s, 5726m/s and 6320m/s. The geologic sections of the area are shown in Fig (6.11).

Attempts have been made to correlate the seismic layers with the various stratigraphic units.

1. INTRODUCTION

Seismic investigation utilizes the fact that elastic waves also called seismic waves travel with different velocities in different rocks. By generating seismic waves at a point and observing the times of arrival of these waves at a number of other points on the surface of the earth, it is possible to determine the velocity distribution and locate subsurface Interfaces where the Waves are reflected or refracted. The aim of this thesis is to map the P wave velocity distribution of the layers beneath the uppermost crust along the main Ethiopian rift.

1.1 Previous works

There was no detailed seismic study over the area. Recently the Ethiopian Afar Geo Scientific Lithospheric Experiment(EAGLE) group made a large scale seismic study in three phases. With a consequence, a number of geological studies has been conducted and reported.

1.2 Physiographic outlines

The area of study runs from Awassa to Nazareth along the main Ethiopian rift in a 104km profile line. This research is based on the data from the Ethiopian Afar Geoscientific Lithospheric Experiment (EAGLE), which was undertaken in the years 2001-2003 in Ethiopia.

1.3 **Aim of the present study.**

The major objectives of this work are,

1. To understand the velocity structure of the uppermost crust along the main Ethiopian rift.
2. To correlate the seismic layers with the various stratigraphic units and to understand the geologic history of the region.
3. To map the major structural features like faults, fractures etc
4. To prepare a geodynamic and geologic model of the uppermost crust along the main Ethiopian rift.

1.4 **Methodology**

Seismic refraction method has been employed to produce the seismic model along the main Ethiopian Rift. In processing the data the GRM method was mainly used.

The geoseismic and the geologic model were then interpreted using the seismic P wave velocities to correlate the results obtained with previous regional geological works.

2. GENERAL GEOLOGY

The geology of the area is mainly controlled by activities, which control the geology of the country. These geologic activities are, the formation of the Precambrian basement and associated deformation, the deposition of Mesozoic sedimentary succession, the uplifting volcanism and rifting of the Cenozoic and much more recently the quaternary climate variation and associated lake level changes. All these geologic activities leave their imprint in the geology of the study area. The Cenozoic uplift and associated rifting have been associated with volcanic activity. This activity leads to the eruption of volcanic rocks such as basalts (the older trap series basalt and the much more recent basaltic scoria cones), ignimbrite, trachite, rhyolites

and pumiceous and pyroclastic. The faulting also formed structural depressions and calderas, which are latter filled by lakes. The lakes occupied a wide area in the quaternary before they shrunk to the present day size due to climate change. The shrinking lakes left behind lacustrine sediments that cover the low-lying plain. Nevertheless, the dominant rocks in the study area are volcanic products of the Cenozoic and the quaternary to recent fluvio-lacustrine sediments. The Ethiopian volcanics were divided in to two main series: trap series or (plateau series) and Aden volcanic series or (Rift series)(Banford, 1869;Mohr, 1971).

TRAP SERIES

An early Cenozoic extrusion of flood lavas occurred from fissures and centers covered the great part of the Mesozoic sedimentary rocks.

ADEN VOLCANIC SERIES

Aden volcanic series rocks post date the formation of the rift system in Ethiopia,i.e post Miocene or Pilo-Quaternary.

2.1 Stratigraphic outline of the volcanic products.

In the central sector of the Main Ethiopian Rift, the rift shoulders are made of volcanic and pyroclastic rocks where as large areas of the rift floor are covered by volcano-lacustrine and flvio-lacustrine deposits. Di Paola(1972) reconstructed the stratigraphic succession of major products in to eight major categories.

2.1.1 Basalt and ignimbrite of the plateau trap series

The trap series is mostly made up of basaltic lava flows with subordinately associated ignimbrite units intercalated between the basic lavas. Also silicic lava flows and lava domes are sometimes associated with the basalt of the trap series. The age varies from 40 to 25ma for the basalt and from 37 to 27 Ma for the rhyolite. This rock unit is exposed around east of lake Langano. The unit consists sub horizontal flows of a few meter thick lava with rather well preserved scoraceous level between the flows.

2.1.2 Trachytic flows and domes

The eastern plateau is covered by Pliocene to early Pleistocene (4.6-1.6Ma) shield volcanoes (chilalo, 4005m, kecha 4245m badda, 4179m). The unit consists of trachite with subordinate basalts and mugearites (Di Paola, 1972; Giday Woldegabriel et al., 1990; Benvenuti et al., 2002)

2.1.3 Rift pyroclastic formation.

The formation covers most of the rift floor and includes three principal rock types: typical ignimbrite, sillar and layered pumice. They are the most ancient formation outcropping on the floor of the rift valley. They are early to middle Pliocene (4.3-3Ma) formations. (Giday Woldegabriel et al., 1990).

2.1.4 Old rhyolitic lava flows and domes.

These are alkaline and per alkaline rhyolitic lava flows and domes which was erupted later as viscous lava after the magma has lost its gases during previous explosive activity. Thin and subordinated layers of unwelded pumice and ash are often found interbedded in these silicic lava flows and domes. This lava was erupted from late Pliocene to middle Pleistocene and in some places crop out as remnants of large calderas and often shows vertical sheeting or flow folding.

2.1.5 Basalt Hyaloclastites

Limited occurrences of sub aqueous basaltic volcanism are found in the rift valley. Hyaloclastite occur both as small monogenic rug and as horizontal layer of local extension. They are exposed around lake chitu on the southwestern shore of lake shalla and around lake ziway. These rocks consist of fine glassy material, generally yellowish to brown in colour. They contain small boulder of basaltic lava.

2.1.6 Recent Basalt

Recent basaltic cones are elongated parallel to the main tectonic trend of the rift and were produced by fissural eruption. The alignment of the spatter cones parallel the rift marks the feeding fissures. These flows consist of lots of scoria cones. Two main field of recent basalt having an area of 800 km² are located close to western and eastern escarpment. The lava fields close to the Butajira-Slite in the west and the spatter cones east of lake Ziway in the east.

2.1.7 Recent rhyolitic pyroclastic and obsidian lava flows.

These rocks are the latest volcanic products in the region. Many young volcanoes and calderas such as the bora-bericco complex, the Alutu volcano and Chabbi volcanoes that grew inside the Corbetti caldera, are made up of rhyolitic lava flows, pumice falls and ashes.

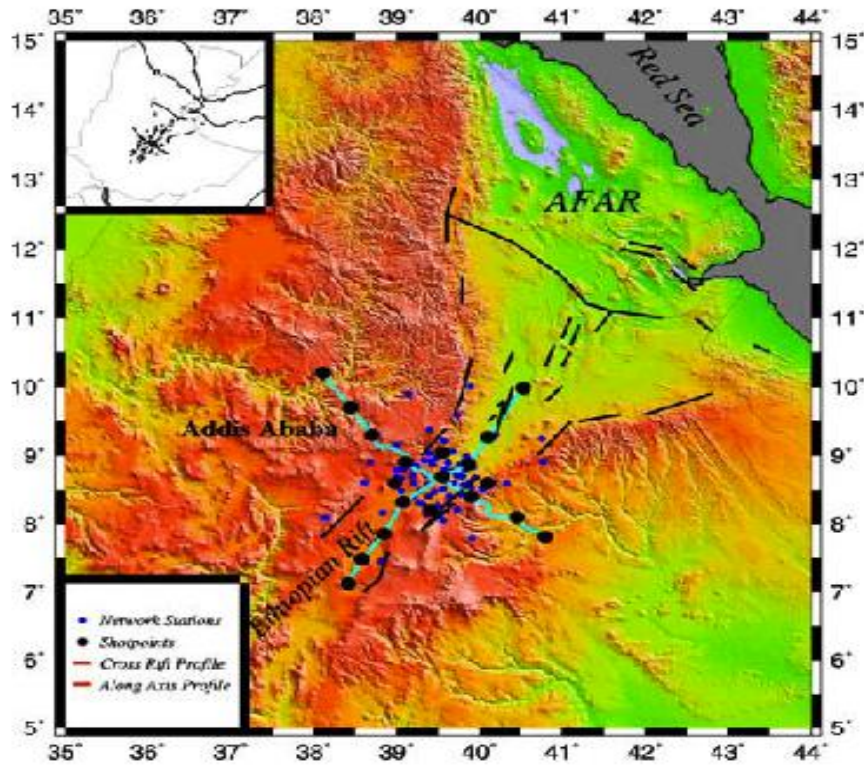
2.1.8 Lacustrine sediments

Lacustrine sediments are common in the rift valley. they are the result of drying up of a big lake that occupie the floor of the rift valley during pluvial periods. The lacustrine sediments include silts, clays and volcano clastic sediments.

3. PROJECT EAGLE

3.1 Introduction

The Ethiopian Afar Geoscientific Lithospheric Experiment(EAGLE) is a major initiative to investigate how the African continent is splitting along the Ethiopian Rift in the Horn of Africa. The project involves the Universities of Leicester, Leeds and Royal Holloway from the United Kingdom, Stanford University, California and the University of Texas, at El Paso both from the United States. There is significant involvement from other European and US institutions. These groups are working with principal counterparts from the Geophysical Observatory and the Department of Geology and Geophysics at the University of Addis Ababa together with many scientists from the Ethiopian Geological Survey and the Petroleum Operations Department of the Ethiopian Ministry of Mines. The Commission of Science and Technology of the Federal Democratic Republic of Ethiopia coordinate the whole project. The work involves seismic recording from controlled explosions and natural Earthquakes, to provide an image of the top 100km or so of the Earth over a region encompassing the north-eastern part of the main Ethiopian Rift where it enters Afar (Fig 3.1). As continental rift zones evolve to seafloor spreading they do so through progressive episodes of lithospheric stretching, heating and magmatism, yet the actual process of continental break-up is poorly understood. The East African rift system in northeastern Ethiopia is central to our understanding of this process as it lies at the transition between continental and oceanic rifting (Ebinger and casey,2001).



Fig(3.1) study area showing profiles 1,2 and 3

This research is based on the data from the Ethiopian Afar Geoscientific Lithospheric Experiment (EAGLE), which was undertaken in the years 2001-2003 in Ethiopia. The study focuses on the determination of the velocity structure of the uppermost crust along the main Ethiopian Rift which runs from Awassa to Nazareth. Seismic waves generated by explosive charges detonated in boreholes of 50 meters deep and 50 inch diameter are used to study the area. Those waves refracted from geological interfaces back to the surface are recorded by the Seismograph Texans, which are deployed on the surface of the Earth separated with an average distance of 1km.

The time of first arrivals of refracted waves are then processed using the appropriate geophysical soft wares and is interpreted in terms of the geological boundaries and structures of the region.

4. Basics of the methods used in this work

4.1 Seismic waves

The particles in the surface of the Earth are in constant vibrations due to continuous disturbances from natural as well as artificial sources. Those disturbances also called Seismic waves have long been observed for a variety of reasons, and the number of applications continues to grow. Generally seismic waves can be considered as messengers that convey information about the Earth's interior. Therefore systematically studying these waves has obvious implications for man's safety, as well as for his curiosity concerning the structure and evolution of the Earth and is the primary concern of exploration seismology. In the application of seismic techniques to the determination of geological structures, the source of seismic radiation is usually man-made, such as an explosive charge.

4.2 Types of seismic waves

When a stress is applied to an elastic body (as when it is hit with a hammer) or when the stress is suddenly released (as when a previous state of stress is altered by fracturing), the corresponding strain (deformation pattern) is propagated outwards as an elastic wave. There are two principal types of elastic waves: body and surface waves.

4.2.1 Seismic body waves

We can categorize seismic body waves as P and S waves depending on the way the wave propagates. As the name indicates body waves can travel directly through a mass of some substance in any direction.

P waves also called compressional or longitudinal waves have vital importance in exploration seismology. It propagates in an elastic medium causing particles to vibrate back and fourth in line with the direction of the advancing wave.

On the other hand S waves also called shear waves or transverse waves propagates in a medium causing particles to vibrate perpendicular to the advancing wave.

4.2.2 Surface waves

Surface waves can be explained in terms of combinations of body waves that interfere with one another in distinctive ways when traveling near the surface of a mass. Surface waves are of two types.

- a. Rayleigh waves
- b. Love waves

An important characteristic of surface waves is their dispersion. By dispersion, we mean the dependence of velocity on wavelength. For example, in a layer with elastic properties gradually varying with depth, the velocity of love waves for short wavelengths is equal to V_s in the upper part of the layer and for long wavelengths it approximates to the velocity of shear wave (V_s) in the lower part.

4.3 Seismic wave velocities

The capacity of a material to be temporarily deformed by passing seismic waves can be described by its physical properties. Therefore the physical properties of a material like elasticity and density determines the velocity of seismic waves through it.

The velocities of P and S waves in different materials are given by the following equations.

$$V_p = \sqrt{\frac{k + \frac{4}{3}\mu}{\rho}} \quad \text{and} \quad V_s = \sqrt{\frac{\mu}{\rho}}$$

Where:- V_p is the velocity of P wave

V_s = the velocity of S wave

K is the bulk modulus

μ is the shear modulus

ρ is the bulk density of the rock(material)

4.4 Principles of refraction seismic

The three principles (law's) which are back bones of refraction seismic are

- a. Fermat's principle
- b. Snell's law
- c. Huygen's principle

➤ Fermat's principle.

According to Fermat's principle, seismic waves generated either naturally or by an artificial source always follows the path requiring the least time.

Of all waves that are generated during an explosion, P waves are the first ones to arrive on the seismograph obeying the three basic laws while traversing inside the earth.

Refraction of seismic wave occurs when they intercept a media of different acoustic property as shown in Fig(4.1).

The way a seismic wave refracted could be explained using snell's law, which states that “the direction of the reflected and refracted waves traveling away from a boundary depends on the direction of the incident waves and their speeds”

$$\text{Mathematically, } \frac{V_1}{V_2} = \frac{\sin i}{\sin r} \quad (\text{Snell's law})$$

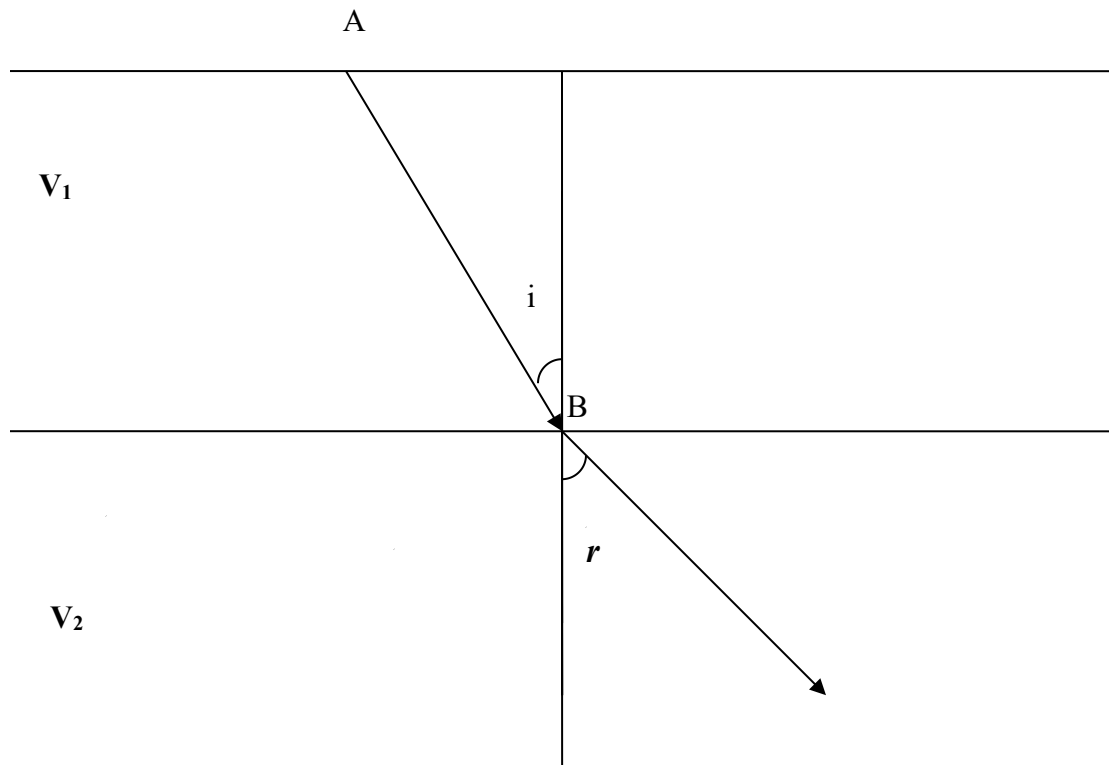
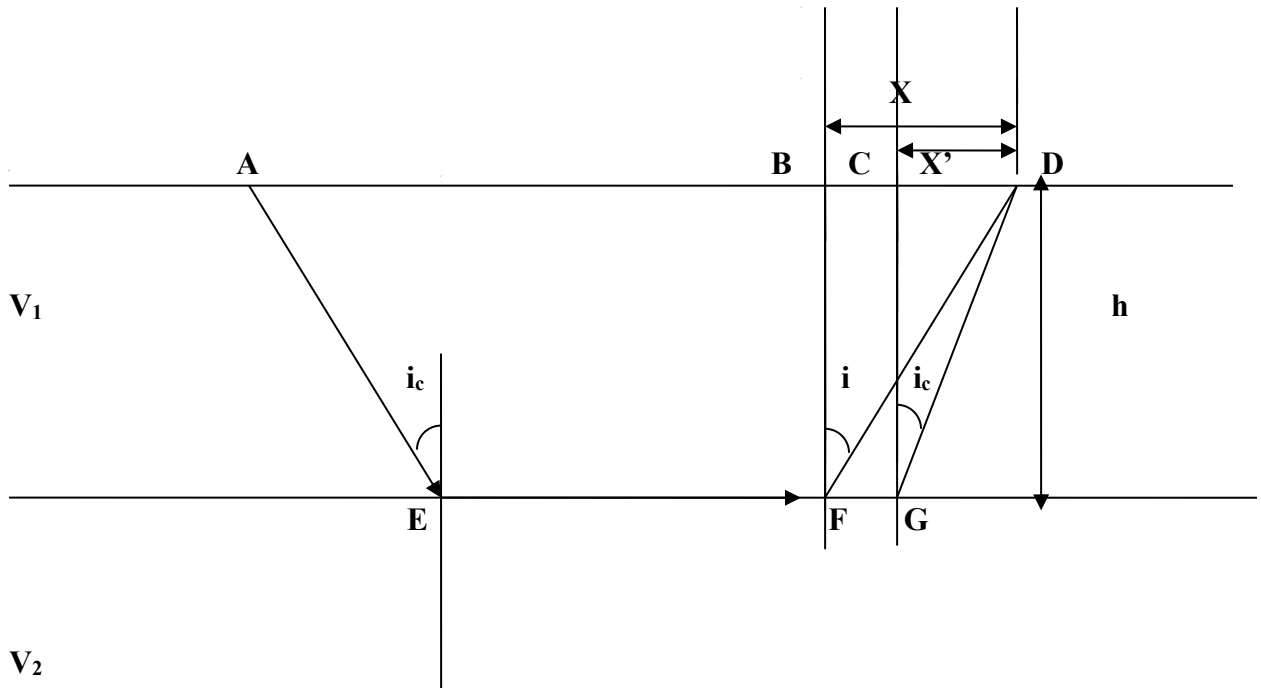


Figure (4.1) diagram showing refraction of seismic wave

Where, V_1 is the velocity of the seismic wave in medium one and,
 V_2 is its velocity in medium two.

➤ **The following are the three consequences of Snell's law**

1. If $v_1 = v_2 \Rightarrow i = r$, there will be no refraction.
2. If $v_1 > v_2 \Rightarrow i > r$, then the refracted wave bends towards the normal and there will be no way to bring it up to the surface.
3. If $v_1 < v_2 \Rightarrow i < r$, then we will have a refracted wave coming up to the surface.



Fig(4.2)Diagram showing critically refracted seismic wave

When $r = 90^\circ$, $\Rightarrow \sin i_c = v_1/v_2$ (critical angle of incidence) and the refracted wave travels along the interface separating the two media Fig 4.2.

We ignore all other waves except the Head wave which is the one striking the interface at an angle of i_c .

The critically refracted wave is made of infinite amount of point explosions, which simulate infinite amount of explosions inside the earth in all direction. Hence energy from within the earth is again released and reaches the surface to be recorded. This principle is called Huygen's principle and is stated formally as follows. "Every point on a wave front is a source of a new waves that travels away from it in all directions".

After the seismic energy has been critically refracted and started to rise up according to Huygen's principle, the seismic wave that arrives first on the surface is the refracted wave,

which has the same angle as the critical angle of incidence. Mathematically this can be shown as follows.

Considering the seismic radiation propagating along the two paths (path AEFD and path AEGD) as shown in fig (Fig 4.2) we have,

T_1 = Time to travel through path one (AEFD) and

T_2 = Time for path two (AEGD)

From the diagram we can easily see that,

$$T_1 = AE/V_1 + EF/V_2 + FD/V_1 \quad \text{and}$$

$$T_2 = AE/V_1 + EG/V_2 + GD/V_1.$$

Therefore, the difference between T_2 and T_1 will give us,

$$\begin{aligned} T_2 - T_1 &= (AE/V_1 + EG/V_2 + GD/V_1) - (AE/V_1 + EF/V_2 + FD/V_1) \\ \Rightarrow T_2 - T_1 &= (AE/V_1 + EF/V_2 + FG/V_2 + GD/V_1) - (AE/V_1 + EF/V_2 + FD/V_1) \\ \Rightarrow T_2 - T_1 &= FG/V_2 + GD/V_1 - FD/V_1 \\ \Rightarrow T_2 - T_1 &= FG/V_2 + (GD - FD)/V_1 \end{aligned} \quad \text{Eqn 4.1}$$

But from simple trigonometry, we have the following.

$$\tan i_c = x'/h \Rightarrow x' = h \tan i_c$$

$$\tan i = x/h \Rightarrow X = h \tan i$$

$$X - X' = h(\tan i - \tan i_c)$$

$$FD = h/\cos i \quad \text{and}$$

$$GD = h/\cos i_c$$

Finally substituting the above expressions in equation 4.1, we will have.

$$T_2 - T_1 = \frac{h}{V_1 \cos i} [\cos(i_c - i) - 1] \quad \text{Eqn 4.2}$$

Since the term in the bracket is always negative, the result $T_2 - T_1$ will be negative, showing that T_2 is less than T_1 .

Therefore the refracted seismic radiation, which arrives first at the geophone on the surface, follows the path AEGD.

Generalized Reciprocal method(GRM)

This method was first developed and later introduced by the Australian scientist Derecke Palmer in 1980. It is the powerful method of all seismic refraction methods. The GRM, as described in palmer(1980,1981), centers on the key aspect that a geophone spacing, termed the XY spacing, can be found whereby up going rays from forward and reverse shots depart from the same point on a refractor. Once this spacing is found, it is theoretically possible to map refracting surfaces accurately and to determine an average velocity for the layers overlying the refractor. The average velocity allows depth to the refractor to be calculated; irrespective of whether undefined velocity inversions or hidden layers are present. The travel times at two geophones, separated by a variable distance XY (Fig 4.3) are used in refractor velocity analysis and time depth calculations. At the optimum XY spacing, the upward traveling segment of the rays to each geophone emerge from near to the same point on the refractor. This results in the refractor velocity analysis being the simplest and the time-depths showing the most detail.

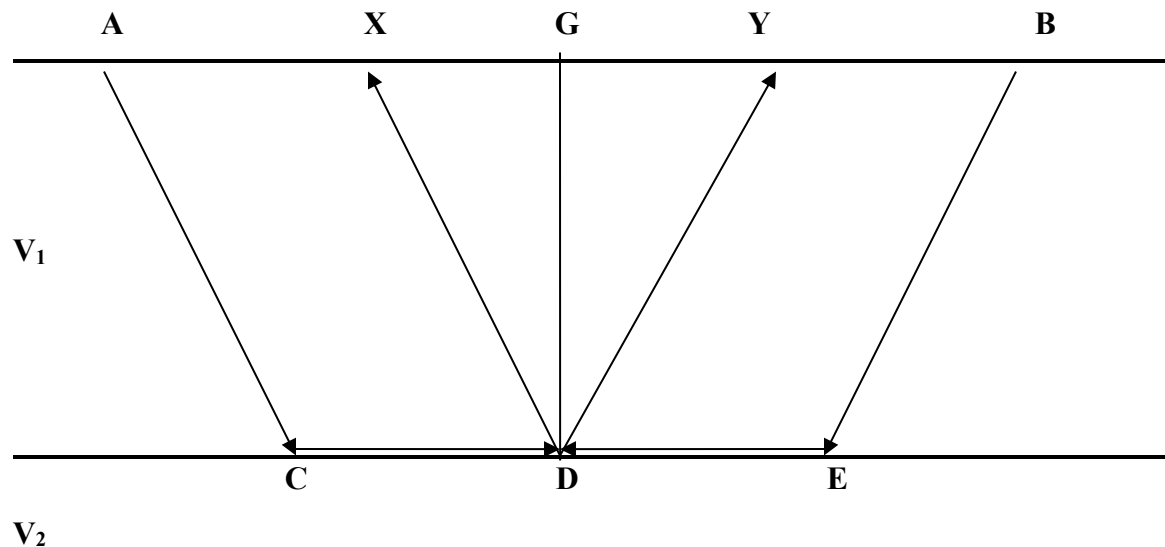


Figure (4.3) Diagram showing the forward and reverse shots with the variable distance XY

➤ Undetected layers

The hidden layer problem occurs where energy from a refractor of higher seismic velocity arrives at the surface before energy from an overlying refractor.

Another type of undetected layer is the velocity inversion problem, which occurs when a layer has a lower seismic velocity than the layer above it.

The GRM provides a means of recognizing and accommodating undetected layers through the two basic approaches of computing and observing the optimum XY values.

There are two distinct approaches for the determination of the optimum XY value, which is probably the most difficult, and the most important aspect of the GRM.

1. Direct calculation of XY values.
2. Observation of XY values.
- 3.

➤ Direct calculation of xy values

We can directly calculate the optimum XY from the seismic velocities and thickness of the layers using the formula,

$$XY_{\text{optimum}} = 2 \sum_{j=1}^{n-1} Z_j G \tan i_{jn} \text{-----Eqn 4.3}$$

Where, $i_{jn} = \sin^{-1}(V_j/V_n)$

Hence an optimum XY value can be calculated for any seismic velocity versus depth section.

➤ Observation of XY values

- The velocity analysis function(T_v)

The velocity analysis function T_v is defined by the equation,

$$T_v = 1/2(t_{AY} - t_{BX} + t_{AB}) \quad \text{Eqn (4.4) Where,}$$

t_{AY} = forward travel time from shot A to the geophone at Y
 t_{BX} = reverse travel time from shot B to the geophone at X
 t_{AB} = the reciprocal time between the two shot points A and B

The value of this function is referred to G, which is mid way between X, and Y.

The values of T_v calculated using the above equation are plotted against distance for Different XY values until we get the optimum xy value.

The inverse of an apparent refractor velocity V_n' is defined as the slope of a line fitted to the T_v values for the optimum XY.i.e $d/dx(Tr) = 1/V_n'$ and it can be shown (palmer 1980) that $V_n \approx V_n' \cos \theta_{n-1}$, where V_n is the true refractor velocity and θ_{n-1} is the dip of the refractor.

- **The generalized time depth**

Using the symbols in (Fig 4.2) the generalized time depth at G is defined by the equation,

$$t_g = \frac{1}{2} \left[t_{AY} + t_{BX} - \left(t_{AB} + \frac{XY}{v_n} \right) \right] \text{-----Eqn 4.5}$$

The time depth function measures the depth to the refractor in units of time.

- **The depth conversion factor**

For plane layering between the forward and reverse arrival times, eqn (4.5) can be shown to reduce to,

$$t_g = \sum_{j=1}^{n-1} Z_j G / V_{jn}$$

Where,

$$V_{jn} \approx V_n' V_j' / (v_n'^2 - V_j'^2)^{1/2}$$

The insensitivity of the depth conversion factor to dip angles makes the GRM an extremely convenient method for dealing with irregular refractors, including those overlain by a layer with in which the velocity varies continually with depth.

Note that t_g is measure of depth in units of time and ZG is measure of depth in units of space.

Therefore, $ZG = Tg.DCF$, where DCF =depth conversion factor.

The optimum XY value obtained using the above two equations is known as the observed XY value.

If the depth section is to be consistent with the travel time data, the computed and observed XY values must agree. If these values do not agree, then undetected layers are indicated.

i.e If $XY_{calc} = XY_{obs}$ (no hidden layer)

If $XY_{calc} > XY_{obs}$ (velocity inversion)

If $XY_{calc} < XY_{obs}$ (thin layer)

For a hidden layer problem the average velocity above the refractor is used to calculate the depth to the refractor with out defining all layers.

The average velocity is defined in equation (4.6) below.

$$\bar{V} = [V'_n XY / (XY + 2t_g V'_n)]^{1/2} \text{ -----Eqn 4.6}$$

And the depth to the refractor is defined as,

$$Z_{ref} = T G \frac{\bar{V}}{\cos i} \text{ where, } i = \sin \frac{\bar{V}}{V'} \text{ ---Eqn 4.7}$$

5. Data acquisition

5.1 The controlled source seismic study

Controlled Source Seismology involves the detonation of seismic explosive charges in deep boreholes distributed along profiles of seismic recording instruments deployed at regular intervals along the Earth's surface. The seismic waves generated by the charges travel down into the crust of the Earth and are reflected and refracted back to the surface from geological interfaces at depth. The seismic instruments record the resulting seismic

waves as they reach the surface along the profiles. The time of arrival of the reflections and refractions at the recording instruments will be interpreted by the research seismologist in terms of the distribution of geological boundaries deep within the Earth. This image is used to constrain our understanding of the geological processes that form these deep geological structures. In the case of the EAGLE project, the processes are those that have generated the Ethiopian Rift at its junction with the region of Afar.

5.2 Report on shot details

Preparations for the controlled source effort began in April 2001 and 91 broadband instruments were first deployed at a nominal 5km interval along the 450km cross rift profile in November – December 2002. The controlled experiment took place in January of 2003 and involved 18 sources in boreholes (typically 1 mt dynamite at 50 m depth) and 2 sources in lakes (Shalla and Arenguade) and the deployment of ~1000 short-period recorders. An environmental impact assessment required by the Ethiopian government incorporating a full report on the lake flora, fauna, and chemistry both before and after the shots were detonated, should provide data of immense value for future seismic studies elsewhere in the world. The controlled source seismic study was undertaken in January 2003. The details of this part of the project, which involved more than 70 scientists deploying nearly 1000 recorders over an 8 day period, which instruments recorded the seismic waves generated from the detonation of 19 charges in both boreholes and lakes involved a major logistical exercise. The results will provide a very high-resolution image of the top 30 km of the Earth both across and along the axis of the Main Ethiopian Rift and also a detailed 3-D image of the subsurface beneath the central Nazret volcanic segment. They will determine the amount of extension that has already occurred across the Rift; provide understanding about the geological controls on the process of continental break-up, illuminate the magmatic processes that are occurring deep within the crust of the Earth in this volcanic region, information about the distribution of near surface sedimentary and basement rocks and so aid in understanding of the distribution of mineral resource in the study area. They will also provide further information about the possible distribution of geothermal resource beneath the controlled source project area.

5.3 Recording instrument deployment phase

Phase III consisted of the deployment of a further 1100 seismic instruments during a controlled source seismic project involving 20 seismic charges being fired into one 450km cross-rift profile (Profile 1), one 450km axial profile (Profile 2), and a dense 2-D array of instruments in a 150km diameter circle around the profiles' intersection (Profile 3), all centered on the magmatically active Nazret region Fig (5.1). The crust and upper mantle velocity models to be derived will be used to provide estimates of total crustal thinning across the rift, to assess the role of basement in the location of major faults and magmatic segments, and to determine whether significant under plating takes place during the syn-rift stage. The high-density array around the intersection of the two profiles was designed to image any magma bodies beneath the Boset magmatic segment.

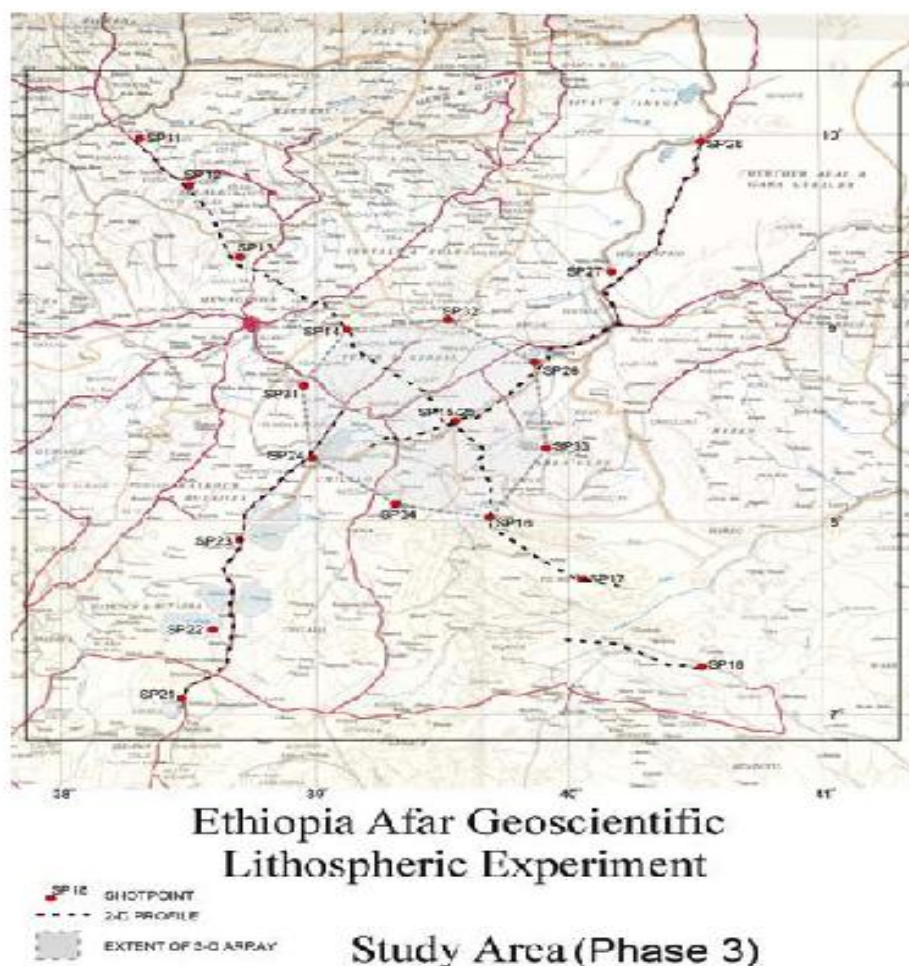


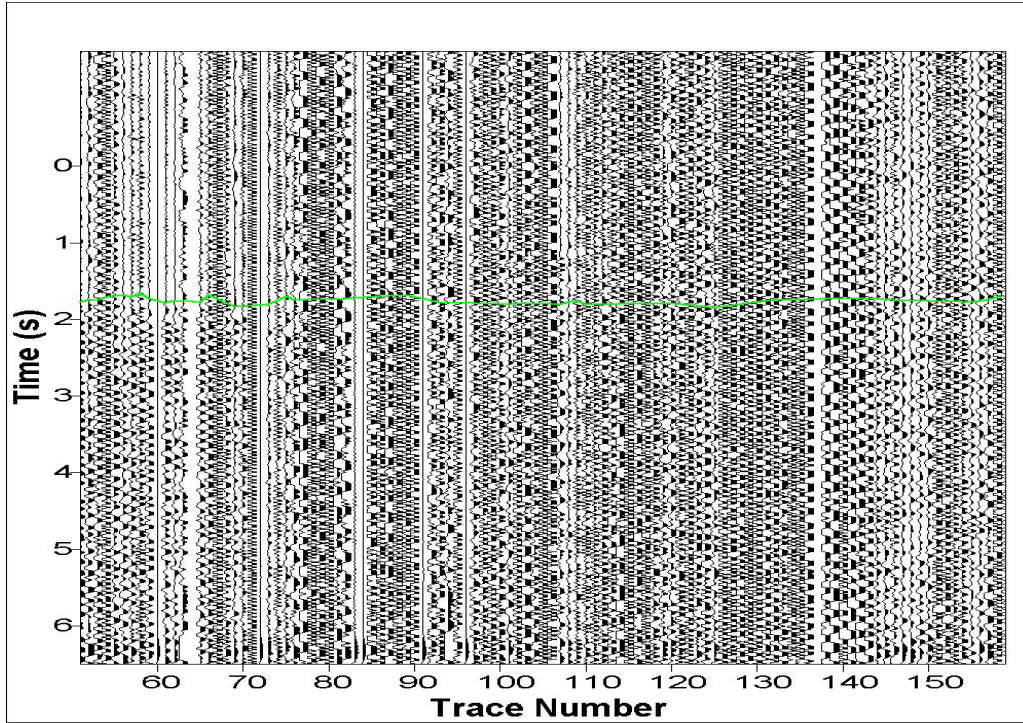
Fig (5.1)

Shot point	Name	Easting	Northing	Elevation (m)
21	Awassa	441578	785718	1706
22	Lake shalla	538072	967557	1797
23	Ziway	469359	870684	1653
24	Koka	500547	920584	1602
25	Doni B	568063	943141	1236

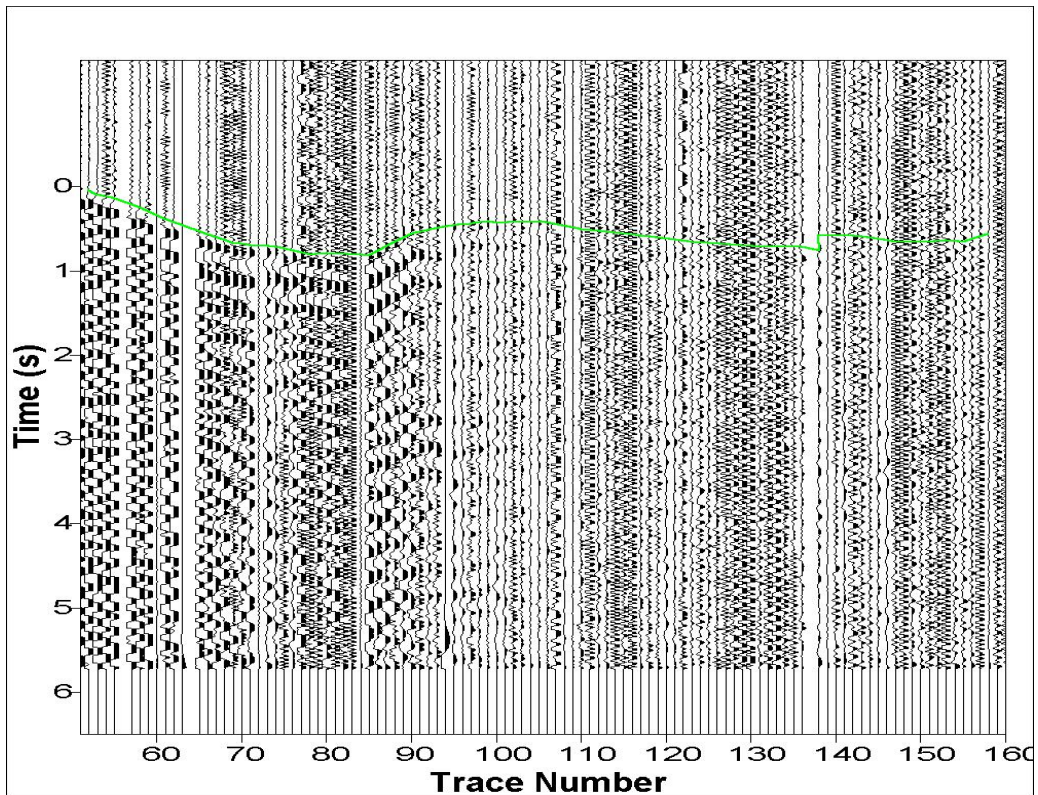
Table (5.1) shot points with their position and elevation.

5.1 DATA PROCESSING

The objective of a seismic refraction survey is to use the information contained in travel time data to determine the depth and configuration of refracting layers. Simply, the interpreter studies the travel time curves, assigns each arrival to a particular refractor, and determines the velocity of the refractor from the slope of the travel time curve. While there are similarities among the various methods of seismic refraction analysis, there are substantial differences in the field procedures, computations, and assumptions that each involves. For several reasons the GRM method is used to process the whole data in this study. Fig 6.1-Fig 6.6 show reduced arrival times for each geophone from each shot.



Fig(6.1) Traces from shot 21



Fig(6.2) Traces from shot

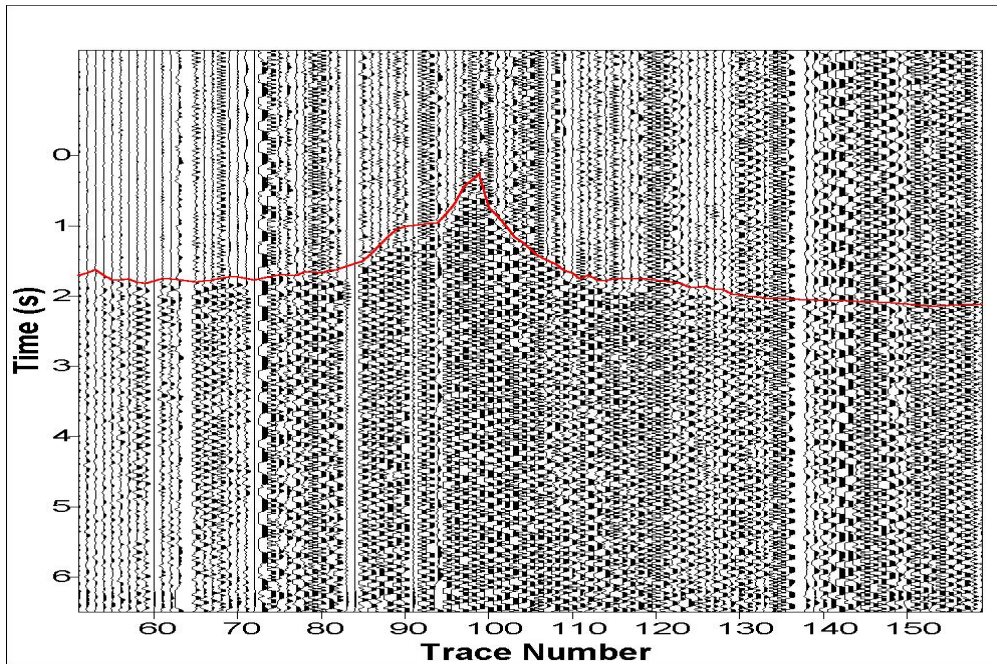


Fig (6.3)traces from shot 23

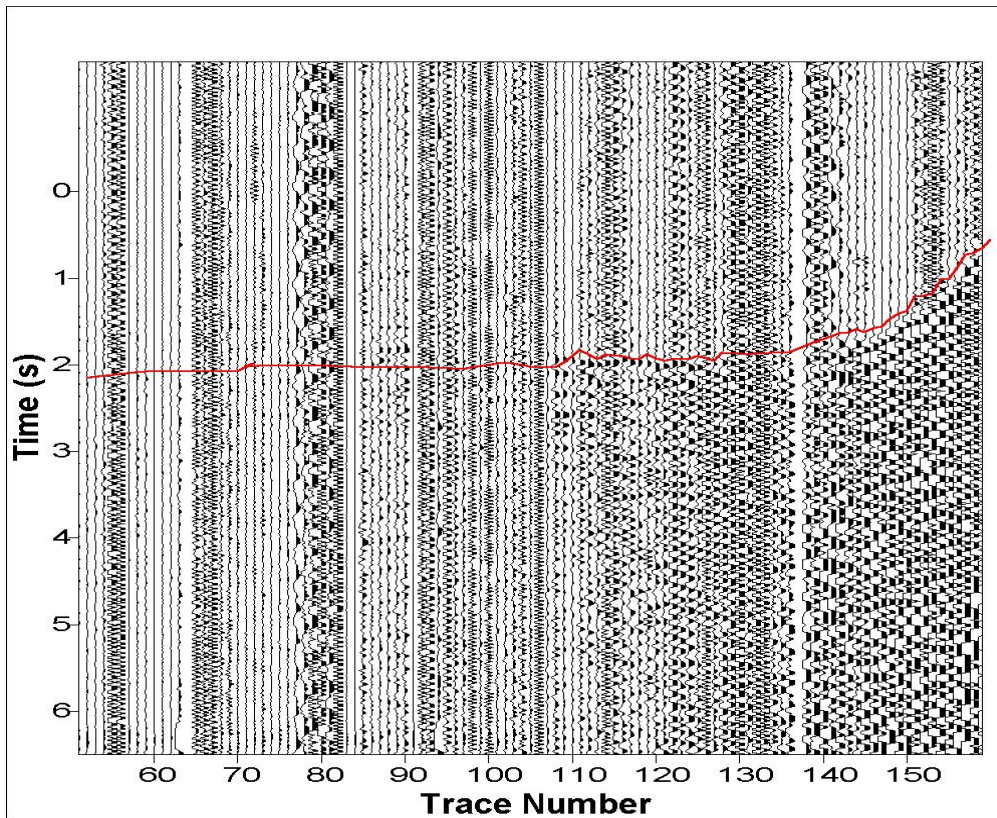


Fig (6.4)Traces from shot 24

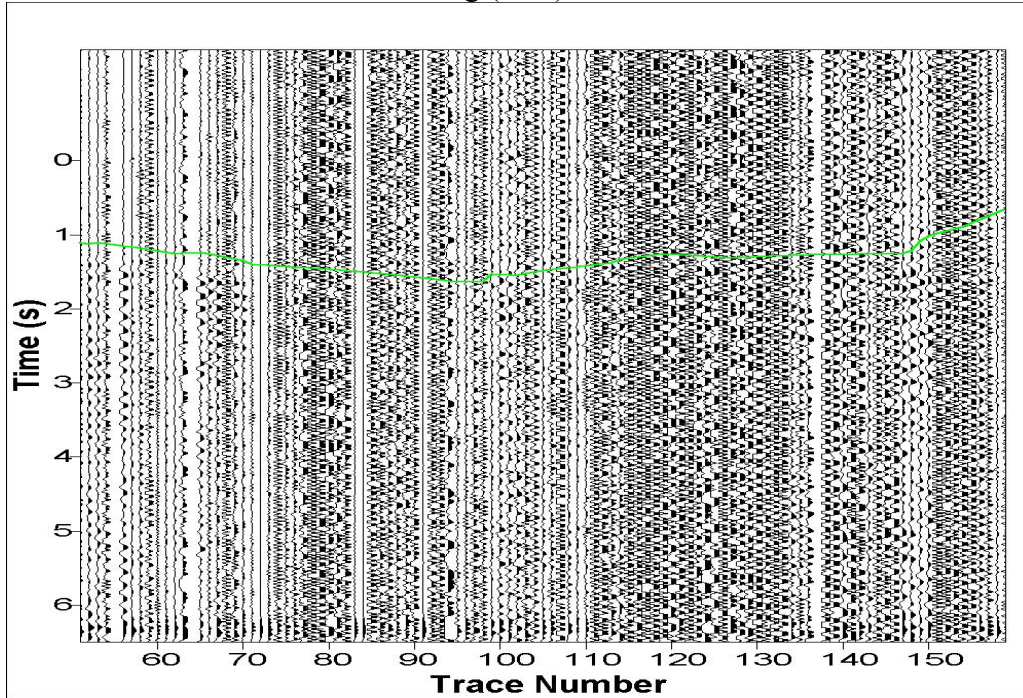


Fig (6.5)Traces from shot 25

6.1 Steps to be employed while processing the data are,

- i. Picking of first arrivals.
- ii. Setting the actual first arrival times.
- iii. Entering the data(actual arrival time of each geophone together with its elevation and geophone position) in to the inverting software.
- iv. Assigning layers to arrival times, on the travel time curves.
- v. Interpreting the assignments to produce the Geoseismic model.

6.1.1 Entering the data

The software type used for analyzing and interpreting the seismic data is GREMIX, which requires,

- ❖ Actual first arrival time
- ❖ Elevation
- ❖ And exact position of each geophone.

The raw data, which was reduced by assuming average crustal velocity of 6000m/s, should be converted and the actual first arrival time must be used while processing the data. The following equation is used, to convert the reduced arrival times in to their actual values.

$$T_t = T_{red} + \frac{X}{V} \text{ Where,}$$

T_t = the actual arrival time of the head waves.

T_{red} = the reduced arrival time as it is picked from the trace.

X = the offset of each geophone from the active shot.

V = crustal velocity = 6000m/s

The input data to the inverting software GREMIX adheres to the shot geophone nomenclature shown in the diagram below.(Fig 6.6).GREMIX uses the first-arrival times nominated from forward and reverse shots to calculate the time depth and velocity-analysis values needed for

the geophysical interpretation. These are plotted with the travel time curves. The calculations can be made for a range of XY values.

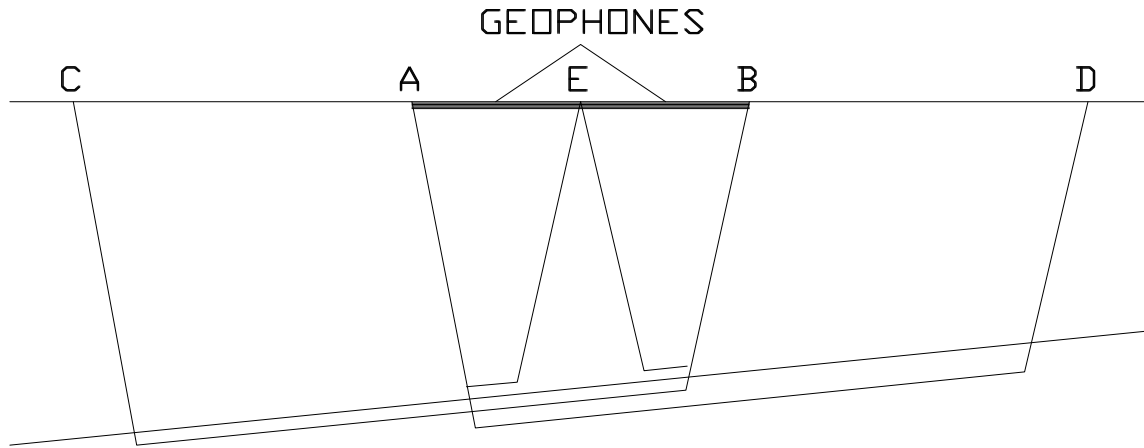


Fig (6.6) Sketch showing the nomenclature of GREMIX

GREMIX recognizes each shot; based on the offset data input it. In this study shot 21 is taken as a reference (with offset zero) to take measurements of offset distances for the remaining shots and sensors on the profile line. There fore, for this study all values that have been recognized as offsets are positive, as indicated in the diagram below. Based on the data input to GREMIX, the following association between the shot points and the letters in the GRIMIX nomenclature was found.

SP 21=C

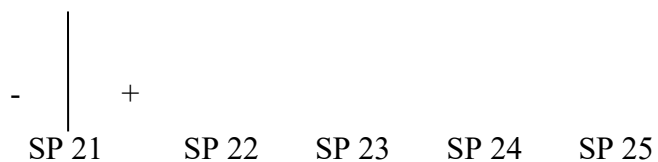
SP22 =A

SP23 =E

SP 24=B

SP 25=D

REFERENCE



Shot type	Shot number	Shot position	Elevation(m)
C	21	0	1706
A	22	26.93	1797
E	23	54.096	1653
B	24	88.399	1602
D	25	120.703	1236

Table 6.1 Input shot data

6.1.2 Data inversion

Following the input data, Gremix automatically give us the travel time curve Fig(6.9) after making a static correction based on the elevation parameters of each geophone.

- Figure showing recorded signals with varying degree of signal to noise ratio.

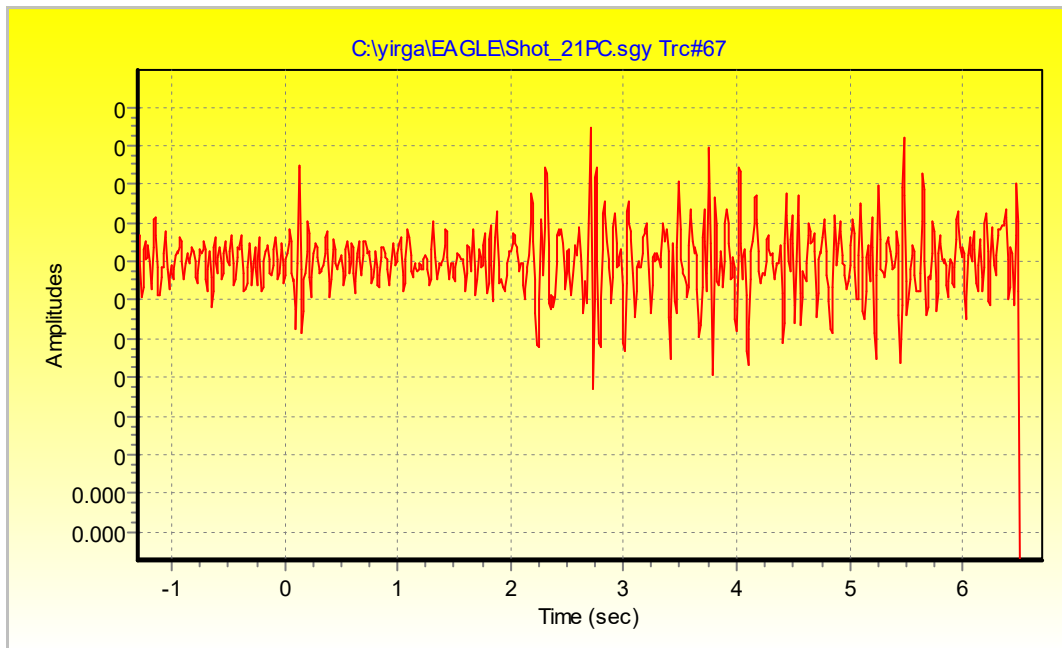


Fig (6.8) Low signal to noise ratio recorded from SP 21.

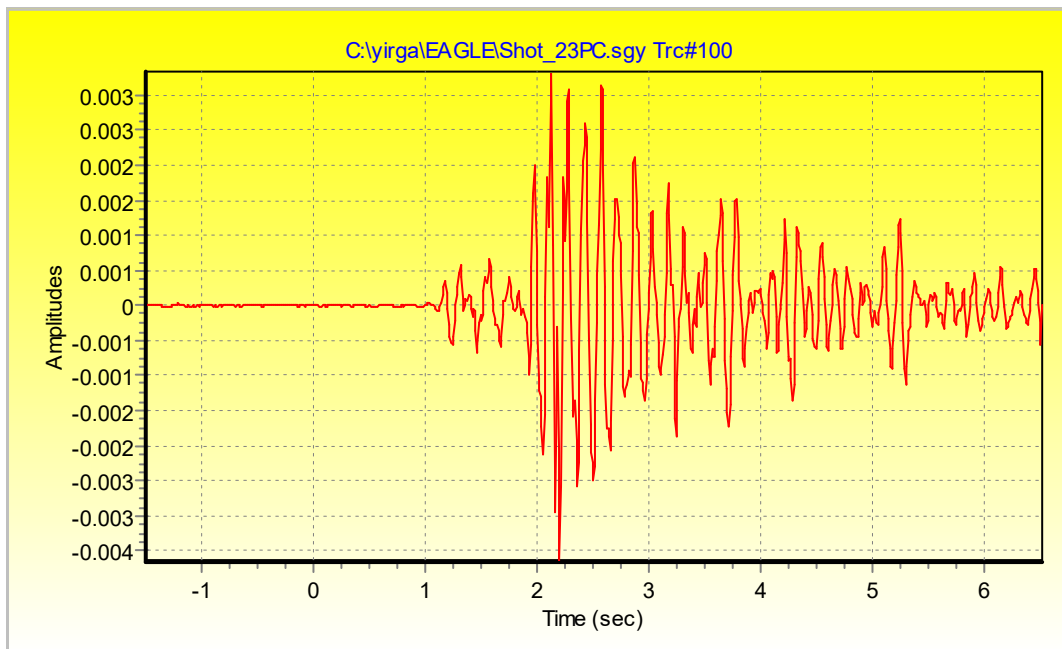


Fig (6.9) High signal to noise ratio recorded from shot point 23.

6.1.3 Assigning layers to arrivals on the travel time curve

Although processing is done by computer, the interpreter has to inspect the data and decide whether arrivals at different geophones are from head waves, which have traveled on the same refractor or different refractors. To calculate a valid velocity analysis or time-depth, we must be able to establish that the waves involved in the first arrivals at the two geophones and that the wave involved in the reciprocal time all traveled on the same refractor.

Basically we use the difference curves to assign layers to arrival times on the travel time curve. For curves with opposite slopes, difference curves show constant slope for areas in which arrivals come from the same refractor. Whereas for curves with slopes of the same sign, difference curves are flat (constant difference) for the same situation.

- The layer assignments for the arrival times on the travel time curves Fig(6.9) were,
 - One refractor each for the arrival times from the two far end shots(SP 21 and SP25)
 - Two refractors were identified for the middle shot on either side of the travel time curve (Sp 23)
 - Three refractors each for the arrival times for the end shots (SP 22 and SP 24).

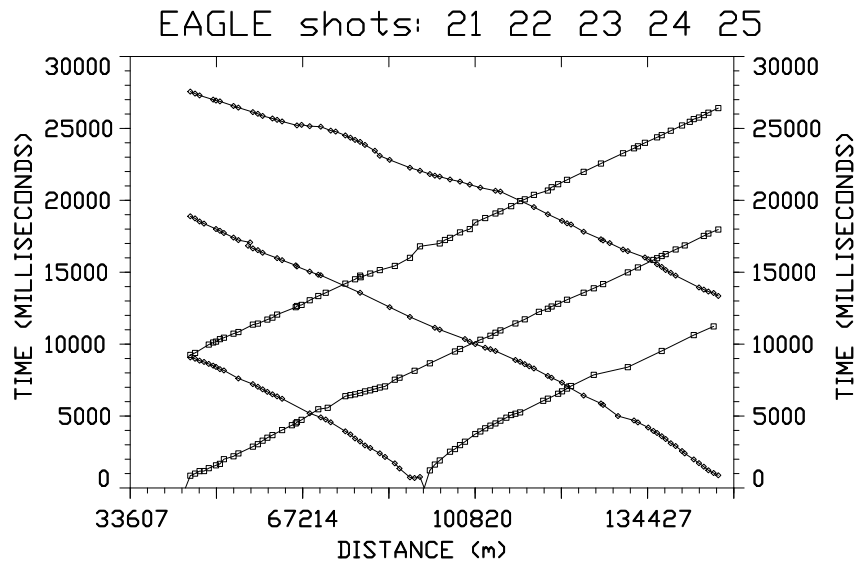


Fig (6.9) assigned arrival curves

- At the out set four layers are identified to a maximum depth of 10km.

7. Interpretation

At the outset, four layers of different acoustic property are identified.

The average seismic velocities of the layers from top to bottom are 2800m/s, 3950m/s, 5726m/s and 6320m/s respectively.

The uppermost layer with velocity 2800m/s has more or less constant thickness. The layers are characterized by undulating refractors may be due to continuous tectonic disturbance.

The low velocity section in the first layer indicates that the uppermost crust has been covered by weathered and fractured low-density material which include quaternary sediments, late Miocene, Early Pliocene, ignimbrites, rhyolites, Oligocene basalts, Mesozoic sediments, etc.

The average thickness of the second layer is 2800m. The layer has rugged topography. It is most likely that a series of Meta sediments and metavolcanics of low-grade metamorphic basement comprises the layer.

The average thickness of the third layer is 3500m. The velocity of this layer is high (5726m/s), Which corresponds to a high grade crystalline basement rock.

The seismic velocity in the fourth layer has a typical crustal velocity greater than 6300m/s. This probably is mafic and ultramafic rock.

Generally the rift shows higher velocities Fig (6.10) in the second, third and fourth layers. This could probably be the result of mafic intrusions beneath the rift magmatic centers.

8. Conclusions And Recommendations

8.1 Conclusions

The result of the research work reveals that there are four major seismic layers down to a depth of 10km along the main Ethiopian rift with average velocities of 2800m/s, 3950m/s, 5726.6 m/s and 6320 m/s respectively. The velocity model constrains velocity anomalies especially in the third layer, which probably is an indication of mafic lower crust.

8.2 Recommendations

Based on the result obtained from the study, the following points are recommended.

1. Other geophysical methods like the gravity method should be employed so as to have an integrated and refined result.
2. Amplitude studies of first arrivals should also be made to understand the attenuating behavior of the rocks..
- 3.** Reflection seismic methods must be employed to delineate the structures beneath The area with high resolution.

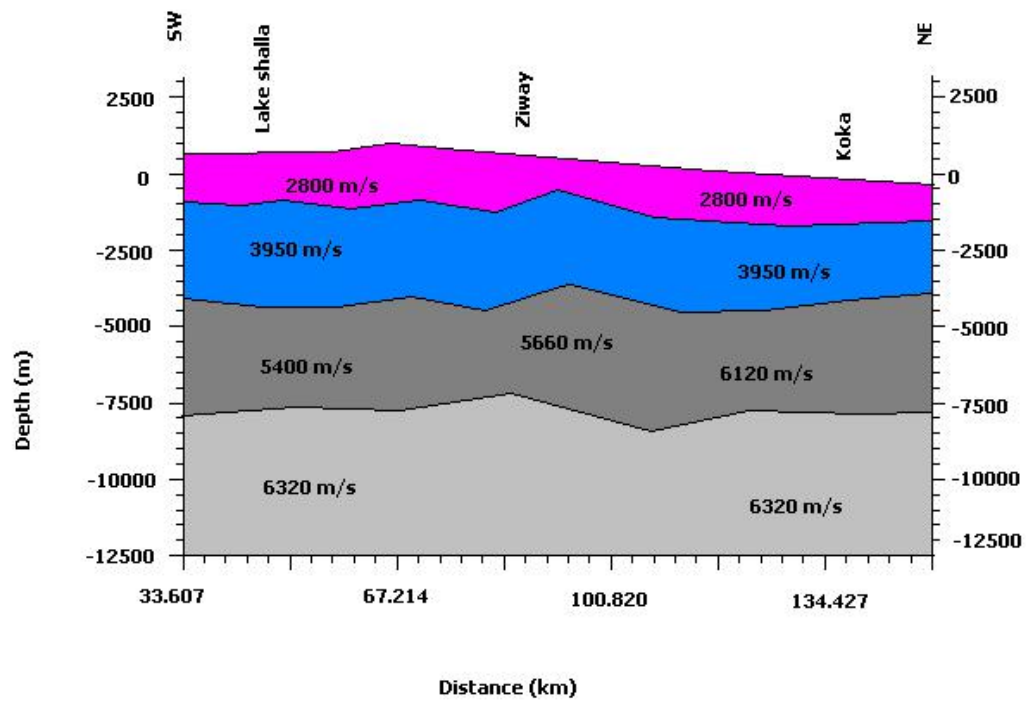
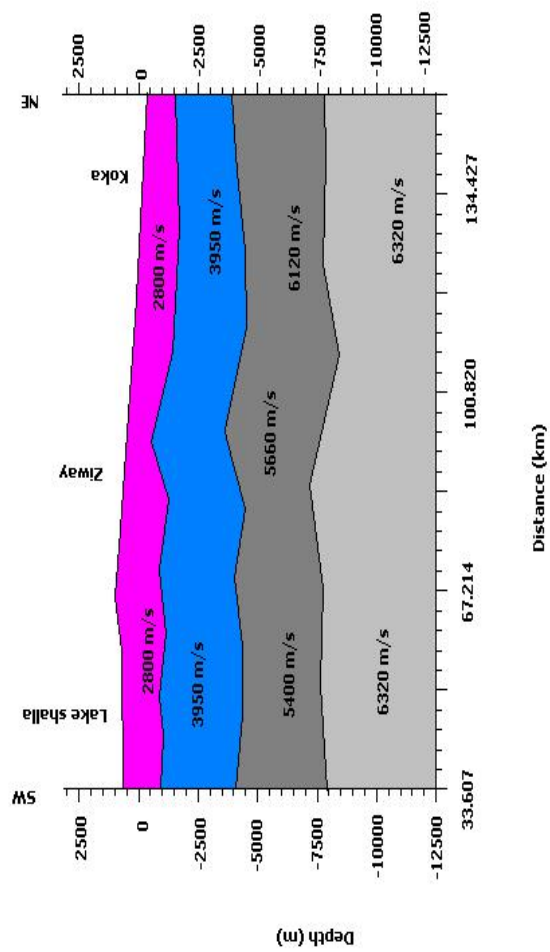


Fig 6.10 Seismic model along the main Ethiopian Rift



- Legend**
- | | |
|---|--|
| 1 | Quaternary sediments, cenozoic volcanic flows and Mesozoic sediments |
| 2 | Meta sediments and Meta volcanics |
| 3 | High grade Metamorphic rock |
| 4 | crustal rocks (mafic ultramafic rocks) |

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APPENDIX

APPENDIX 1 -GEOPHONE STATION, LOCATION AND ALTITUDE						
Order	Trace	Station	Easting	Northing	Elevation	Offset in (km) from shot 21
1	51	2051	464068	824000	1755.95	45.317
2	52	2052	464098	825018	1730.6	46.243
3	53	2053	464216	826014	1727.5	47.12
4	54	2054	464262	826989	1679.55	47.987
5	55	2055	464244	827986	1662.2	48.913
6	56	2056	464345	828980	1643	49.773
7	57	2057	464215	830011	1626.65	50.323
8	58	2058	463742	830862	1650.3	51.087
9	59	2059	463453	831854	1636.75	51.862
10	60	2060	463249	832802	1647.2	52.816
11	61	2061	463382	833788	1653.5	53.737
12	62	2062	463375	834801	1647.7	54.687
13	63	2063	463499	835786	1643.7	55.602
14	64	2064	463510	836776	1634.8	56.579
15	65	2065	463621	837791	1635.4	57.506
16	66	2066	463619	838794	1622.45	58.484
17	67	2067	463753	839798	1627.4	59.41
18	68	2068	463752	840795	1616.6	60.333
19	69	2069	463814	841766	1611.6	61.323
20	70	2070	463949	842774	1612.25	62.232
21	71	2071	463929	843755	1602.75	63.204
22	72	2072	464016	844766	1602.75	64.186
23	73	2073	464158	845760	1598.65	65.108
24	74	2074	464139	846752	1601.2	65.864
25	75	2075	463819	847672	1598.3	66.092
26	76	2076	463484	848030	1591.95	67.002
27	77	2077	462540	849313	1598.4	68.582
28	79	2079	462024	851138	1606.1	70.253
29	80	2080	461667	852101	1605.5	70.736
30	81	2081	461403	853070	1609.35	71.704
31	82	2082	460774	853751	1611.8	72.675
32	83	2083	460718	854771	1605.9	73.643
33	84	2084	460776	855764	1612.95	74.567
34	85	2085	460681	856792	1621	75.527
35	86	2086	460796	857716	1636.55	76.519

36	87	2087	461334	858566	1650.45	77.421
37	88	2088	461826	859459	1654.4	78.406
38	89	2089	462390	860239	1652.4	79.351
39	90	2090	462860	861129	1650	80.381
40	91	2091	463480	861935	1654.9	81.289
41	92	2092	463934	862876	1656	82.24
42	93	2093	464470	863667	1655	83.258
43	94	2094	465082	864475	1655.25	84.165
44	95	2095	465633	865372	1665.35	85.168
45	96	2096	466151	866161	1666.2	86.09
46	97	2097	466597	867074	1649.85	88.117
47	98	2098	467202	867851	1651.9	89.025
48	100	2100	467930	869746	1640.45	91.999
49	101	2101	467830	870727	1646.35	92.97
50	102	2102	468158	871725	1648.75	93.922
51	104	2104	468206	873719	1646.75	95.926
52	105	2105	468354	874690	1649.65	96.886
53	106	2106	468530	875630	1648.75	97.873
54	107	2107	468614	876667	1651.4	98.837
55	108	2108	468698	877669	1641.8	99.774
56	109	2109	468863	878621	1641.1	100.845
57	110	2110	468929	879629	1645.15	101.837
58	111	2111	469039	880601	1645.4	102.776
59	112	2112	469219	881524	1643.7	103.797
60	113	2113	469278	882621	1648.8	104.783
61	114	2114	469395	883618	1651.9	105.75201
62	115	2115	469500	884565	1656.6	106.90501
63	116	2116	469660	885578	1660.95	107.84001
64	117	2117	469685	886595	1666.4	108.67201
65	118	2118	469853	887553	1657.3	109.59301
66	119	2119	470036	888697	1668.25	110.49101
67	120	2120	470157	889633	1669.05	111.36501
68	121	2121	470769	890328	1670.4	112.27301
69	122	2122	471459	891089	1668.15	113.24601
70	123	2123	472091	891843	1669.9	114.08801
71	124	2124	472787	892551	1668.5	114.99401
72	125	2125	473460	893300	1671.1	115.73101
73	126	2126	474150	894107	1662.85	117.00001
74	127	2127	474818	894785	1655.95	117.75001
75	128	2128	475439	895542	1655.5	118.72801

76	129	2129	475951	896154	1665.1	119.51101
77	130	2130	476794	897218	1651.55	120.44901
78	131	2131	477625	897741	1652.3	121.97101
79	132	2132	478244	898570	1662.95	122.90201
80	133	2133	478811	899204	1665.75	123.94001
81	134	2134	479453	899982	1665.6	125.33901
82	136	2136	481045	901052	1667.15	127.04201
83	137	2137	481696	901813	1665.15	127.69201
84	138	2138	482258	902715	1661.25	128.67101
85	139	2139	483302	903830	1665.25	129.63801
86	140	2140	483353	904253	1664.75	130.57301
87	141	2141	484307	905280	1683.1	131.77801
88	142	2142	484710	905824	1689.85	132.53901
89	143	2143	485121	906718	1691.1	133.90701
90	144	2144	485964	907439	1691.4	134.56101
91	145	2145	486452	908257	1686.75	135.55501
92	146	2146	487016	909333	1689.25	136.27801
93	147	2147	487385	910005	1692.4	137.18101
94	148	2148	487847	911291	1694.55	137.95401
95	149	2149	488366	911796	1692.4	138.94701
96	150	2150	489235	912531	1695.4	139.89601
97	151	2151	489817	913089	1690.35	141.13701
98	152	2152	490652	913735	1681.95	141.66401
99	153	2153	491221	914343	1681.6	142.65101
100	154	2154	492216	915023	1674.5	143.42101
101	155	2155	492782	915816	1661.65	144.44101
102	156	2156	493828	916735	1656	145.38201
103	157	2157	494333	917104	1650.95	146.27601
104	158	2158	495156	917835	1639.25	147.29901
105	159	2159	495756	918419	1627.5	148.20701
106	160	2160	496415	919255	1607.4	149.19201

APPENDIX 2-OFFSET OF EACH GEOPHONE FROM EACH SHOT					
Station	OFFSET FROM SP21	Offset from SP 22 (km)	Offset from SP 23 (km)	Offset from sp 24 (km)	Offset from SP 25 (km)
2051	45.317	6.594	-45.993	-102.35	-157.45
2052	46.243	6.699	-44.99	-101.37	-156.62
2053	47.12	6.869	-44.014	-100.44	-155.86
2054	47.987	7.115	-43.025	-99.522	-155.13
2055	48.913	7.593	-42.026	-98.56	-154.33
2056	49.773	8	-41.016	-97.646	-153.65
2057	50.323	8.129	-40.235	-97.035	-153.34
2058	51.087	8.592	-39.297	-96.233	-152.82
2059	51.862	9.177	-38.392	-95.437	-152.27
2060	52.816	10.045	-37.395	-94.476	-151.46
2061	53.737	10.88	-36.396	-93.548	-150.74
2062	54.687	11.784	-35.405	-92.597	-149.95
2063	55.602	12.655	-34.426	-91.684	-149.23
2064	56.579	13.607	-33.406	-90.711	-148.43
2065	57.506	14.512	-32.417	-89.795	-147.72
2066	58.484	15.482	-31.406	-88.826	-146.92
2067	59.41	16.401	-30.424	-87.918	-146.21
2068	60.333	17.325	-29.458	-87.013	-145.5
2069	61.323	18.318	-28.442	-86.039	-144.69
2070	62.232	19.234	-27.482	-85.16	-144.03
2071	63.204	20.214	-26.476	-84.215	-143.27
2072	64.186	21.204	-25.473	-83.256	-142.49
2073	65.108	22.142	-24.507	-82.375	-141.83
2074	65.864	22.933	-23.681	-81.694	-141.44
2075	66.092	23.192	-23.413	-81.523	-141.44
2076	67.002	24.221	-22.442	-80.824	-141.29
2077	68.582	25.915	-20.887	-79.468	-140.47
2079	70.253	27.728	-19.334	-78.091	-139.68
2080	70.736	28.325	-18.99	-77.82	-139.72
2081	71.704	29.333	-18.111	-76.973	-139.11
2082	72.675	30.326	-17.218	-76.097	-138.44
2083	73.643	31.339	-16.385	-75.274	-137.87
2084	74.567	32.271	-15.542	-74.427	-137.2
2085	75.527	33.174	-14.54	-73.424	-136.26

2086	76.519	34.12	-13.521	-72.404	-135.31
2087	77.421	34.969	-12.559	-71.443	-134.39
2088	78.406	35.917	-11.556	-70.437	-133.46
2089	79.351	36.81	-10.541	-69.425	-132.48
2090	80.381	37.813	-9.508	-68.388	-131.55
2091	81.289	38.686	-8.556	-67.435	-130.64
2092	82.24	39.594	-7.539	-66.421	-129.66
2093	83.258	40.582	-6.487	-65.368	-128.68
2094	84.165	41.462	-5.542	-64.422	-127.79
2095	85.168	42.448	-4.544	-63.414	-126.89
2096	86.09	43.339	-3.558	-62.433	-125.93
2097	88.117	45.351	-1.703	-60.441	-124.22
2098	89.025	46.285	-1.521	-59.672	-123.72
2100	91.999	49.288	3.246	-56.977	-121.68
2101	92.97	50.268	4.132	-56.098	-121.01
2102	93.922	51.224	5.018	-55.228	-120.34
2104	95.926	52.255	6.033	-54.34	-119.69
2105	96.886	53.254	7.021	-53.479	-119.06
2106	97.873	54.22	7.958	-52.617	-118.4
2107	98.837	55.22	8.961	-51.776	-117.8
2108	99.774	56.195	9.928	-50.94	-117.18
2109	100.845	57.135	10.848	-50.107	-116.54
2110	101.837	58.222	11.945	-49.22	-115.92
2111	102.776	59.223	12.942	-48.376	-115.3
2112	103.797	60.171	13.89	-47.59	-114.72
2113	104.783	61.198	14.907	-46.717	-114.07
2114	105.752	62.2	15.924	-45.943	-113.54
2115	106.905	63.173	16.887	-45.125	-112.92
2116	107.84	64.332	18.037	-44.164	-112.19
2117	108.672	65.274	18.978	-43.406	-111.63
2118	109.593	66.076	19.707	-42.485	-110.77
2119	110.491	66.963	20.526	-41.455	-109.79
2120	111.365	67.833	21.348	-40.476	-108.88
2121	112.273	68.677	22.148	-39.483	-107.94
2122	113.246	69.558	22.999	-38.478	-106.99
2123	114.088	70.505	23.924	-37.415	-106
2124	114.994	71.323	24.727	-36.467	-105.1
2125	115.731	72.21	25.607	-35.488	-104.2
2126	117	72.931	26.326	-34.695	-103.46
2127	117.75	74.177	27.574	-33.344	-102.23

2128	118.728	74.9	28.31	-32.387	-101.25
2129	119.511	75.865	29.285	-31.365	-100.33
2130	120.449	76.635	30.067	-30.51	-99.533
2131	121.971	77.56	31.008	-29.51	-98.616
2132	122.902	79.046	32.559	-27.627	-96.72
2133	123.94	79.967	33.504	-26.629	-95.804
2134	125.339	80.999	34.552	-25.594	-94.908
2136	127.042	82.802	36.394	-23.733	-93.26
2137	127.692	84.077	37.71	-22.338	-91.971
2138	128.671	84.723	38.372	-21.668	-91.375
2139	129.638	85.701	39.355	-20.763	-90.64
2140	130.573	86.659	40.358	-19.65	-89.573
2141	131.778	87.591	41.304	-18.744	-88.804
2142	132.539	88.794	42.518	-17.617	-87.866
2143	133.907	89.553	43.286	-16.901	-87.263
2144	134.561	90.921	44.648	-15.75	-86.352
2145	135.555	91.571	45.324	-15.033	-85.684
2146	136.278	92.559	46.36	-13.897	-84.605
2147	137.181	93.279	47.11	-13.108	-83.87
2148	137.954	94.178	48.059	-12.05	-82.856
2149	138.947	94.949	48.859	-11.234	-82.105
2150	139.896	95.94	49.914	-10.032	-80.938
2151	141.137	96.888	50.883	-9.121	-80.127
2152	141.664	98.128	52.185	-7.751	-78.83
2153	142.651	98.655	52.744	-7.137	-78.239
2154	143.421	99.643	53.781	-6.063	-77.217
2155	144.441	100.414	54.585	-5.264	-76.454
2156	145.382	101.435	55.632	-4.355	-75.571
2157	146.276	102.378	56.597	-3.595	-74.763
2158	147.299	103.272	57.521	-2.927	-73.949
2159	148.207	104.298	58.578	-2.418	-73.032
2160	149.192	105.208	59.521	2.285	-72.193

APPENDIX 3-FIRST ARRIVAL RECORDS FOR EACH
GEOPHONE FROM EACH SHOT

Trace	First arrival from SP 21	First arrival from SP 22	First arrival from SP 23	First arrival from SP 24	First arrival from SP 25
51	1.697	0.009	1.409	18.8907	1.318
52	1.682	0.026	1.485	18.7435	1.318
53		0.06	1.515	18.5287	1.318
54		0.102	1.591	18.379	
55	1.818	0.119	1.621		
56	1.818		1.667		1.394
57	1.788	0.264	1.712	18.0055	1.379
58	1.833	0.213	1.712	17.8798	1.409
59	1.803	0.29	1.773	17.7272	
60					
61	1.773	0.375		17.4093	1.439
62	1.727	0.366	1.712	17.2328	1.455
63				17.0737	
64				16.8405	
65	1.773	0.554	1.818	16.6568	1.515
66	1.682	0.563	1.803	16.5363	1.53
67		0.554	1.788	16.354	1.5
68	1.652	0.605	1.773		
69	1.621	0.622	1.773		1.576
70	1.682		1.788	15.9813	1.591
71		0.656	1.788	15.8388	1.606
72					
73		0.673			
74	1.591	0.665		15.4797	
75	1.636	0.707		13.5872	1.636
76	1.561	0.699			1.701
77	1.621		1.712	15.0457	1.744
78			1.727	14.9332	
79	1.636	0.852		14.8262	
80		0.852	1.742	14.791	1.833
81	1.621		1.742		
82			1.697		1.773
83					1.801

84					
85	1.621	0.852	1.515		1.803
86		0.767	1.47		1.803
87	1.606	0.682	1.348		1.803
88	1.697	0.614	1.303	13.5725	1.818
89	1.727	0.58	1.197		1.773
90	1.712	0.486	1.197		
91		0.435			1.672
92	1.742	0.384	1.152		1.485
93		0.298	1.076		
94				12.57	1.515
95	1.742	0.452	0.955		
96		0.443	0.758		
97	1.773		0.455	11.8915	1.561
98		0.435	0.439		
99	1.788		0.5		1.576
100		0.452	0.682		1.545
101			0.939	11.1377	1.545
102	1.742		1.076	11.0077	1.591
103	1.818		1.212		
104	1.803		1.348		1.606
105		0.452	1.379		
106	1.758	0.46	1.485		1.667
107			1.561	10.354	
108	1.621			10.1692	1.667
109	1.652		1.758	10.0063	
110		0.435	1.773		1.667
111	1.636		1.833	9.74967	
112		0.384	1.848	9.65017	
113	1.621	0.426	1.833	9.52117	1.742
114	1.606	0.426	1.864		1.788
115			1.879		
116	1.621		1.909		
117		0.426	1.879	8.89883	
118	1.682		1.833	8.77317	
119	1.652	0.418		8.61	
120				8.4595	
121	1.667			8.307	1.697
122		0.494			
123			1.939		

124	1.515	0.426	1.939	7.79367	1.667
125	1.621	0.46		7.6615	
126	1.621	0.443	1.955		
127			2	7.32183	1.697
128	1.636	0.443	2.03	7.0755	1.682
129			2.076		1.727
130					
131	1.652	0.401		6.4225	1.697
132					
133		0.384	2.106		
134	1.667			5.87483	1.712
135		0.375		5.7735	1.667
136				5.541	1.697
137					
138				5.2485	
139	1.667				1.652
140		0.392	1.515		1.667
141	1.654			4.69417	
142	1.667	0.409		4.55883	
143	1.682				1.636
144				4.2025	1.621
145		0.426		3.96817	
146	1.667	0.429		3.82067	1.576
147	1.667	0.431	1.515	3.59933	1.561
148		0.433		3.40233	1.47
149	1.682			3.111	1.47
150		0.435		2.91417	1.409
151	1.682			2.56483	
152		0.426		2.4015	
153	1.682				
154	1.742		1.53	1.98333	
155	1.682			1.72583	1.348
156	1.697	0.46		1.47817	1.333
157	1.712	0.477		1.22983	1.333
158			1.47	1.039	1.379
159	1.712	0.435	1.424	0.88083	1.318

APPENDIX – 4 FIRST ARRIVAL DATA AS OBTAINED IN THE RAW
FORMAT (Example taken from SP 22)

Trace No	First arrival
51.029	0.009
51.975	0.026
52.987	0.06
54.015	0.102
54.994	0.119
56.986	0.264
57.998	0.213
59.043	0.29
61.034	0.375
61.98	0.366
64.991	0.554
65.987	0.563
66.982	0.554
68.066	0.605
69.04	0.622
70.942	0.656
72.978	0.673
73.996	0.665
75.013	0.707
75.987	0.699
79.031	0.852
80.004	0.852
85.049	0.852
86.044	0.767
87.04	0.682
88.035	0.614
89.053	0.58
90.004	0.486
91	0.435
92.018	0.384
93.058	0.298
95.035	0.452
96.031	0.443

98	0.435
100.013	0.452
105.035	0.452
106.031	0.46
110.035	0.435
111.982	0.384
112.044	0.443
113.018	0.426
113.991	0.426
116.978	0.426
118.969	0.418
122	0.494
123.969	0.426
125.009	0.46
126.004	0.443
128.04	0.443
131.027	0.452
131.032	0.401
133.021	0.384
135.01	0.375
140.046	0.392
141.971	0.409
144.955	0.426
150.055	0.435
152.044	0.426
154.964	0.46
156.054	0.477
157.049	0.46
159.07	0.435

APPENDIX 5 -ACTUAL X-T DATA THAT WAS INVERTED TO GIVE THE GEOSEISMIC MODEL								
Trace No	Station	Elevation	Offset from Sp21	Actual time from SP 21	Actual time from SP 22	Actual time from SP 23	Actual time from SP 24	Actual time from SP 25
51	2051	1755.95	45.317	9.24983	1.108	9.0745	18.8907	27.5595
52	2052	1730.6	46.243	9.38917	1.1425	8.9833333	18.7435	27.4217
53	2053	1727.5	47.12		1.20483	8.8506667	18.5287	27.2948
54	2054	1679.55	47.987		1.28783	8.7618333	18.379	
55	2055	1662.2	48.913	9.97017	1.3845	8.6253333		
56	2056	1643	49.773	10.1135		8.503		27.002
57	2057	1626.65	50.323	10.1752	1.61883	8.4178333	18.0055	26.9363
58	2058	1650.3	51.087	10.3475	1.645	8.2615	17.8798	26.8793
59	2059	1636.75	51.862	10.4467	1.8195	8.1716667	17.7272	
60	2060	1647.2	52.816					
61	2061	1653.5	53.737	10.7292	2.18833		17.4093	26.5622
62	2062	1647.7	54.687	10.8415	2.33	7.6128333	17.2328	26.4463
63	2063	1643.7	55.602				17.0737	
64	2064	1634.8	56.579				16.8405	
65	2065	1635.4	57.506	11.3573	2.97267	7.2208333	16.6568	26.1347
66	2066	1622.45	58.484	11.4293	3.14333	7.0373333	16.5363	26.016
67	2067	1627.4	59.41		3.2875	6.8586667	16.354	25.869
68	2068	1616.6	60.333	11.7075	3.4925	6.6826667		
69	2069	1611.6	61.323	11.8415	3.675	6.5133333		25.6912
70	2070	1612.25	62.232	12.054		6.3683333	15.9813	25.5952
71	2071	1602.75	63.204		4.025	6.2006667	15.8388	25.4848
72	2072	1602.75	64.186					
73	2073	1598.65	65.108		4.36333			
74	2074	1601.2	65.864	12.5683	4.48717		15.4797	
75	2075	1598.3	66.092	12.6513	4.57233		13.5872	25.209
76	2076	1591.95	67.002	12.728	4.73583			25.2485
77	2077	1598.4	68.582	13.0513		5.1931667	15.0457	25.1555
78			69.401				14.9332	
79	2079	1606.1	70.253	13.3448	5.47333		14.8262	
80	2080	1605.5	70.736		5.57283	4.907	14.791	25.1202
81	2081	1609.35	71.704	13.5717		4.7605		
82	2082	1611.8	72.675			4.5666667		24.8467
83	2083	1605.9	73.643					24.7795
84	2084	1612.95	74.567					

85	2085	1621	75.527	14.2088	6.381	3.9383333		24.5122
86	2086	1636.55	76.519		6.45367	3.7235		24.3548
87	2087	1650.45	77.421	14.5095	6.51017	3.4411667		24.2005
88	2088	1654.4	78.406	14.7647	6.60017	3.229	13.5725	24.0618
89	2089	1652.4	79.351	14.9522	6.715	2.9538333		23.853
90	2090	1650	80.381	15.1088	6.78817	2.7816667		
91	2091	1654.9	81.289		6.88267			23.4458
92	2092	1656	82.24	15.4487	6.983	2.4085		23.0952
93	2093	1655	83.258		7.06167	2.1571667		
94	2094	1655.25	84.165				12.57	22.8128
95	2095	1665.35	85.168	15.9367	7.52667	1.7123333		
96	2096	1666.2	86.09		7.66617	1.351		
97	2097	1649.85	88.117	16.4592		0.7388333	11.8915	22.264
98	2098	1651.9	89.025		8.14917	0.6925		
99			90.075	16.8005		0.7651667		22.0555
100	2100	1640.45	91.999		8.66667	1.223		21.8253
101	2101	1646.35	92.97			1.6276667	11.1377	21.7137
102	2102	1648.75	93.922	17.3957		1.9123333	11.0077	21.6468
103			94.938	17.641				
104	2104	1646.75	95.926	17.7907		2.5181667		21.4497
105	2105	1649.65	96.886		9.48867	2.7053333		
106	2106	1648.75	97.873	18.0702	9.66333	2.9785		21.3002
107	2107	1651.4	98.837			3.2156667	10.354	
108	2108	1641.8	99.774	18.25			10.1692	21.0907
109	2109	1641.1	100.845	18.4595		3.7488333	10.0063	
110	2110	1645.15	101.837		10.3055	3.93		20.8828
111	2111	1645.4	102.776	18.7653		4.148	9.74967	
112	2112	1643.7	103.797		10.5837	4.3325	9.65017	
113	2113	1648.8	104.783	19.0848	10.7927	4.487	9.52117	20.6645
114	2114	1651.9	105.752	19.2313	10.9548	4.6785		20.6073
115	2115	1656.6	106.905			4.8851667		
116	2116	1660.95	107.84	19.5943		5.072		
117	2117	1666.4	108.672		11.4387	5.1635	8.89883	
118	2118	1657.3	109.593	19.9475		5.254	8.77317	
119	2119	1668.25	110.491	20.0672	11.7235		8.61	
120	2120	1669.05	111.365				8.4595	
121	2121	1670.4	112.273	20.3792			8.307	19.529
122	2122	1668.15	113.246		12.2448			
123	2123	1669.9	114.088			6.0601667		
124	2124	1668.5	114.994	20.6807	12.461	6.2068333	7.79367	19.0328

125	2125	1671.1	115.731	20.9095	12.6152		7.6615	
126	2126	1662.85	117	21.121	12.8058	6.5506667		
127	2127	1655.95	117.75			6.7183333	7.32183	18.5722
128	2128	1655.5	118.728	21.424	13.0872	6.9108333	7.0755	18.4033
129	2129	1665.1	119.511			7.0871667		18.3158
130	2130	1651.55	120.449					
131	2131	1652.3	121.971	21.9805	13.5753		6.4225	17.817
132	2132	1662.95	122.902					
133	2133	1665.75	123.94		13.8838	7.8646667		
134	2134	1665.6	125.339	22.5568			5.87483	17.293
135			125.756		14.1753		5.7735	17.2103
136	2136	1667.15	127.042				5.541	17.0255
137	2137	1665.15	127.692					
138	2138	1661.25	128.671				5.2485	
139	2139	1665.25	129.638	23.2733				16.5808
140	2140	1664.75	130.573		14.9905	8.399		16.4677
141	2141	1683.1	131.778	23.617			4.69417	
142	2142	1689.85	132.539	23.7568	15.3345		4.55883	
143	2143	1691.1	133.907	23.9998				16.028
144	2144	1691.4	134.561				4.2025	15.9017
145	2145	1686.75	135.555		15.8525		3.96817	
146	2146	1689.25	136.278	24.38	15.9755		3.82067	15.5543
147	2147	1692.4	137.181	24.5305	16.1273	9.5248333	3.59933	15.3703
148	2148	1694.55	137.954		16.2578		3.40233	15.1542
149	2149	1692.4	138.947	24.8398			3.111	14.9597
150	2150	1695.4	139.896		16.583		2.91417	14.7635
151	2151	1690.35	141.137	25.2048			2.56483	
152	2152	1681.95	141.664		16.8685		2.4015	
153	2153	1681.6	142.651	25.4572				
154	2154	1674.5	143.421	25.6455		10.6275	1.98333	
155	2155	1661.65	144.441	25.7555			1.72583	13.9432
156	2156	1656	145.382	25.9273	17.523		1.47817	13.7935
157	2157	1650.95	146.276	26.0913	17.689		1.22983	13.6578
158	2158	1639.25	147.299			11.233	1.039	13.551
159	2159	1627.5	148.207	26.4132	17.9697	11.344167	0.88083	13.3502

APPENDIX 5

P-wave velocities of different materials

Earth materials	V_p (m/s)
Water	1450-1530
Petroleum	1300-1400
Loess	300-600
Soil	100-500
Snow	350-3000
Solid glacier ice	3000-4000
Sand loose	200-2000
Sand (dry, loose)	200-1000
Sand (water saturated loose)	1500-2000
Glacial moraine	1500-2700
Sand and gravel (near surface)	400-2300
Sand and gravel at 2km depth	3000-3500
Clay	1000-2500
Estuarine mud's/clay	300-1800
Flood plain alluvium	1800-2200
Permafrost quaternary sediments	1500-4900
Limestone (soft)	1400-4500
Limestone (hard)	2800-7000
Dolomites	2500-6500
Anhydrite	3500-5500
Rock salt	4000-5500
Gypsum	2000-3500

Shale	2000-4100
Granites	4600-5200
Basalts	5500-7000
Peridotite	7800-8400
Serpentite	5500-6500
Gneiss	3500-7600
Marbles	3780-7000
Sulphide ores	3950-6700
Pulverised fuel ash	600-1000
Landfill refuse	400-750
Concrete	3000-3500
Disturbed soil	180-380

