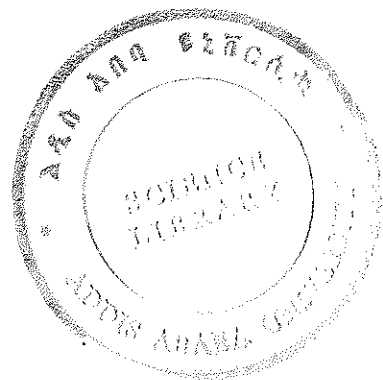


THEORETICAL & EXPERIMENTAL ANALYSIS OF THE
DIRECT & GLOBAL SOLAR RADIATION
OF ADDIS ABABA REGION

B Y

YOSEPH TELAHUN



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A B S T R A C T

The thesis submitted is divided into three parts; theoretical, computational, and experimental.

In the first part, general theory of attenuation of solar radiation in the atmosphere due to molecular scattering, water vapour and aerosol absorption is presented. The theoretical consideration includes integral turbidity factor and its relation to different absorption mechanisms. The calculated value due to the above absorption mechanisms and the integral turbidity factors for Addis Ababa region, are shown in Chapter Two.

In the second part of the thesis different empirical formulas relating global solar radiation, geographical and meteorological parameters of a given locality are collected. The most general formula by Sayigh et.al. [29] has been chosen and applied to Addis Ababa region. With constants given by Sayigh et. al. which are claimed to be universal, considerable discrepancies were obtained. Because of this reason, an effort was made to determine a new set of nine constants of the Sayigh's equation which are applicable to Addis Ababa region.

In this process the Gauss - Newton method of least squares for non-linear regression modified by H.O.Hartly [30] was used. The regression was made using one year data of the global solar radiation supplied in Chapter 3.

In the third experimental part, a Pyranometer based on the thermoelectric principle was constructed. The instrument is simple and in-expensive. The description and comparison measurements between the experimental and local meteorological station Pyranometers were taken, the results obtained are reported in Chapter 4.

CHAPTER 1

I N T R O D U C T I O N

1.1 The Solar Energy

The Earth and its atmosphere receives continuously 1.7×10^{17} W of radiation from the sun. A world population of 10 milliard with a total power need per person of 10 KW would require about 10^{11} KW. Taking only the irradiance on only 1% of the Earth's surface could be converted into useful energy with 10 percent efficiency, solar energy could provide the energy needs of all the people on the Earth.

Solar energy is the world's most abundant permanent source of energy. The amount of solar energy intercepted by the planet Earth is 5000 times greater than the sum of all other inputs (terrestrial, nuclear, geothermal and gravitational energies,) of this 30% is reflected to space, 41% is converted to low temperature heat and radiation to space and 23% powers the evaporation.

Total terrestrial radiation is only about one-third of the extraterrestrial during a year and 70% of that falls on the oceans. However, the remaining 5.4×10^{19} J that falls on land is a prodigious amount of energy about 5000 times the total energy usage of the world in 1978. However, only a small fraction of this total can be used because of physical and socio-economic constraints.

1.2 Solar Radiation

When we speak about radiation, reference is made to the electromagnetic emission directly from the sun's disc. This radiation reaches the earth about eight minutes after the emission process.

The sun's electromagnetic radiation covers an extremely wide range of wavelengths; from ($10^{-4}\mu - 10^{10}\mu$). Of the whole spectrum, we can sense only those that have approximately 0.1 to 100 microns. Within this limit, cause our body to heat up. Hence these waves are called "thermal radiation". Visible light ray occupy a very narrow band of the thermal radiation spectrum.

1.3 Solar Radiation at the Earth's Surface

From the point of view of utilization of solar energy we are more interested in the energy received at the Earth's surface than in the extraterrestrial energy. Solar energy received at the surface of the Earth is entirely different due to various reasons. Before studying this, it is important to know the following terms.

The Solar Radiation reaches the ground in two forms,

a) Direct Solar Radiation

Solar Radiation intercepted by the surface of the

Earth with negligible direction change and scattering in the atmosphere. Direct radiation is also referred to as beam radiation.

b) Diffuse Radiation

- Solar radiation scattered by aerosols, dust, and by Rayleigh mechanism, and does not have unique direction.

c) Global Radiation

- Global radiation is the total of direct and diffuse radiation, global radiation is sometimes referred to as total radiation.

1.4 The Solar Constant

Referring to Figure 1.1, the geometry of sun-Earth relationships, eccentricity of the Earth's orbit is such that the distance between the Earth and the sun varies $\pm 3\%$. The characteristics of the sun and its spatial relationship to the Earth result in a nearly fixed intensity of solar radiation outside of the Earth's atmosphere. The "Solar Constant", I_0 , is the energy from the sun, per unit time, received on a unit area of surface perpendicular to the radiation, in space, at the Earth's mean distance from the sun.

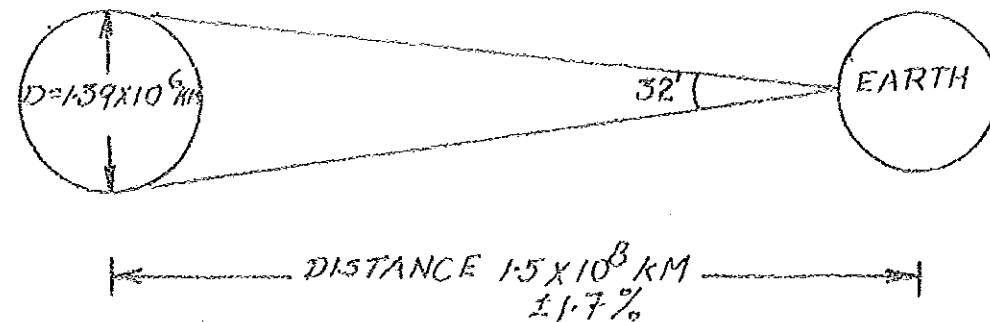


Fig. 1.1 Schematic of sun-Earth relationships (not to scale.)

Until recently, estimates of solar constant had to be made from ground based measurements of solar radiation after it had been transmitted through the atmosphere, and thus in part absorbed and scattered by components of the atmosphere. Extrapolations from terrestrial measurements, which were made from high mountains, had to be based on estimates of atmospheric transmission in various portion of solar spectrum. Pioneering studies were done by C.G. Abbot and his colleagues at Smithsonian institute.

However recently, the availability of very high altitude aircrafts, balloons and space crafts has permitted direct measurement of solar intensity ^{on} most or all of the Earth's atmosphere. These measurements have been reviewed and summarized and a new standard value of solar constant is given by Thekaekara [1]. And more recently rocket and satellite measurements make it possible to estimate the value of I_0 to an accuracy

of ± 0.5 percent or better. Following the adoption of the World Radiometric reference (WRR) [2], the present and the most reliable value of I_0 is $1370 \pm 6 \text{ W-M}^{-2}$, (about $1.95 \text{ cal CM}^{-2} \text{ min}^{-1}$) in SI units $\pm 6 \text{ W - M}^{-2}$.

For any mean distance between sun-Earth, the Extra terrestrial radiation can be calculated by considering the Earth's orbital eccentricity factor $e(n)$ that is,

$$I_0' = I_0 e(n) \quad (1.1)$$

Where approximately $e(n) = 1 + 0.034 \cos \left[\frac{2\pi n}{365} \right]$ (1.2) in which 'n' is the day number counted from January 1st. These values are available on a table, FRANK KREITH [3]. In its annual revolution round the sun, the Earth moves not in a circle but in an ellipse. At different times it moves nearer or farther from the sun. Early in January, the Earth passes through the point in its orbit when it is nearest to the sun, known as perhelion, and at the beginning of July, it passes through the farthest point, called the aphe^lion.

1.5 Spectral Distribution of Solar Radiation

The extra atmospheric distribution of energy in the solar radiation spectrum can be found either by direct measurements, or by extrapolation beyond the surface of the atmosphere, of surface spectrometric measurement data. So far direct measurements are taken by means of rockets.

In 1977, the WMO working groups on radiation measurement system of the commission for instruments and methods of observation (CI MO) recommended the use of spectral distribution, summerized by Smith and Gottlieb [4] and the accuracy of the data is about $\pm 2\%$.

The spectral irradiance reaching the Earth's surface, according to the above convention, is in the range of (0.29 - 4.00) μm in which it constitutes 99.125% of the total solar radiation.

The spectral radiation, according to their constituent is described as follows:-

Wave length App. range (μm)	Name	%
(0.29 - 0.40)	Ultra Violet	07.672
(0.40 - 0.75)	Visible	43.806
(0.75 - 4.0)	Infrared	37.697

Table 1.1 - Main Spectral Distribution of Solar Radiation

From the point of view of utilization of solar energy, we are interested in the energy received at the Earth's surface than in the extraterrestrial energy. Due to the following reasons, the solar radiation received by the Earth's surface is entirely different.

1.6 Attenuation of Solar Radiation

The variation in solar radiation reaching the Earth than received at the outside of the atmosphere is due to absorption and scattering in the atmosphere. All atmospheric components contribute to a greater or lesser degree to the attenuation of the direct solar radiation on its path to the Earth's surface. The attenuation of the direct flux of radiation takes place as the result of the mechanisms of absorption and scattering which affects simultaneously all parts of solar spectrum. In the upper layer of the atmosphere, the main process are absorption of the X-ray and ultraviolet regions of the solar spectrum and scattering in the violet and blue ranges.

As the radiation penetrates to lower layer of the atmosphere, the attenuation affects the longer wavelength portion of the solar radiation.

The basic atmospheric gases-nitrogen and oxygen mainly attenuate radiation in the ultraviolet visible regions of the spectrum through molecular scattering. Solar radiation is scattered not only by gas molecules and water vapour, but also by solid drop particles, fog and ice crystals. Diffuse radiation constitutes a significant part of the radiation flux, while part of which goes back to space.

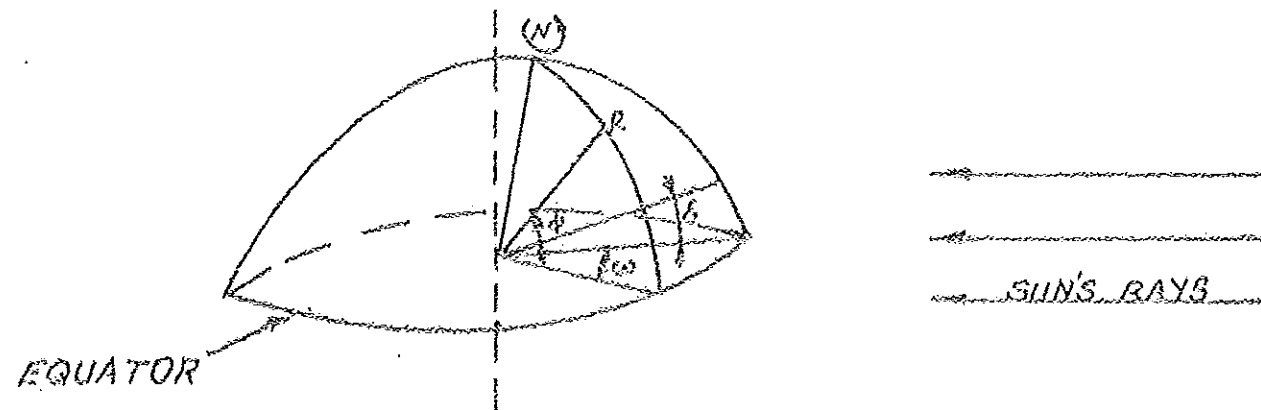
1.7 The geometry of Sun-Earth - Time System

In order to understand, calculation of solar radiation, some basic terms in solar radiation geometry are given.

Basic Sun Earth Angles. The latitude ϕ of a point on the surface of the Earth is its angular distance north or south of the equator measured from the center of the Earth.

The hour angle (ω) is the angle through which the Earth must turn to bring the meridian of a point directly in line with the sun's ray. At solar noon, the hour angle is zero and it expresses the time of the day with respect to solar noon. One hour angle equivalent to $\frac{2\pi}{24} = 0.262$ rad or $\frac{360^\circ}{24} = 15^\circ$ (1.3).

The Sun's Declination (δ) is the angular distance of the sun's rays north (or south) of the equator. It is the angle between a line extending from the center of the sun to the center of the Earth, and the projection of this line upon the Earth's equatorial plane.



The variation of the sun's declination angle δ (the maximum value of δ is at summer solstice* (+23.5°) and winter solstice(-23.5°)) roughly can be found using the equation;

$$\delta = 23.5 \sin \left(\frac{2\pi K}{365} \right) \quad (1.4)$$

Where $k = 0$, at March 21st

For solving practical problems in actinometry, the position of the sun is determined most conveniently by system of horizontal coordinates, using the sun's elevation and azimuth. The sun's altitude angle is determined by the direction towards the sun and the plane of the horizon. The azimuth of the sun ψ , is the angle between the local meridian and the vertical plane of the sun.

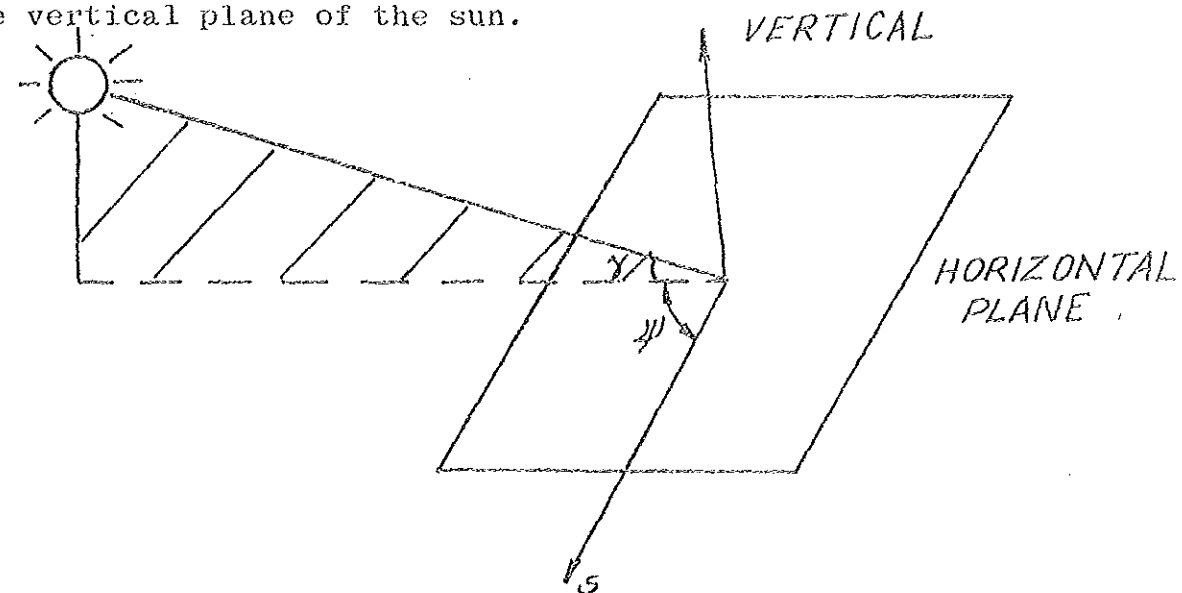


Fig. 1.3 Diagram showing solar altitude angle γ and solar-azimuth angle ψ .

*Solstice means "sun standing still" of min. or max. declination.

Solar attitude and azimuth are not fundamental angles, however they must be related to the fundamental angular quantities, hour angle, latitude and solar declination.

Considering the spherical triangle of the above diagram will have sides, $(90 - \delta)$, $(90 - L)$ and $z = (90 - \gamma)$

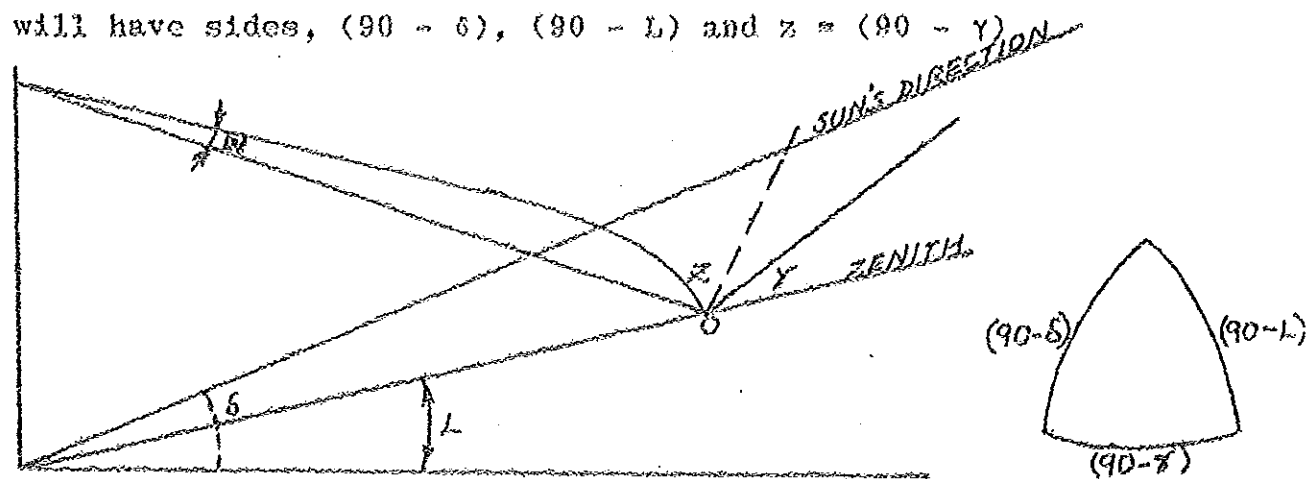


Figure 1.4 - Shows δ , L , ω of the Solar Geometry $Z = (90 - \gamma)$

Using spherical trigonometries according to KELLS KERN and BLAND [5] and using Napier's rule for spherical triangles, "The cosine of any side of a spherical triangle is equal to the product of the cosines of the two other sides increased by the product of the sines of the two other sides and the cosine of the angle included between them."

Symbolically considering the spherical triangle of figure (1.4)

$$\cos(90-\gamma) = \cos(90-L) \cos(90-\delta) + \sin(90-\delta) \sin(90-L) \times \cos(\omega) \quad (1.5)$$

Similarly the azimuthal angle Ψ can be computed as,

$$\sin(\Psi) = \frac{\cos(\delta) \times \sin(\omega)}{\cos(\gamma)} \quad (1.7)$$

Length of the day; Another important factor for solar radiation is the length of the day. This can be calculated as follows; at the time of sun rise or sun set, the zenith angle $z = 90^\circ$, and hence substituting into (1.5) and solving for (ω)

$$\omega = \cos^{-1}(-\tan\phi \tan\delta) \quad (1.8)$$

The length of the day $T_d = 2\omega$

$$2\omega = T_d = 2 \cos^{-1}(-\tan\phi \tan\delta)$$

$$\text{and hence } T_d = \frac{2}{15} \cos^{-1}(-\tan\phi \tan\delta) \quad (1.9)$$

the length of the day T_d is a function of latitude and solar declination.

1.8 The Equation of Time

It is noted that the time specified in all of the sun angle relationships is solar time, which does not coincide with local time.

It is necessary to convert standard time to solar time by applying two corrections. First, there is a constant

correction for any difference in longitude between the location and the meridian on which the local standard time is based. The second correction is from equation of time which takes into account the various perturbations in the Earth's orbit and rate of rotation which affect the time the sun appears to cross the observers meridian. This correction is obtained from published charts. Thus, solar time can be obtained from standard time by this

$$\text{Solar Time} = \text{Standard time} + E + 4 (\text{LST} - \text{LOC}) \quad (1.10)$$

Where E → the equation of time, tables in minutes. (6)

LST → the standard meridian of the local time zone.

LOC → the longitude of the location in question,
in degrees west.

Values of the equation of time for any day in the year can be estimated from a graph known as the Anelmma . and this can be found on solar energy books.

CHAPTER 2

ATTENUATION OF DIRECT SOLAR RADIATION MOLECULAR SCATTERING AND TURBIDITY FACTOR

2.1 Introduction

In recent years requests have come from physicists, engineers and from biologists for more information concerning the quantity and quality of solar radiation reaching the Earth's surface. Increasing awareness of the depletion of the non-renewable energy sources has stimulated investigations into new methods of converting solar energy into useful power.

Among the major factors governing the "heat balance" at the Earth's surface are:

- a) direct solar radiation reaching the Earth's surface after penetration of the atmosphere,
- b) diffuse radiation reaching the Earth's surface,
- c) radiation leaving the Earth's surface (albedo).

The sky radiation (diffuse) includes:

- i) radiation from the extraterrestrial source including stars, moon, planets and the dust particles causing zodiacal light,
- ii) radiation due electric excitation of the atoms and molecules in the atmosphere,

iii) scattered radiation passing through the atmosphere,

iv) "thermal" radiation from the atmosphere.

The contribution of (i) and (ii) to heat balance are negligible as compared to the contribution of (iii) and (iv).

Solar radiation reaches the Earth's surface either directly or after being scattered by atmospheric particles. This paper reports the analysis of the measurements in Addis Ababa, of the molecular scattering of the direct beam radiation, the integral turbidity factor calculation and the contribution of attenuation due to water vapour and aerosol concentrations.

The spectral distribution of solar radiation was first investigated qualitatively by Langley (1881), but there have been few systematic investigations of direct beam spectral irradiance in spite of reports of changes in atmospheric transmission.

Various authors have reported examples of calculated spectral distribution based on various model atmospheres, Unsworth and McCartney (7); measured narrow band spectra has also been published, both in energy units, Dunkel man and Scolnic (8).

In the calculation of the attenuation of the direct solar radiation factors and relations, in the atmosphere, were found by considering the theoretical transfer of heat in the atmosphere.

2.2 Solar Radiation Transfer in the Atmosphere

All the atmospheric components contribute to a greater or lesser degree to the attenuation of the direct solar radiation on its path to the Earth's surface. The attenuation of the direct flux of radiation takes place as a result of the mechanisms of absorption which affect simultaneously all parts of the solar spectrum.

In the upper layers of the atmosphere, the absorption of solar radiation is caused by oxygen, oz-one and nitrogen oxides and, in the lower stratosphere and troposphere, by water vapour, carbondioxide, aerosols and other minor components. The variations in the concentration of water vapour and oz-one and optical characteristics of aerosols under the influence of various processes of interaction between solar radiation field, the atmosphere and the underlying surface result in constant fluctuations in spectral transmissions in the atmosphere.

2.3 Attenuation of Direct Radiation

Radiation at normal incidence received at the surface of the Earth from the sun is subject to variation due to ,

1. variations in distance from Earth to sun,
2. variations in atmospheric scattering by air molecules, water vapour, and dust, and
3. variations in atmospheric absorption by O_2 , O_3 , H_2O and CO_2 .

Scattering, which results in attenuation of direct radiation by air molecules, water vapour, and dust, has been the subject of a number of studies and an approximate methods have been developed to estimate the magnitude of the effect.

2.4 The Equation of Radiative Transfer

Due to great complexity of the processes that determine the radiative transfer, the transfer equation is complex. In the atmospheric radiative transfer we usually confine the problem to consideration of radiative transfer for a stationary field of unpolarized radiation, without taking account of refraction. Kondrayatev [9] .

The thermal emissions were considered unpolarized and the refraction in the atmosphere is neglected in the general consideration of the problem.

In the derivation of the general transfer equation for a stationary field of unpolarized radiation, we take an arbitrarily directed ray r -direction and consider an environmental element in the form of a cylinder of unit section, the axis of which coincides with the direction of the ray. Let the ray cut the basis, perpendicularly as shown in the Figure 2.1 E.M. Feigel'son [10] .

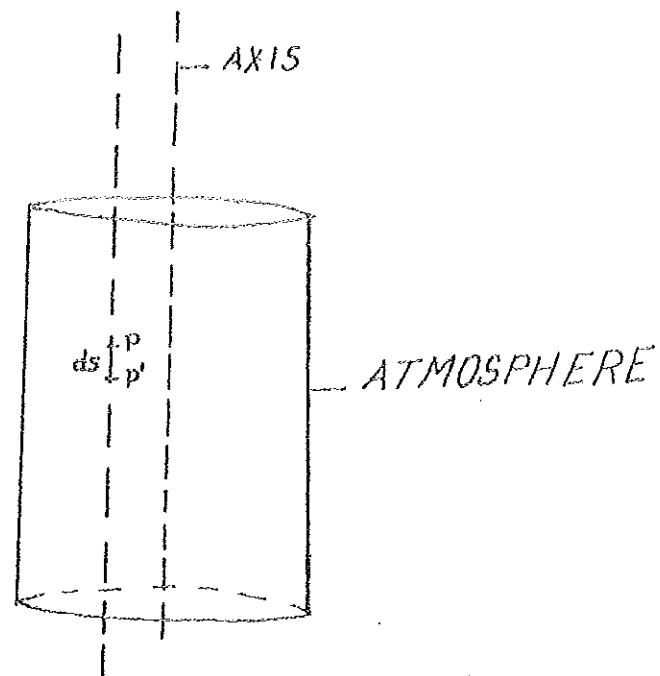


Fig.2.1: Cylindrical Cross Section of Environmental Atmosphere

Then, the intensity of radiation at point p and p' be

$$I_{\lambda}(p', r) = I_{\lambda}(p, r) + \frac{\partial I_{\lambda}}{\partial s} ds \quad (2.1)$$

The variation of intensity from point p and p' is caused by

- a) attenuation due to absorption of radiant energy, which can be expressed as

$$-k_{\lambda}(p) I_{\lambda}(p, r) \rho(p) ds \quad (2.2)$$

Where ρ , is the density and k_{λ} is the mass absorption coefficient at point p .

- b) Attenuation due to scattering radiant energy by considered environmental element, which can be presented in the form

$$-\delta_{\lambda}(p) I_{\lambda}(p, r) \rho(p) ds \quad (2.3)$$

where δ_{λ} is scattering coefficient.

- c) Increase of the intensity of radiation owing to the considered element of the medium emitting in the direction r , which can be written as,

$$n_{\lambda} \rho(p) ds \quad (2.4)$$

- d) Increase of the intensity of radiation resulting from scattering process due to which the rays of various direction passing through elementary cylinder add a part of their energy to the ray of the direction r .

If we consider the ray of the direction ' r ' through an element of medium, if the part of this energy equal to

$$\delta_{\lambda}(p) I_{\lambda}(p, r') \rho ds \quad (2.5)$$

will be scattered by this element and a part

$$\delta_{\lambda}(p) I_{\lambda}(p, r') \frac{\gamma_{\lambda}(p, r', r)}{4\pi} \rho ds \quad (2.6)$$

will go in the direction r .

Here the expression $(\frac{1}{4\pi})\gamma_{\lambda}(p, r, r')$ characterizes the scattering function, and this means the total quantity of the scattered radiation, the portion of $(\delta_{\lambda} I_{\lambda} \rho)$ equal to

$(\frac{1}{4\pi} \gamma \delta_{\lambda} I_{\lambda} \rho ds)$ departs in the direction r . Integrating the whole expression over all possible directions, we have,

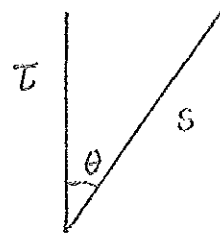
$$\frac{\delta_{\lambda}(p)}{4\pi} \int I_{\lambda}(p, r') \gamma_{\lambda}(p, r', r) d\omega' \rho ds \quad (2.7)$$

Then summing over all (a, b, c, d) , and

taking into consideration all obtained results and reducing by ds we get

$$\frac{1}{\rho} \frac{dI_\lambda}{ds} = \eta_\lambda + \frac{\delta_\lambda}{4\pi} \int I_\lambda(p, r') \gamma_\lambda(p, r', r) d\omega' - (k_\lambda + \sigma_\lambda) I_\lambda \quad (2.8)$$

As mentioned by the relation, substituting a change of variable as shown in the diagram,



$$\cos \theta ds = d\tau$$

Fig. 2.2 Atmospheric Thickness and Sun's Ray

$$\frac{1}{\rho} \frac{dI_\lambda}{ds} = \frac{1}{\rho} \frac{dI_\lambda}{dz} \frac{dz}{ds} = \frac{1}{\rho} \cos \theta \frac{dI_\lambda}{d\tau} \quad (2.9)$$

where τ is the atmospheric thickness.

$$\cos \theta \frac{1}{\rho} \frac{dI_\lambda}{d\tau} = \eta_\lambda + \frac{\delta_\lambda}{4\pi} \int I_\lambda(p, r') \gamma_\lambda(p, r', r) d\omega' - (k_\lambda + \sigma_\lambda) I_\lambda \quad (2.10)$$

and the relation $I_\lambda(p, r) \rightarrow I_\lambda(\tau, r)$

$$\rightarrow \cos \theta \frac{1}{\rho} \frac{dI_\lambda}{d\tau} = \eta_\lambda + \frac{\delta_\lambda}{4\pi} \int I_\lambda(\tau, r') \gamma_\lambda(\tau, r', r) d\omega' - (k_\lambda + \sigma_\lambda) I_\lambda \quad (2.11)$$

If the medium is in thermodynamic equilibrium, the emissivity and the absorptivity of the medium are related by Kirchoff's law,

$$\eta_\lambda = K_\lambda B_\lambda(T) \quad (2.12)$$

$$\text{where } B_{\lambda}(T) = \frac{2hc^2}{\lambda^5} \frac{1}{(\text{Exp}(hc/K_{\lambda}T) - 1)} \quad (2.13)$$

The lower atmosphere (troposphere) generally satisfies the condition of the local equilibrium. Then the eqn. becomes,

$$\cos \theta \frac{dI_{\lambda}}{d\tau_{\lambda}} = -I_{\lambda}(\tau, \theta) + \frac{B_{\lambda}(T)}{4\pi} \int I_{\lambda}(\tau, x') \gamma(\tau, x, x') d\omega' \quad (2.14)$$

Where, $a_{\lambda} = \mathcal{G}_{\lambda}^{\delta} + \mathcal{G}_{\lambda}^{\kappa}$ and $\mathcal{G}_{\lambda}^{\delta} = (a_{\lambda} - \mathcal{G}_{\lambda}^{\kappa}) = \beta_{\lambda} a_{\lambda}$, $a_{\lambda} I_{\lambda} = I_{\lambda}(\tau, x)$

In a cloudless atmosphere the absorption by atmospheric gases is neglected and the equation becomes,

$$\cos \theta \frac{dI}{d\tau} = -\frac{1}{4\pi} \int I(\tau, x') \gamma(\tau, x, x') d\omega' - I(\tau, x) \quad (2.15)$$

Separating the unknown $I(\tau, x)$ between the diffuse and the direct radiation, we get,

$$I(\tau, x) = I_s(\tau, x) + I_r(\tau, x) \quad (2.16)$$

Where $I_s(\tau, x)$ is the intensity of the parallel beam of the sun rays determined from the above equation, for the direct beam we have $\frac{d\omega}{4\pi} = 0$, this implies,

$$\cos \theta \frac{dI}{d\tau} = -I(\tau, x) = -I_s(\tau, x) \quad (2.17)$$

Solving the equation using the boundary condition

$$I_s(\tau^*, x) = \pi s \delta(x, x_0) \quad (2.18)$$

$$\rightarrow \frac{dI}{d\tau} = -I_s \text{ See } \theta$$

∴ $I_s(\tau, r) = c e^{-\sec \theta \tau}$ using the boundary condition

$$I_s(\tau, r) = \pi s \delta(r-r_0) e^{-\sec \theta (\tau^* - \tau)} \quad (2.19)$$

Where $\sec \theta = m$; m , the atmospheric mass and θ is the zenith angle determined by the relation of equations (1.5)

and $(\pi s \delta(r-r_0))$ where $r=r_0$ is the extraterrestrial solar radiation.

Denoting $\pi s \delta(r-r_0) = I_0$ extraterrestrial radiation and $m = \sec \theta$, the equation becomes,

$$I(\tau, r) = I_0 (e^{-m\tau}) \quad (2.20)$$

And this equation is known as Lambert-Beer equation.

2.5 Spectral Distribution of Direct Solar Radiation

The attenuation of monochromatic radiation in the direct beam by the atmosphere may be described by Lambert-Beer's equation, thus, the total energy integrated over all wavelength in the direct solar beam at normal incidence on the Earth's surface,

$$I_{N,R} = \int_{0.29\mu m}^{4.0\mu m} I(\lambda) d\lambda \quad (2.21)$$

$$\rightarrow I_{N,R} = \int_{0.29\mu m}^{4.0\mu m} I_0(\lambda) \text{Exp}\{-m(\theta)\tau(\lambda)\} d\lambda \quad (2.22)$$

Where $I_0(\lambda)$ extraterrestrial radiation and $m(\theta)$ is atmospheric air mass in the direction of the sun. The radiation in the interval (0.29 - 4.0) μm constitutes 99.195% of the solar radiation according to WMO(4)bulletin.

The integrand of eqn. 2.22 can be written as

$$I(\lambda) = I_0(\lambda)e^{-\tau(\lambda)m(\theta)} \quad (2.23)$$

Where the atmospheric thickness

$$\tau(\lambda) = \tau_R(\lambda) + \tau_g(\lambda) + \tau(\lambda) \quad (2.24)$$

and denoting $p(\lambda) = e^{-\tau(\lambda)}$ (2.25) where $p(\lambda)$ is the transparency coefficient.

$$I_{N,R} = \int_{0.29\mu\text{m}}^{4.0\mu\text{m}} I_0(\lambda) \text{Exp}\{-m_h(\theta)(\tau_R(\lambda) + \tau_g(\lambda) + \tau(\lambda))\} d\lambda \quad (2.26)$$

Where, $\tau_R(\lambda)$, $\tau_g(\lambda)$, and $\tau(\lambda)$ are respectively the vertical attenuation coefficients for molecular scattering, absorption by atmospheric gases (water vapour, carbon dioxide, ozone, oxygen) and aerosol extinction.

Using eqn.(2.25) we get,

$$I(\lambda) = I_0(\lambda) p_{\lambda}^{m(\theta)} \quad (2.27)$$

then integrating over all possible wavelengths we get,

$$I_m = \int_{0.29\mu\text{m}}^{4.0\mu\text{m}} I(\lambda) d\lambda = \int_{0.29\mu\text{m}}^{4.0\mu\text{m}} I_0(\lambda) p_{\lambda}^{m(\theta)} d\lambda \quad (2.28)$$

Then the transparency coefficient, averaged over the whole spectrum

$$I_m = P_m^m \int_{0.29\mu\text{m}}^{4.0\mu\text{m}} I_o(\lambda) d\lambda = I_o P_m^m \quad (2.29)$$

$$P_m^m = \frac{\int_{0.29\mu\text{m}}^{4.0\mu\text{m}} I_o(\lambda) P_\lambda^m d\lambda}{\int_{0.29\mu\text{m}}^{4.0\mu\text{m}} I_o(\lambda) d\lambda} \quad (2.30)$$

By taking into account of the radiation due, to, molecular scattering and absorption of radiation by water vapour and aerosol as,

$$\tau(\lambda) = \int K_\lambda \rho dh + \int a_{\omega, \lambda} \rho_\omega dh + \int a_{d, \lambda} \rho_d dh. \quad (2.31)$$

Where ρ , ρ_ω and ρ_d are molecular, water vapour and aerosol concentration respectively. From (2.31) it can be written as,

$$\tau(\lambda) = T_\lambda \int K_\lambda \rho dh \quad (2.32)$$

$$\text{where } T_\lambda = 1 + \frac{\int a_{\omega, \lambda} \rho_\omega dh}{\int K_\lambda \rho dh} + \frac{\int a_{d, \lambda} \rho_d dh}{\int K_\lambda \rho dh} \quad (2.33)$$

Therefore, the monochromatic direct solar radiation flux at

a given solar zenith distance (given atmospheric mass) can be presented in the form,

$$I_m(\lambda) = I_0(\lambda) \text{Exp} \left\{ - I_{\lambda m} \int K_{\lambda} \rho dh \right\} \quad (2.34)$$

$$\text{Setting } q(\lambda) = \text{Exp} \left\{ - \int K_{\lambda} \rho dh \right\} \quad (2.35)$$

Then for full flux of direct solar radiation I_m , becomes,

$$I_m = \int_{0.29\mu m}^{4.0\mu m} I_0(\lambda) q(\lambda) m^{T_{\lambda}} d\lambda \quad (2.36)$$

If q_m and T_m are the averaged values over the whole spectrum we get,

$$I_m = q_m^{mT_m} \int_0^{\infty} I_0(\lambda) d\lambda = I_0 q_m^{mT_m} \quad (2.37)$$

Where, T_m is the integral turbidity factor.

$$\text{And hence, } \frac{I_m}{I_0} = q_m^{mT_m} \quad (2.38)$$

Comparing with the expression of the transparency coefficient, we get,

$$P_m^m = \frac{I_m}{I_0} = q_m^{mT_m} \quad (2.39)$$

Solving for the integral turbidity factor T_m from (2.39)

$$T_m = \frac{1}{m \ln q_m} \ln \left(\frac{I_m}{I_0} \right) \quad (2.40)$$

denoting $I_{ideal} = I_0 q_m^m$, where I_{ideal} is only molecular scattering where q_m is as defined in (2.35).

It can be shown that the following relation holds namely,

$$\frac{1}{m \ell_n q_m} = \frac{1}{\ell_n \left(\frac{I_{ideal}}{I_0} \right)} \quad (2.41)$$

Substituting (2.41) into (2.40) and simplifying we get

$$I_m = \frac{\ell_n(I_m) - \ell_n(I_0)}{\ell_n(I_{ideal}) - \ell_n(I_0)} \quad (2.42)$$

Where I_{ideal} , is the radiation received by the Earth depleted due to molecular scattering. (Rayleigh scattering mechanism).

In order to calculate the integral turbidity factor (2.42), it is important to calculate I_{ideal} , the solar radiation due to molecular attenuation.

2.6 Molecular Scattering of Solar Radiation (Rayleigh Scattering)

The most important contribution to this field was made by Lord Rayleigh in early 1870, who contended that air molecules were the causes of light scattering. The molecular scattering of Rayleigh is the scattering of light caused by density fluctuations.

The following fundamental assumptions were taken into consideration.

- a) The dimensions of the scattering particles (Rayleigh particles) which scatter the radiant energy in this case are roughly less or equal to 0.1λ , as compared to the wavelength (λ) of the solar spectrum.
- b) The scattering particles and the medium are not conductors and do not contain free electric charges.
- c) The particles scatter light independently to each other.

The Rayleigh volume scattering coefficient of standard air, at a particular wavelength λ may be written as

$$\beta(\lambda) = \frac{32\pi^3(n_\lambda - 1)^2}{3\lambda^4 N} \left| \frac{6 + 3\rho_{n\lambda}}{6 - 7\rho_{n\lambda}} \right| \quad (2.43)$$

Where N_s is the number density (mole/cm³), ρ_n is the depolarization factor and n_s is the refractive index of standard air.

The spectral solar radiation intensity reaching the Earth's surface after depletion due to scattering by air molecules may be written as,

$$I_\lambda(0, h) = I_{0\lambda} \exp[-\beta(\lambda) m_\lambda(0)H] \quad (2.44)$$

Where ($H = 7996\text{m}$) is the vertical height of the imaginary homogenous atmosphere at NTP.

Denoting the optical thickness of the atmosphere as

$$\tau(s, s') = \int_{s'}^s \beta ds \quad (2.45)$$

where s' is at the surface of the Earth and s at the top of the atmosphere, then we can write,

$$I_{\text{Ideal}}(\theta, h) = \int_{0.29\mu\text{m}}^{4.0\mu\text{m}} I_0(\lambda) \text{Exp} \{-\tau(\lambda)m(\theta)\} d\lambda \quad (2.46)$$

Where, $\tau(\lambda)$ is Rayleigh molecular attenuation coefficient and $I_{\text{Ideal}}(\theta, h)$ & $I_0(\lambda)$, are spectral energy at the wavelength λ , on the Earth's surface and outside the Earth's surface respectively.

And the airmass $m(\theta)$ is the absolute airmass which takes account, the stations altitude h , by means of the formula Kodrayatev [11] .

$$M_h(\theta) = \frac{P_h}{P_0} (m(\theta)) \quad (2.47)$$

Where P_h is the pressure at the station and P_0 pressure at sea level.

Molecular attenuation coefficients $\tau(\lambda)$, calculated from Rayleigh's theory via (2.44) and (2.45), were tabulated by Pendorf [12] and the evidence supporting the accuracy of theory was reported by Chandra M. [13] .

The integrated effect of gaseous absorption on the attenuation of solar radiation has been investigated by several

authors, MacDonald [14] ; Roach [15] ; and Yamamoto [16] .

Thus, the total energy integrated over all wavelengths in the direct solar beam at normal incidence on the Earth's surface, after having traversed the Rayleigh atmosphere is calculated via (2.46) .

2.7 Atmospheric Turbidity

Several atmospheric turbidity coefficients have been introduced during the past decades in order to quantify the influence of atmospheric aerosol content on a direct radiation received at the Earth's surface. The most currently used are Linke factor T , and Angstrom Coefficient β . Turbidity usually occurs in conditions of high humidity, and in the presence of hygroscopic aerosol particles. Variations in atmospheric turbidity are also possible in conditions of low relative humidity if air contains dry aerosol particles, natural or man made-origin.

In practical calculations of the attenuation of solar radiation in the atmosphere due to aerosol scattering, extensive use is made of the method proposed by Angstrom, expressed as,

$$a_D(\lambda) = \beta_0 \lambda^{-\alpha} \quad (2.48)$$

Where, β is the Angstrom turbidity coefficient, can be determined by measuring direct solar radiation in the band (0-0.630) μm , and α is an index related to dimensions of

the scattered particles and the nature of their size distribution. This is an empirical relationship obtained from the observational data.

Wide band direct beam solar irradiance can be similarly analysed in terms of aerosol attenuation (turbidity).

H.A. McCARTNEY and M. HUNSWORTH [17] used an integral turbidity coefficient, $\tau(\Delta\lambda)$, to describe aerosol attenuation in a finite wave band, $\Delta\lambda$, by analogy to Lambert-Beer's law (2.20).

$$\tau(\lambda) = \frac{|\ln(I_*(\Delta\lambda)) - \ln\{I(\Delta\lambda)\}|}{m_h(\theta)} \quad (2.49)$$

Where $I(\Delta\lambda)$ is the direct irradiance in the wave band $\Delta\lambda$ and $I_*(\Delta\lambda)$ is irradiance emerging from an aerosol free atmosphere.

Linke Turbidity Factor

In practice, the characteristic of atmospheric transparency is widely used; the turbidity factor (Trübungs Faktor) denoted by T was proposed by Linke (Sutton) [18].

The turbidity factor is defined as the ratio of the coefficient of total extinction of solar radiation in a real atmosphere to the coefficient of extinction by the component of ideal atmosphere, i.e., by molecular scattering and selective absorption by ozone, oxygen and carbon dioxide.

The turbidity factor gives a clear idea of radiation extinction, since it characterizes variable extinction in the real atmosphere relative to constant extinction in ideal atmosphere.

Recalling (2.33) and denoting,

$$W = \frac{\int a_{w,\lambda} \rho_w dh}{\int k_\lambda \rho dh}, \quad R = \frac{\int a_{d,\lambda} \rho_d dh}{\int k_\lambda \rho dh} \quad (2.50)$$

We get, $T = (1 + W + R)$ (2.51)

Corresponding to selective (water vapour) and aerosol extinction of solar radiation, the turbidity factor, as might be expected, shows marked variations with time and place. The greatest changes in variation of atmospheric turbidity are those which arise from atmospheric pollution.

From the definition factor it follows that $T = 1$ if the attenuation of solar radiation is caused by molecular scattering only. The second term on the right-hand side of (2.51), the value of W , is called humid turbidity factor, and characterizes the influence of absorption and scattering by water vapour on the attenuation of solar radiation.

The third term R , is known as the residual turbidity factor, determining the effect of radiation attenuation due to absorption and scattering by dust and aerosol particles. It is possible to calculate the aerosol and water vapour attenuation as,

$$R = T - 1 - W \quad (2.52)$$

Where $W = 0.5 e_o^{0.43}$, and e_o absolute humidity Sivkov [19].

2.8 Results

The calculations for radiation depleted due to Rayleigh scattering mechanism and the integral turbidity factor were calculated using (2.46) and (2.42).

a) Molecular Scattering of Solar Radiation

The attenuation of solar radiation due to molecular scattering can be theoretically calculated for any locality. And this value can have a slight difference from place to place, considering the atmosphere to be ideal or the constitution of the gases in the atmosphere has the ideal percentage.

The variation only comes due to the airmass variation calculation from place to place. As it has been seen in (2.47) the airmass is dependent on the pressure value at a given place, which in turn takes into consideration the altitude of a given locality.

In our calculation for the A.A. region of the year 1983, the atmospheric pressure was in the interval (76,660.36 - 77,060.32) Pa. according to NMA* compared to the (101,325) Pa. at sea level. This has introduced a difference in calculation of the airmass at sea level and in Addis Ababa of altitude 2,408 m above sea level.

* NMA, National Meteorological Agency.

The attenuation of the direct solar radiation due to Rayleigh atmosphere was calculated using (2.46). For the range 0.29 to 4.0μ , in steps of 0.01μ between 0.2 and 0.8μ , in steps of 0.1μ from there on upto 2μ , in steps of 0.5μ between 2 and 4μ .

And the values $I_0(\lambda)$, extraterrestrial and $\tau(\lambda)$, attenuation coefficient for the respective wavelength was considered in the calculations.

The integrations for the mentioned equation was carried on by Simpson's integration rule with the above steps. The input data used in the calculations—direct solar radiation, pressure of each day, (source NMA) and declination angle (δ) and Earth's elliptic orbit escentricity $E(I)$ (are shown on appendix A. The direct solar radiation was measured by Eppley pyrliometer. The wavelength interval (λ), the Rayleigh molecular attenuation coefficient $\tau(\lambda)$, extraterrestrial radiation $I_0(\lambda)$ for the spectrum in the interval ($0.29 - 4.0$) μ are shown on appendix A.

Three different airmass at 13:00, 14:00 and 15:00 hrs LST were considered in the calculation and three airmass using (2.19) and (2.47) are tabulated on Appendix B.

The calculation of solar radiation received by the Earth's surface due to Rayleigh scattering mechanism, using (2.46) was carried on by a computer for the days of the year 1983 of A.A. region, and calculated values are reported on Appendix C.

The values are in the interval (931,71 - 1,012,98) $W-M^{-2}$, comparing this with the extraterrestrial radiation $I_0 = 1,370 \pm 6 W-M^{-2}$ shows that (26.05 - 31.95) % of the extraterrestrial radiation was depleted due to molecular scattering. The annual evolution of the monthly mean values of this scattering is shown in Figure (2.3).

b) The Turbidity Factor 'T'

The turbidity factor was also calculated using (2.42). The values for 335 days of the year 1983 of A.A. region were calculated by having the corresponding values of I_m measured, and these are on appendix A, I_{ideal} calculated for each day. In calculating the turbidity factor the value of I_0 (extraterrestrial radiation - the recent value $I_0 = 1,370 \pm 6 W-M^{-2}$) according to WMO Technical Note was adopted.

The values of the turbidity factor for each day of the whole year are reported on Appendix C and the annual evolution of 'T' monthly mean values are shown in Figure (2.4). The integral turbidity factor as might be expected, shows marked variations with time and place. There are usual variations in summer and winter, but the greatest changes are those which arise from the atmospheric pollutions.

FIG.2.4 THE ANNUAL EVOLUTION OF TURBIDITY MONTHLY
MEAN VALUES.

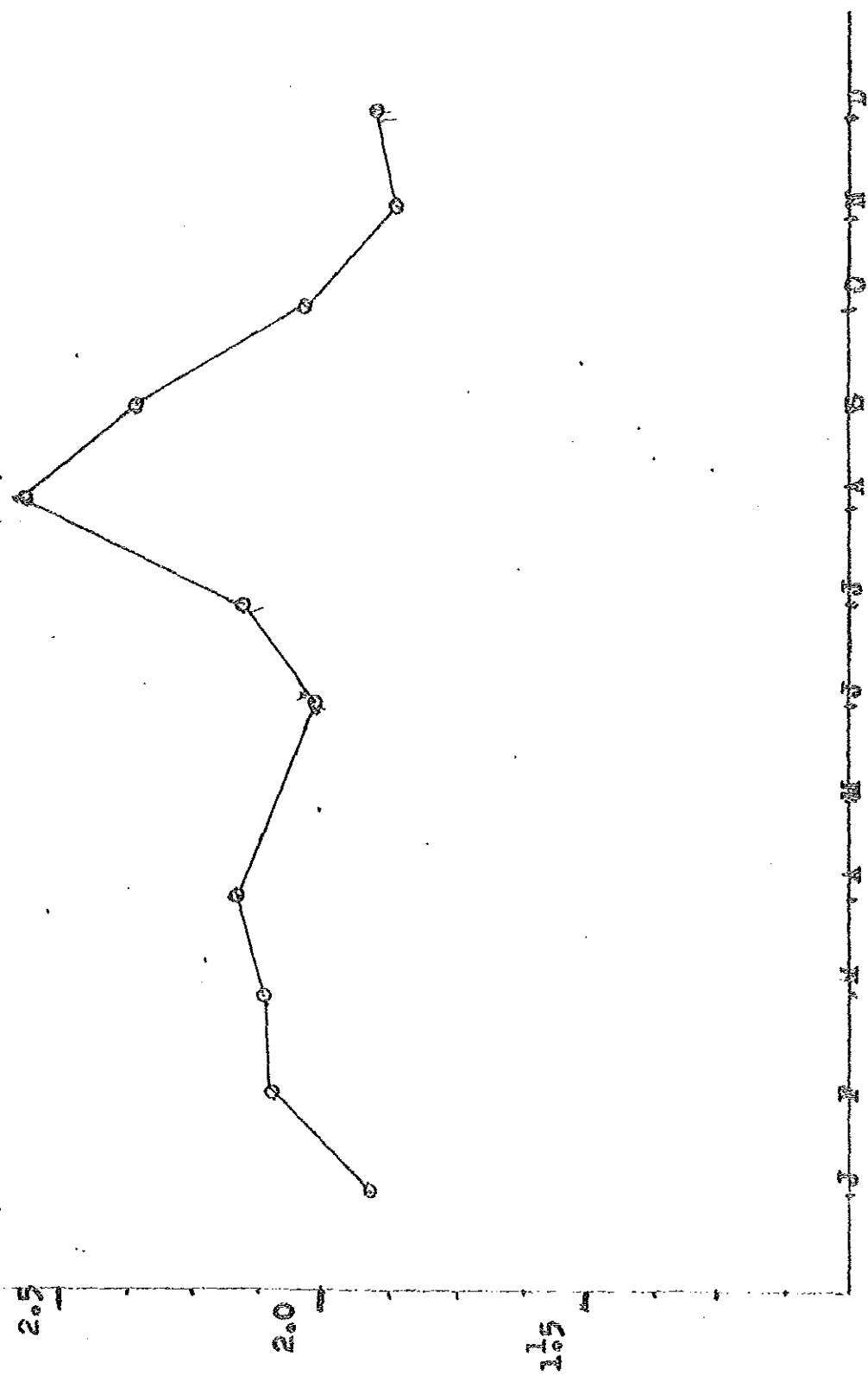
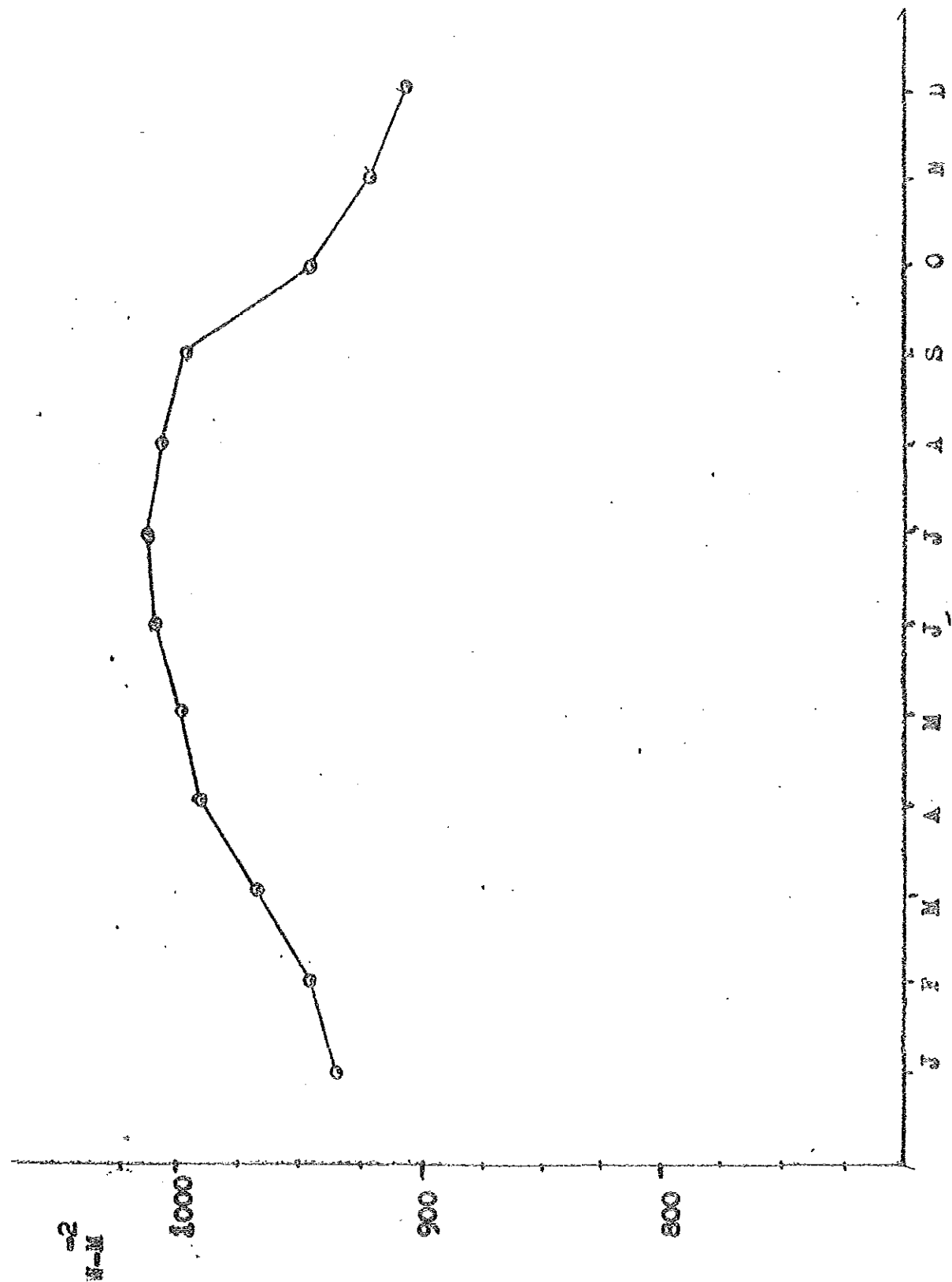


FIG. 2.3 THE ANNUAL EVOLUTION OF THE MONTHLY MEAN VALUE OF RAYLEIGH SCATTERING.



As it has been done for the molecular scattering, the direct radiation measured was considered only for the time (13:00 - 15:00) EST. This has been chosen due to airmass variation is negligible, however, the molecularly scattered radiation received by the Earth was calculated at three different hour angles and averaged values are tabulated on Appendix.C for the whole year. This work could have been continued to evaluate aerosol concentration using (2.52) , however, lack of data of absolute humidity has hindered the process. Correlation of turbidity factors could have also been done by a linear regression $\beta = a + bT$. M. Katz [20], but lack of measured values of I_m in a different wave bands, has also hindered the process.

CHAPTER 3

Estimation of Global Solar Radiation

From Sunshine Hours,

Geographical & Meteorological Parameters

3.1 Estimation of Global Solar Radiation With Various Empirical Formulas

For the proper planning of the utilization of solar energy, it is necessary to know the total radiation falling in a particular region.

There are various methods for estimating the global solar radiation but it seems no single form of equation is applicable to all atmospheric environments and their applicabilities should be tested for a particular region.

The possible approaches, in estimating the global solar radiation can be made employing one of the following:

i) Using equations relating global solar radiation to number of bright sunshine hours.

ii) Using equation relating global values to other meteorological parameters.

Many authors have proposed various equations by using the first, the second approaches. The latter considers a large number of meteorological parameters, and is more dependent on the considered area.

These formulas were tested at various places, most of the equations are, conversion from daily sunshine duration to daily global radiation by Angstrom and Schuepp formula.

$$G = G_0 (a + b (S/S_0)) \quad \text{Angstrom} \quad (3.1)$$

$$G = a + b \left(\frac{S}{S_0} \right) \left(1 + \frac{S}{S_0} \right)^{\frac{1}{2}} \quad \text{Schuepp} \quad (3.2)$$

Where S and S₀ are daily hours of bright sunshine, maximum possible daily hours of bright sunshine hours and G₀, clear sky global radiation, respectively.

Using (3.1) R.H.B. Excel [21], A.J.Biga [22], A.I.Kundish [23] and make use of 3.1 and 3.2 by Cesen [24] working group, P.Kulisk [25], T.D.M.A. Samuel [26] and F. Kasten [27] had used these equations for various regions in the world and reported the regression constants ('a' and 'b') determined for their localities. However, more recently using (3.1) determination of regression constants 'a' and 'b', based on 11 year solar radiation data at Trieste, found substantial temporal and spatial scattering. P. C. Jain [28].

However, estimation of solar radiation from meteorological data, such as humidity, temperature and geographical factor latitude and attitude are more related to solar radiation. One has to consider all the parameters, and

deleting some of these may bound to an error. using these parameters, A. M. Sayigh et.a. [29] in 1975 deduced a formula, that can predict solar radiation on a horizontal plane.

$$G = 1.53 K \text{ Exp } \left\{ \phi \left(D - \frac{R^{1/3}}{100} - \frac{1}{T_{\max}} \right) \right\} \quad (3.3)$$

Where, ϕ is the latitude of the place in radians,
R is relative humidity, D, relative sunshine hours to
12 hrs., T_{\max} , maximum temprature of the day in °C.

and $K = (\lambda z + \text{Cos}\phi \Psi_{ij}) \times 10^2 \text{ (cal + cm}^{-2} \text{ - day}^{-1}\text{)}$

$$\lambda = 0.2/(1+0.1\phi)$$

z, the length of the day, Ψ_{ij} seasonal factor.

The seasonal factor has i = 1,2,3 stands for inland
station, coastal station and hilly station respectively.

J = 1,2,3 ---- 12, number of months starting from January.

The monthly value is given on table (3.1)

	1	2	3	4	5	6	7	8	9	10	11	12
Ψ_1	1.28	1.38	1.54	1.77	2.05	2.30	2.48	2.41	2.36	1.73	1.38	1.17
Ψ_2	1.46	1.77	2.05	2.15	2.05	2.05	2.10	2.17	2.14	1.96	1.60	1.43
Ψ_3	1.60	1.81	2.00	2.17	2.25	2.26	2.24	2.20	2.10	1.92	1.74	1.60

Table 3.1 Monthly seasonal factor values

The selection of the seasonal factor values using the above criteria and humidity (annual) of the given locality should be taken.

The equation (3.3) by Sayigh et. al has been used to obtain daily global solar radiation, by considering the corresponding data of meteorological parameters for the Addis Abeba region. The results were compared with the daily measured value of global solar radiation.

However, discrepancies ranging (20 - 300)% were obtained between the calculated and measured global solar radiation.

Therefore, the following general form of the Sayigh et. al. equation was considered.

$$G(X_i, \theta_k) = \alpha K \exp(-A_1 \phi^A) (BID^B + CIR^C + EI T_{max}^E) \quad (3.4)$$

$$X_i = (K, \phi, D, E, T_{max}) \quad (3.5)$$

and $\theta_k = (\alpha, A_1, A, B_1, B, C, EI, E)$ (3.6) are general constants.

Due to the above reasons an effort was done to determine the general constants using the global solar radiation, and meteorological parameters of the year 1983 for the Addis Abeba region.

A mathematical method modified by H. O. Hartley [30], "The modified Gauss - Newton method for fitting of non-linear regression functions by least squares" was used in determining the above nine constants for Addis Ababa region.

3.2 The Formulation of the Regression

The formulation of the non-linear regression problem was, given for n sets of "observed" $(k + 1)$ tuples, $Y_h; X_{1h}, X_{2h}, \dots, X_{kh}$.

$$h = (1, 2, 3 \dots n)$$

$$G(X, \theta) = G(X_1, \dots, X_k, \theta_1, \dots, \theta_m) \quad (3.7)$$

It is required to determine, a set of θ_i for which the sum of the squares:

$$Q(\theta) = \sum_{h=1}^n (y_h - G(xh, \theta))^2 = \text{Min.} \quad (3.8)$$

Where, Y_h is the measured global solar radiation.

i) Introducing the first derivative of $G(X_h, \theta_k)$ with regard to θ_i as.

$$\frac{\partial G}{\partial \theta_i} = G_i(X, \theta), \text{ we get, nine sets of equations.}$$

$$G_1(X, \theta) = \frac{\partial G(X, \theta)}{\partial \alpha} = K \exp A11^A (B1D^B + C1R^C + E1 T^E) \quad (3.9)$$

$$G_2 (X, \theta) = \frac{\partial G (X, \theta)}{\partial A_1} = (\alpha K (L^A (B_1 D^B + C_1 R^C + E_1 T^E))) \times \text{Exp} \{ (A_1 L^A (B_1 D^B + C_1 R^C) + D_1 T^D) \} \quad (3.10)$$

$$G_3 (X, \theta) = \frac{\partial G (X, \theta)}{\partial B_1} = (\alpha K (A_1 L^A D^B) \text{Exp} \{ (A_1 L^A (B_1 D^B + C_1 R^C + E_1 T^E) \} \quad (3.11)$$

$$G_4 (X, \theta) = \frac{\partial G (X, \theta)}{\partial C_1} = (\alpha K (A_1 L^A R^C) \text{Exp} \{ (A_1 L^A (B_1 D^B + C_1 R^C + E_1 T^E) \} \quad (3.12)$$

$$G_5 (X, \theta) = \frac{\partial G (X, \theta)}{\partial E_1} = (\alpha K \text{Exp} \{ (A_1 L^A (B_1 D^B + C_1 R^C + E_1 T^E) \} \times (A_1 L^A T^E) \quad (3.13)$$

$$G_6 (X, \theta) = \frac{\partial G (X, \theta)}{\partial A} = (\alpha K \text{Exp} \{ (A_1 L^A (B_1 D^B + C_1 R^C + E_1 T^E) \} \times A_1 (B_1 D^B + C_1 R^C + E_1 T^E) (L^A \ln(L)) \quad (3.14)$$

$$G_7 (X, \theta) = \frac{\partial G (X, \theta)}{\partial B} = (\alpha K \text{Exp} \{ (A_1 L^A (B_1 D^B + C_1 R^C + E_1 T^E) \} \times (A_1 L^A B_1) \ln (D) D^B) \quad (3.15)$$

$$G_8 (X, \theta) = \frac{\partial G (X, \theta)}{\partial C} = (\alpha K \text{Exp} \{ (A_1 L^A (B_1 D^B + C_1 R^C + E_1 T^E) \} \times (A_1 L^A C_1) \ln (R) R^C) \quad (3.16)$$

$$G_9 (X, \theta) = \frac{\partial G (X, \theta)}{\partial E} = (\alpha K \text{Exp} \{ (A_1 L^A (B_1 D^B + C_1 R^C + E_1 T^E) \} \times (A_1 L^A E_1) \ln (T) T^E) \quad (3.17)$$

Then using the above definitions of (3.8) it is possible to write,

$$\frac{\partial Q}{\partial \theta_i} = Q_{\theta_i}(x; \theta) = -2 \sum_{h=1}^n (Y_h - G(X_h; \theta)) G_i(X_h, \theta) \quad (3.18)$$

3.3 The modified Gauss - Newton Iteration:- We start with initial values of the constants $\theta_0 = (\alpha_0, A1_0, A_0, B1_0, C_0, E1_0, E_0)$

The first step is compute "Corrections" to the elements θ_0 of the starting vector θ_0 . These correction will be proportional to the solutions P_i of the Gauss - Newton equation corresponding to (3.8)

Expanding $G(X, \theta)$ by Taylor's series at $\theta = \theta_0$ and considering upto the linear term and substituting into (3.8) we get:

$$2 \sum_{j=1}^m \left\{ \sum_h G_i(X_h, \theta_0) G_j(X_h, \theta_0) \right\} P_j = -Q_i(X; \theta_0) \quad (3.19)$$

Where $P_j = (\theta_j - \theta_0)$ for our case $i = 1, 2 \dots 9$ (3.20)

and,

$$\begin{aligned} -\frac{1}{2} Q_1(X, \theta_0) = & ((G_1(X_1, \theta_0) G_1(X_1, \theta_0) + \dots + \\ & | \quad G_1(X_{365}, \theta_0) G_1(X_{365}, \theta_0)) P_1 + \\ & | \quad ((G_1(X_1, \theta_0) G_2(X_1, \theta_0)) + \dots + \\ & | \quad (G_1(X_{365}, \theta_0) G_2(X_{365}, \theta_0)) P_2 + \dots + \\ & | \quad (G_1(X_1, \theta_0) G_9(X_1, \theta_0) + \dots + \\ & | \quad G_1(X_{365}, \theta_0) G_9(X_{365}, \theta_0)) P_9 \quad (3.21) \end{aligned}$$

$$\begin{aligned}
 -\frac{1}{2} Q_9 (X, \theta) = & (G_9 (X_1, \theta) G_1 (X_1, \theta) + \dots + G_9 \\
 & (X_{365}, \theta) \times G_1 (X_{365}, \theta)) P_1 + \dots + \\
 & + (G_9 (X_1, \theta) G_9 (X_1, \theta) + \dots + \\
 & G_9 (X_{365}, \theta) G_9 (X_{365}, \theta)) P_9 \quad (3.22)
 \end{aligned}$$

and where Q_i are determined as:

$$\begin{aligned}
 -\frac{1}{2} Q_1 (X, \theta) = & ((Y(1) - G(X_1, \theta) G_1 (X_1, \theta)) + (Y(2) - G \\
 & (X_2, \theta) \times G_1 (X_2, \theta) + \dots + \\
 & (Y(365) - G(X_{365}, \theta)) (G_1 (X_{365}, \theta))
 \end{aligned}$$

$$\begin{aligned}
 -\frac{1}{2} Q_2 (X, \theta) = & ((Y_1 - G(X_1, \theta)) G_2 (X_1, \theta) + \\
 & (Y_2 - G(X_2, \theta)) G_2 (X_2, \theta) + \dots + \\
 & (Y_{365} - G(X_{365}, \theta)) G_2 (X_{365}, \theta))
 \end{aligned}$$

$$\begin{aligned}
 -\frac{1}{2} Q_9 (X, \theta) = & ((Y_1 - G(X_1, \theta)) (G_9 (X_1, \theta) + \\
 & (Y_2 - G(X_2, \theta)) (G_9 (X_2, \theta) + \dots + \\
 & (Y_{365} - G(X_{365}, \theta)) G_9 (X_{365}, \theta)) \quad (3.23)
 \end{aligned}$$

By this we get systems of linear equations of nine variables namely (P_1, P_2, \dots, P_9) . The determinant of linear equation (3.18) thus can always be solved yielding the elements P_i of the vector P as solutions. Considering the function:

$$Q(V) = Q(X, \theta + Vp), \text{ for } 0 \leq V \leq 1 \quad (3.24)$$

and denoting V the value for which $Q(v)$ is minimum on the interval $0 \leq V \leq 1$. Defining the vector

$$1_0^0 = o_0 + V^0 P \quad (3.26)$$

with elements $1_1^0 i = o_0 i + V^0 P_i \quad (3.26)$

After getting P_i , the next stage is to determine, the minimum value of Q we proceed by an approximate method:- we evaluate Q for $V = 0$, $V = \frac{1}{2}$ and $V = 1$ and determine the level of V for which the parabola through $Q(0)$, $Q(\frac{1}{2})$, $Q(1)$ attains its minimum from

$$V_{\min} = \frac{1}{2} + \frac{1}{4} \frac{(Q(0) - Q(1))}{(Q(1) - 2Q(\frac{1}{2}) + Q(0))} \quad (3.27)$$

In this formula: $Q(0) = \sum_{h=1}^{365} (Y_h - G(X_h, o_0))^2$, and (3.28)

$$Q(\frac{1}{2}) = \sum_{h=1}^{365} (Y_h - G(X_h, \frac{1}{2} \theta_i))^2 \quad (3.29)$$

Where $\frac{1}{2} \theta_i = o_0 i + \frac{1}{2} P_i$, $V = \frac{1}{2}$

$$\text{and } Q(1) = \sum_{h=1}^{365} (Y_h - G(X_h, 1\theta_i))^2 \quad (3.30)$$

Where $1\theta_i = o_0 i + P_i$, $V = 1 \quad (3.31)$

By doing so V_{\min} is calculated, then the next set of constants

$$1\theta_i = o_0 + V_{\min} P_i$$

$$\text{i.e., } I^A = E_0 + V_{\min} P1 \quad (3.32)$$

$$IA1 = A10 + V_{\min} P2$$

$$\begin{array}{c} | \\ | \\ | \\ | \end{array}$$

$$IE = E_0 + V_{\min} P9$$

The whole process is repeated with $(I^d, IA1, IA, IB1, IB, IC1, IC, IE1, IE)$ taking the starting trial place of $o \theta i$. Carrying in this manner, several cycles can be done until the new set of constants has no appreciable difference with the preceding cycle set of constants. In other words having V_{\min} and $(P1, P2 \dots P9)$ to be small.

3.4 Data Analysis and Computation

The determination of new general constants of Sayigh's equation was carried on by writing a computer program based on the above method. The program is attached in appendix (D).

The global solar radiation for the year 1983, of Addis Ababa region, measured by Eppley integrator was used. The corresponding day meteorological parameters, sunshine hours

(measured by campbell stokes sunshine recorder), relative humidity R, and maximum temprature of the day Tmax, all measured by NMA,* were taken in the calculation.

Due to power interuption and malfunction of the recording instruments, in the regression the data available were only for three hundred fifteen (315) days of the year. These values with the corresponding day meteorological parameters are in appendix (E). In the determination of new sets of constants the cycle in which the whole iteration to be repeated was seven times. The set of constants determined by each cycle and solutions of the linear equations (3.21) are shown in table (3.2). The regression was started by arbitrary sets of constants which were close to Sayigh's constants.

The final set of constants which were determined by the last (seventh) cycle are also given in table (3.2). It follows from this, the solar radiation in Addis Ababa is given by the empirical expression:

$$G(X, \theta) = 2.65 \times 10^{-4} K (4.184 \times 10^6) \text{Exp} \{1.46 \phi^{0.52} (0.52 (D)^{0.23} - 0.17R^{0.38} + 0.38 T_{\max}^{0.43})\} \quad (3.33)$$

* NMA - National Meteorological Agency

1.5300	1.0000	0.5000	0.5000	0.3000	0.5x10	0.4000	0.6000	0.4000
-1.5405	0.0334	0.0034	-0.0001	0.0061	-0.0184	0.0001	0.0001	0.0081
-0.5244	1.0446	0.5045	0.4993	0.3082	-0.0196	0.4001	0.6108	0.3992
0.5196	0.0648	-0.0597	0.0025	-0.0759	0.0263	-0.0002	0.0030	0.3993
-0.0679	1.1015	0.4521	0.5021	0.2415	0.0035	0.3999	0.6135	0.3993
0.0579	0.0182	0.1332	0.0041	0.1160	-0.1238	-0.0033	-0.2778	0.0124
-0.0197	1.1167	0.5630	0.5055	0.3381	-0.0995	0.3971	0.3821	0.4103
0.0012	0.2959	-0.0697	0.0244	-0.1849	-0.0251	-0.0258	-0.0060	0.1024
-0.0194	11.1840	0.5455	0.5117	0.2665	-0.1029	0.3907	0.3806	0.4360
0.0146	0.2609	0.2008	-0.0318	0.0584	-0.2551	-0.0159	0.0044	0.0382
-0.0157	1.2489	0.5955	0.5038	0.28100	-0.1664	0.3867	0.3817	0.4456
0.0196	0.2589	0.5955	0.5038	0.0533	-0.0078	-0.0065	-0.0000	-0.0077
2.65×10^{-4}	1.4606	0.5260	0.5294	0.2374	-0.1728	0.3813	0.3817	0.4392
α	A1	A	B1	B	C1	C	E1	E
2.65×10^{-4}	1.46	0.52	0.52	0.23	-0.17	0.38	0.38	0.43

Table(3.2): Shows set of constants and solution of (3.18) for each cycles

Where $G(X, \theta)$, is in $J \cdot m^{-2} \cdot day^{-1}$, ϕ , latitude of the place in degrees, D relative sunshine hours to 12 hrs, R relative humidity, T_{max} maximum temprature of the day and K as it was defined in (3.3).

The daily solar radiation calculated by (3.3), (3.34) and measured values are plotted on figures (3.1) - (3.11). A converting factor for the radiation value from $cal \cdot cm^{-2} \cdot day^{-1}$ to $J \cdot m^{-2} \cdot day^{-1}$ was considered since the measured values are in $J \cdot m^{-2} \cdot day^{-1}$.

Global solar radiation calculated via (3.3) and (3.33) were compared with the measured value and the percentage deviation calculated by,

$$\left(\frac{G_{cal} - G_m}{G_m} \right) \times 100 \quad (3.34)$$

are shown in appendix (F).

The new set of constants has made the maximum temprature of the day to play a dominant role in the calculation of the global solar radiation. A change of $1^\circ C$ in the temprature while the other parameters being constant can result a considerable change in the radiation. The relative sunshine hour (D) is the second important variable in the determination of the radiation. The least important is

Fig. 3.1 DAILY MEASURED AND ESTIMATED GLOBAL SOLAR RADIATION OF ADDIS ABEBA REGION FOR THE MONTH JANUARY 1983.

Δ MEASURED
 ● SATOH'S EQUATION
 ○ NEW CONSTANTS

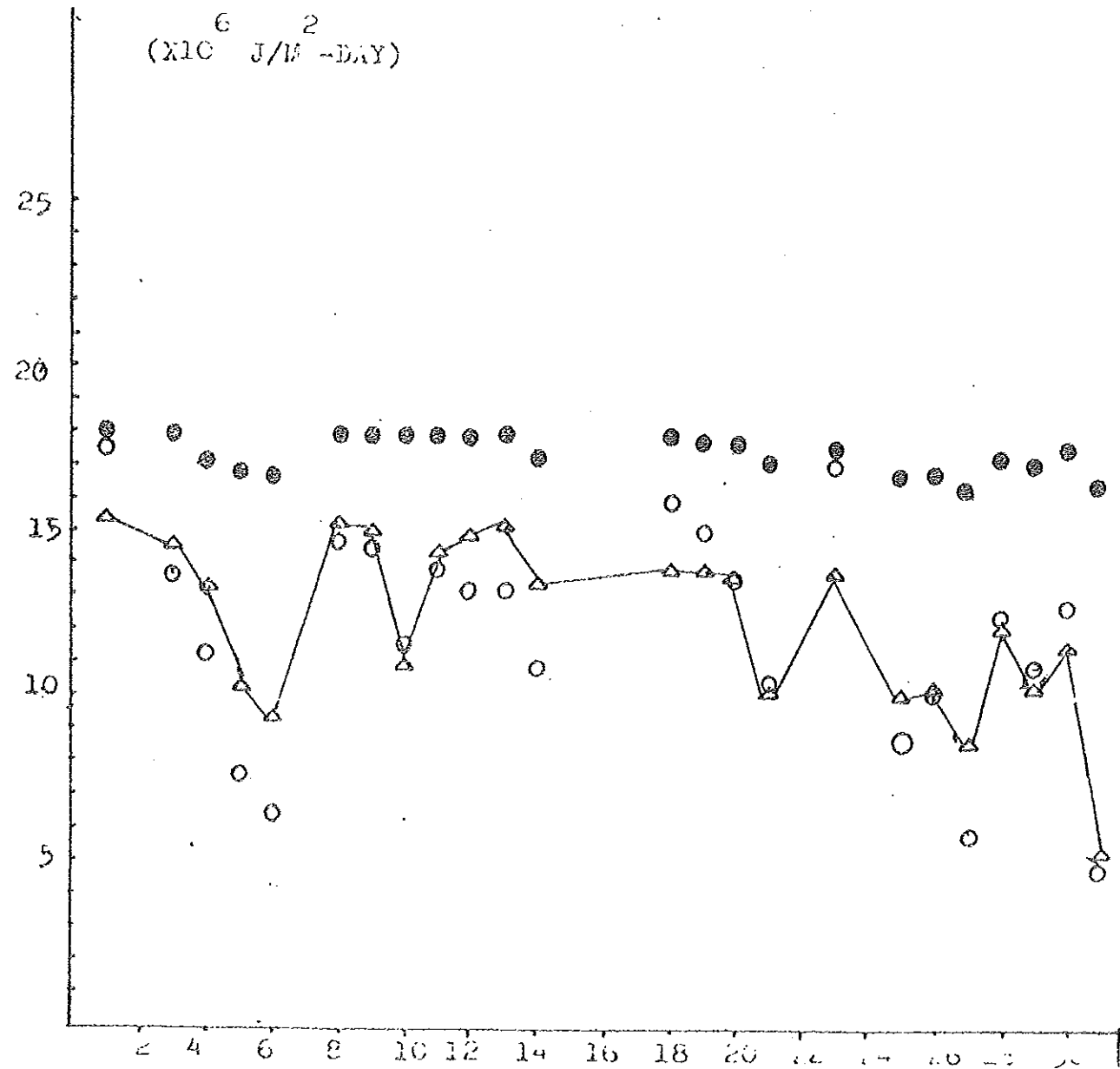


Fig. 3.2 DAILY MEASURED AND CALCULATED SOLAR RADIATION IN LAKH NADIA RADIATION
OF ADDIS ABABA, ETHIOPIA FOR THE MONTH FEBRUARY 1983.

- △ MEASURED
- SAYIGH'S EQUATION
- NEW CONSTANTS

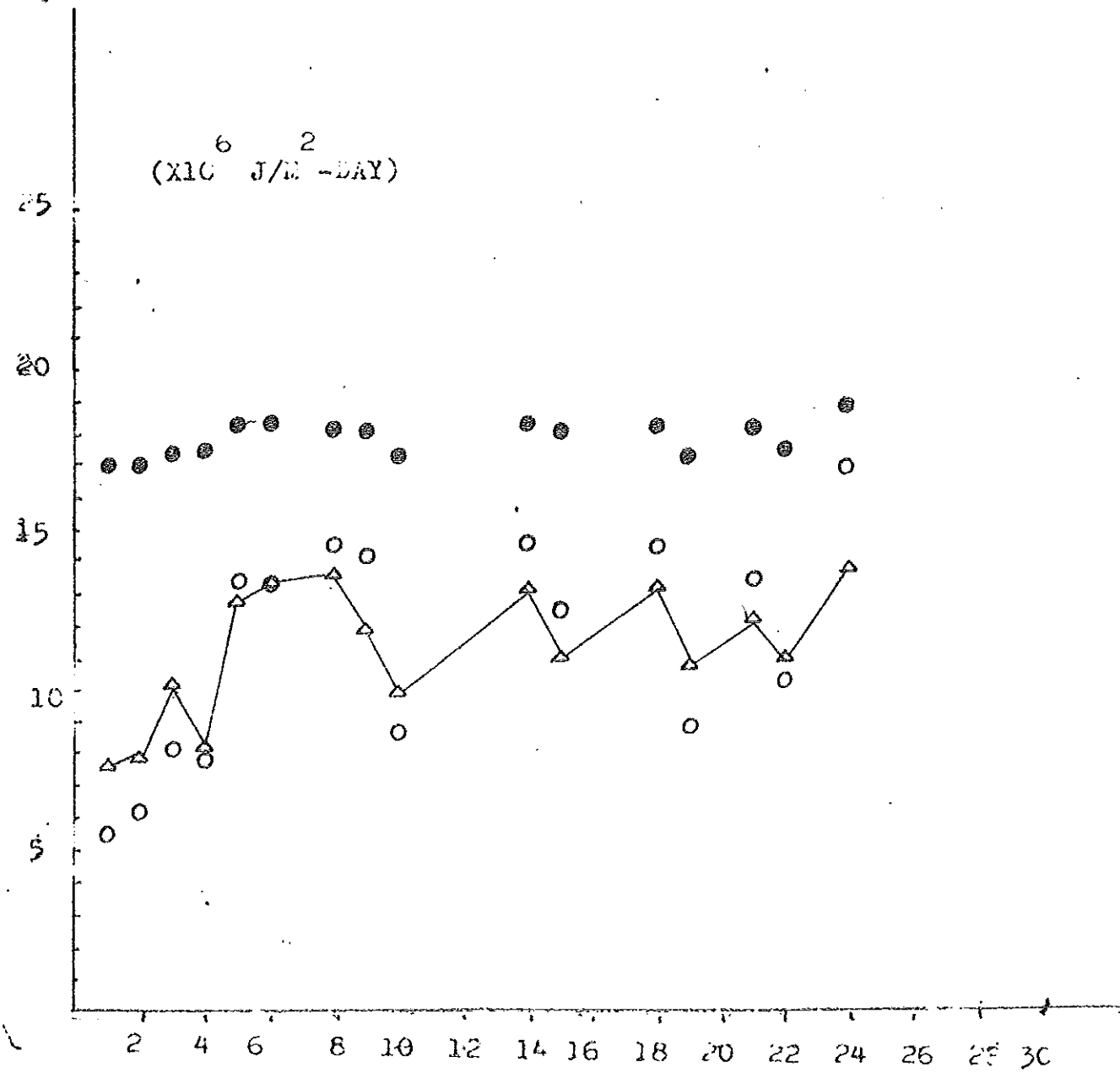


FIG. 3.3 DAILY MEASURED AND ESTIMATED NET LONGWAVE RADIATION
 (1 UNIT IS EQUAL TO 1 W/M²-DAY) MONTH MARCH 1983.

△ MEASURED
 ● DAYLGH'S SURFACE
 ○ NET CONSTANT

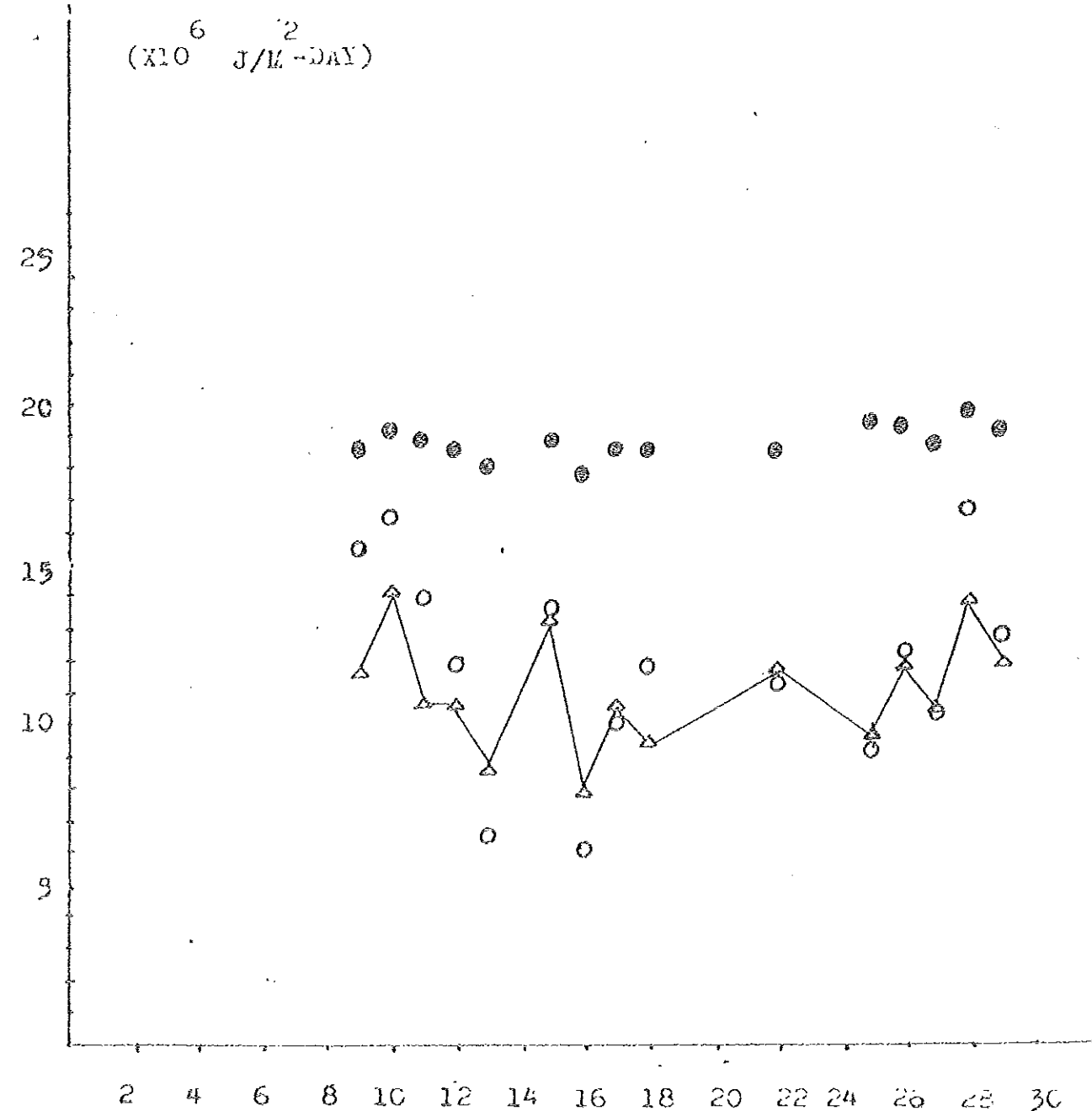


Fig. 3.4 DAILY MEASURED AND ESTIMATED GLOBAL SOLAR RADIATION
OF ADDIS ABABA REGION FOR THE MONTH APRIL 1983.

△ MEASURED
● DAYLIGHT'S EQUATION
○ REF. CONSTANTS

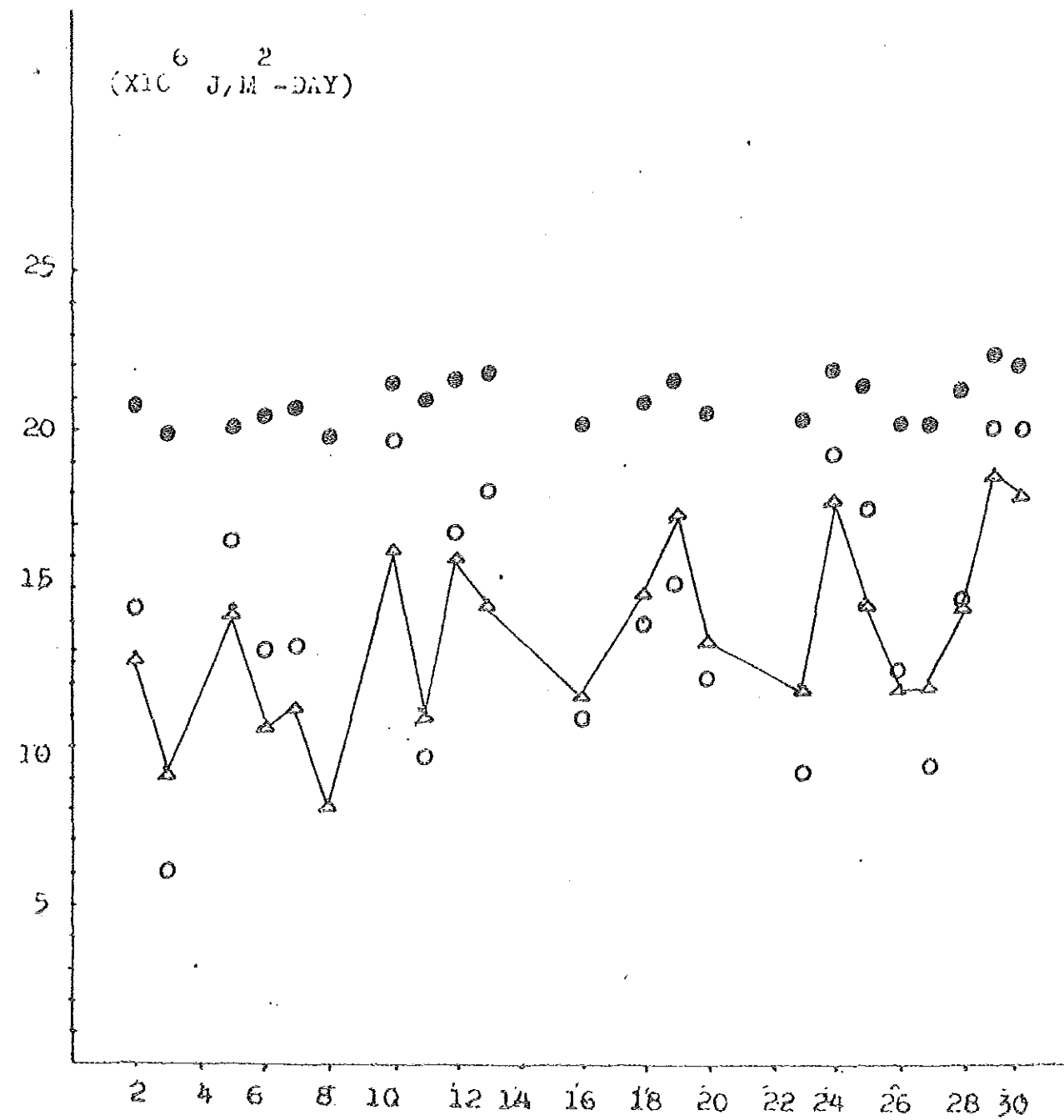


FIG. 5.9 DAILY MEASURED AND CALCULATED GLOBAL SOLAR RADIATION
OF ADDIS ABABA REGION FOR THE MONTH JUNE 1983.

△ MEASURED
● DAYLIGHT'S EQUATION
○ NET CONSTANTS

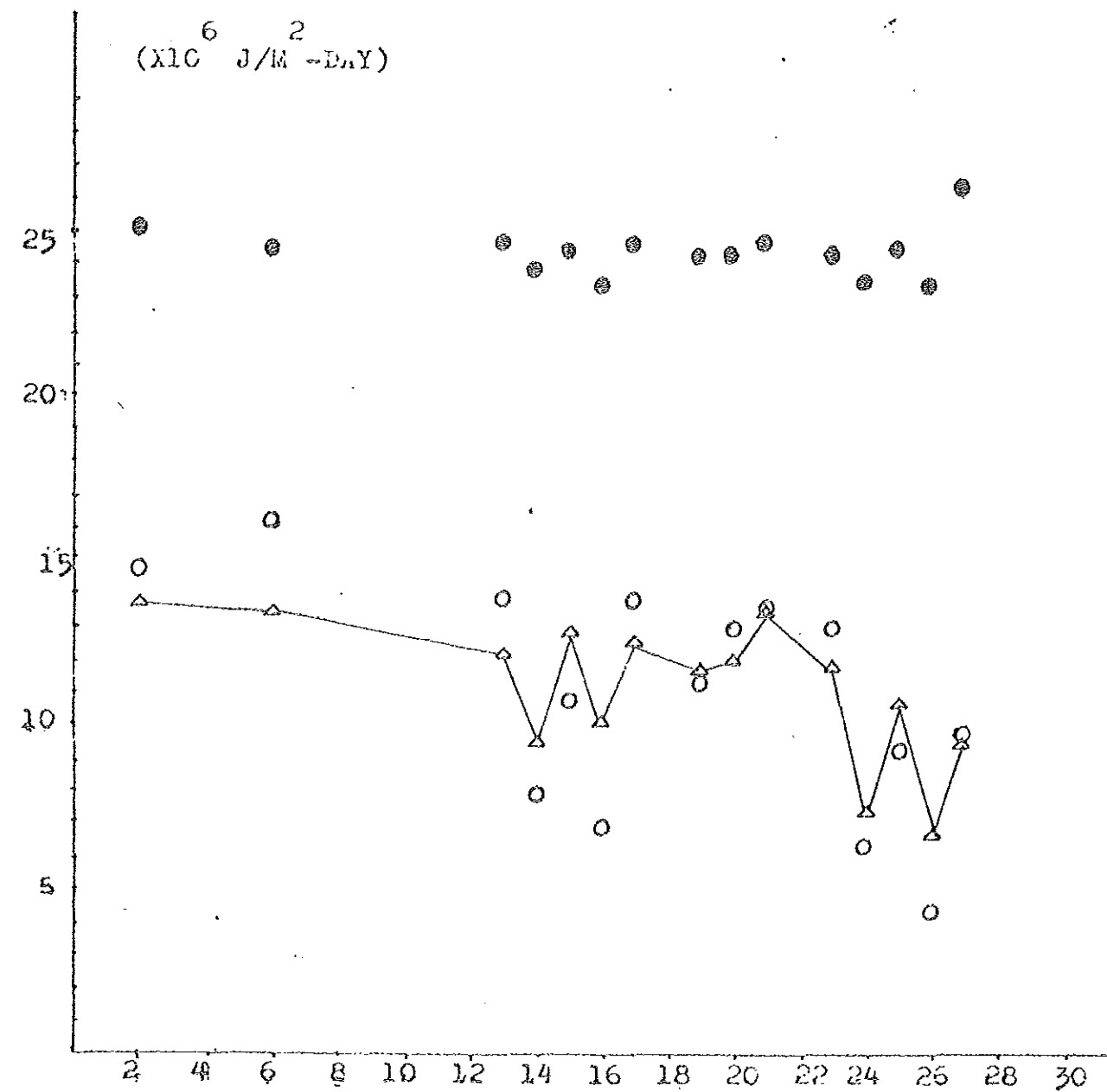


FIG. 3.0 DAILY MEASURED AND ESTIMATED THERMAL SOLAR RADIATION
OF ADOLF HUBER REGION FOR THE MONTH JULY 1983.

△ MEASURED
● SAYIGH'S EQUATION
○ NEW CONSTANTS

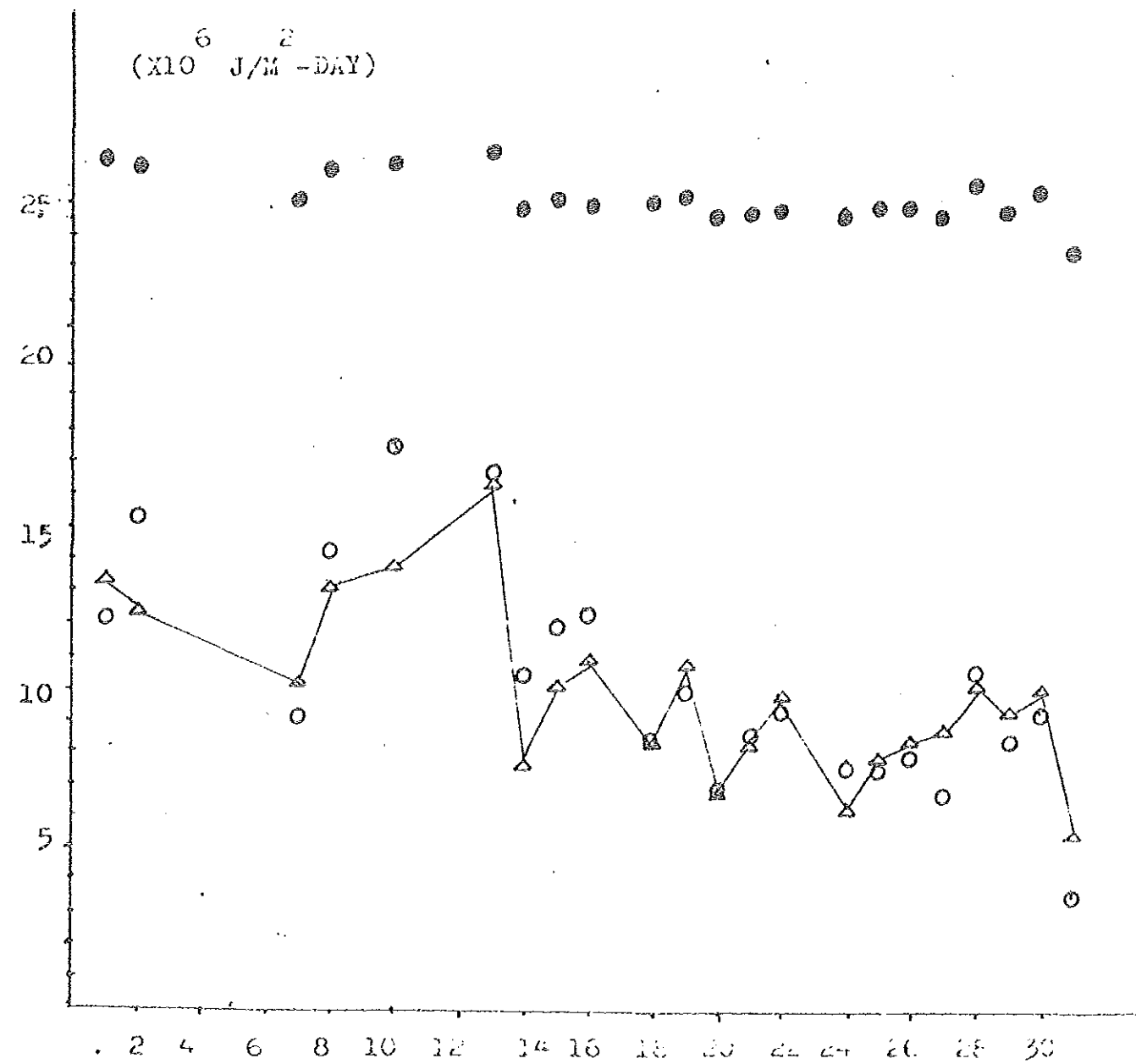


FIG. 5.7 DAILY MEASURED AND ESTIMATED GLACIAL MELT WATER FLOW OF ADLIS AUBA BASIN FOR THE MONTH AUGUST 1983.

▲ MEASURED
 ● SAYIGH'S EQUATION
 ○ NEW CONSTANTS

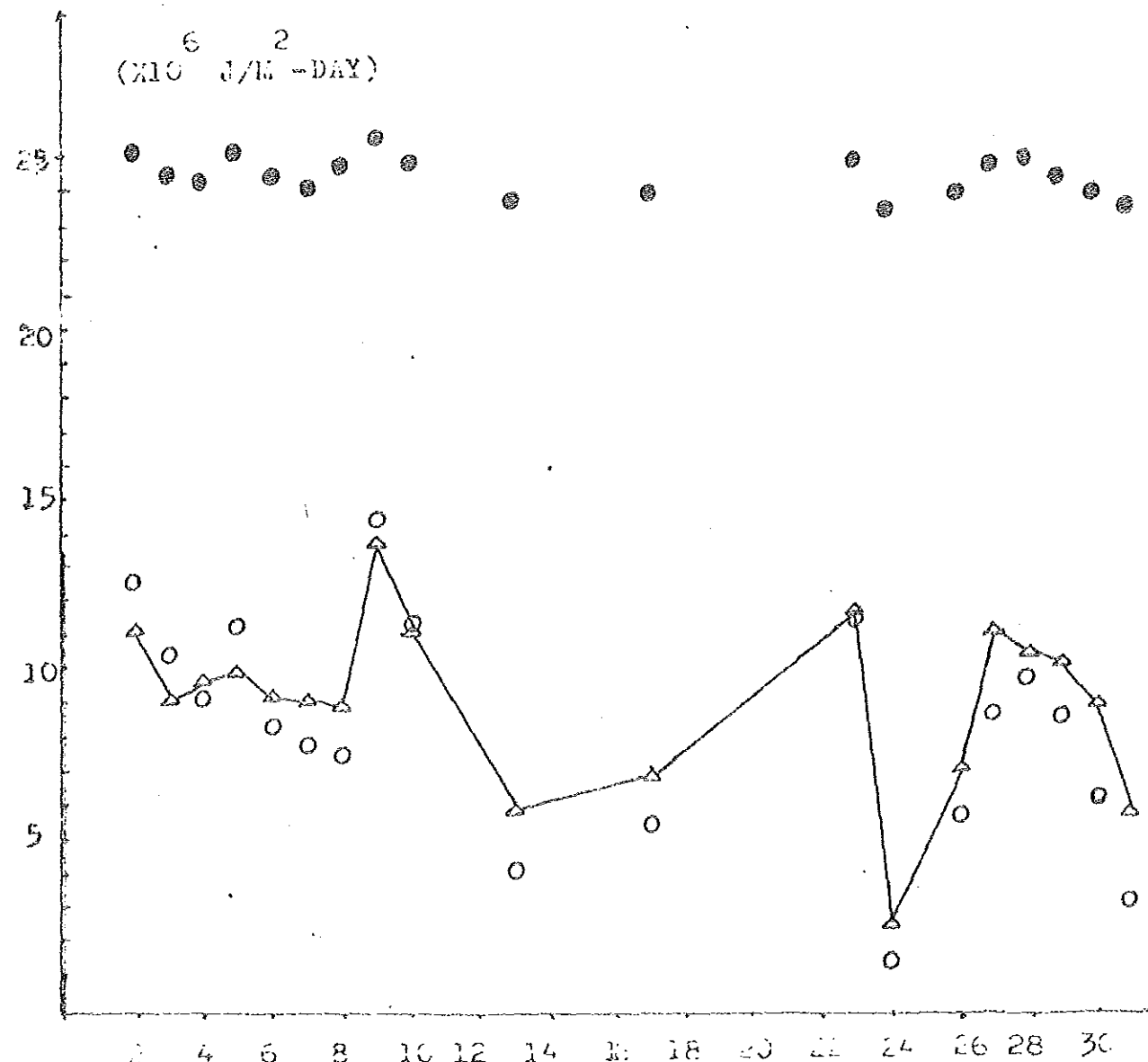


Fig. 2.8 DAILY MEASURED AND ESTIMATED GLOBAL SOLAR IRRADIATION
OF ADDIS ABABA REGION FOR THE MONTH SEPTEMBER 1983.

△ MEASURED
● SAYIGU'S EQUATION
○ NEW CORRECTIONS

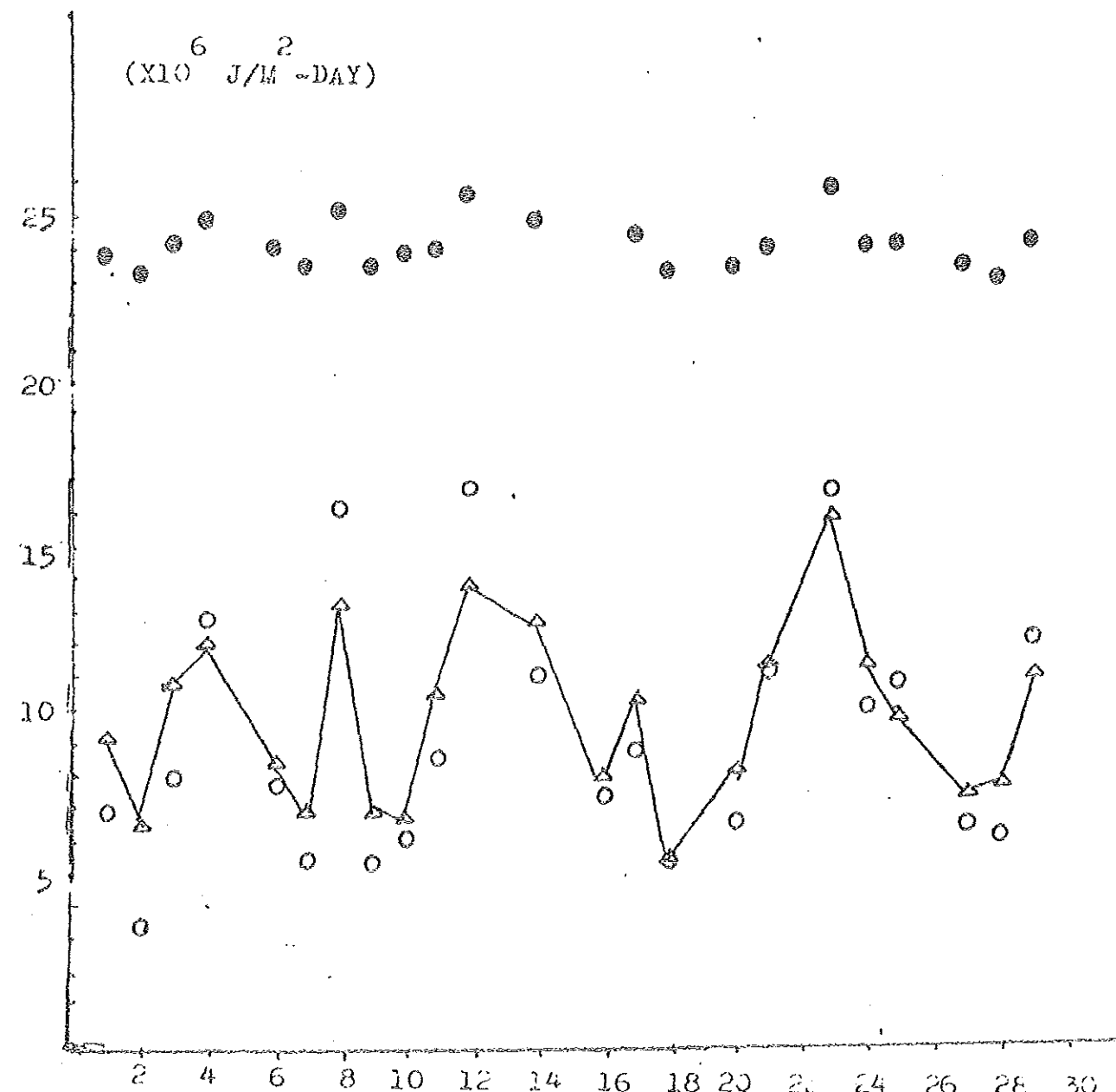


FIG. 3.9 DAILY MEASURED AND ESTIMATED GLOBAL SOLAR RADIATION
OF ADDIS ABABA REGION FOR THE MONTH OCTOBER 1983.

△ MEASURED
● SAYIGH'S EQUATION
○ NEW CONSTANTS

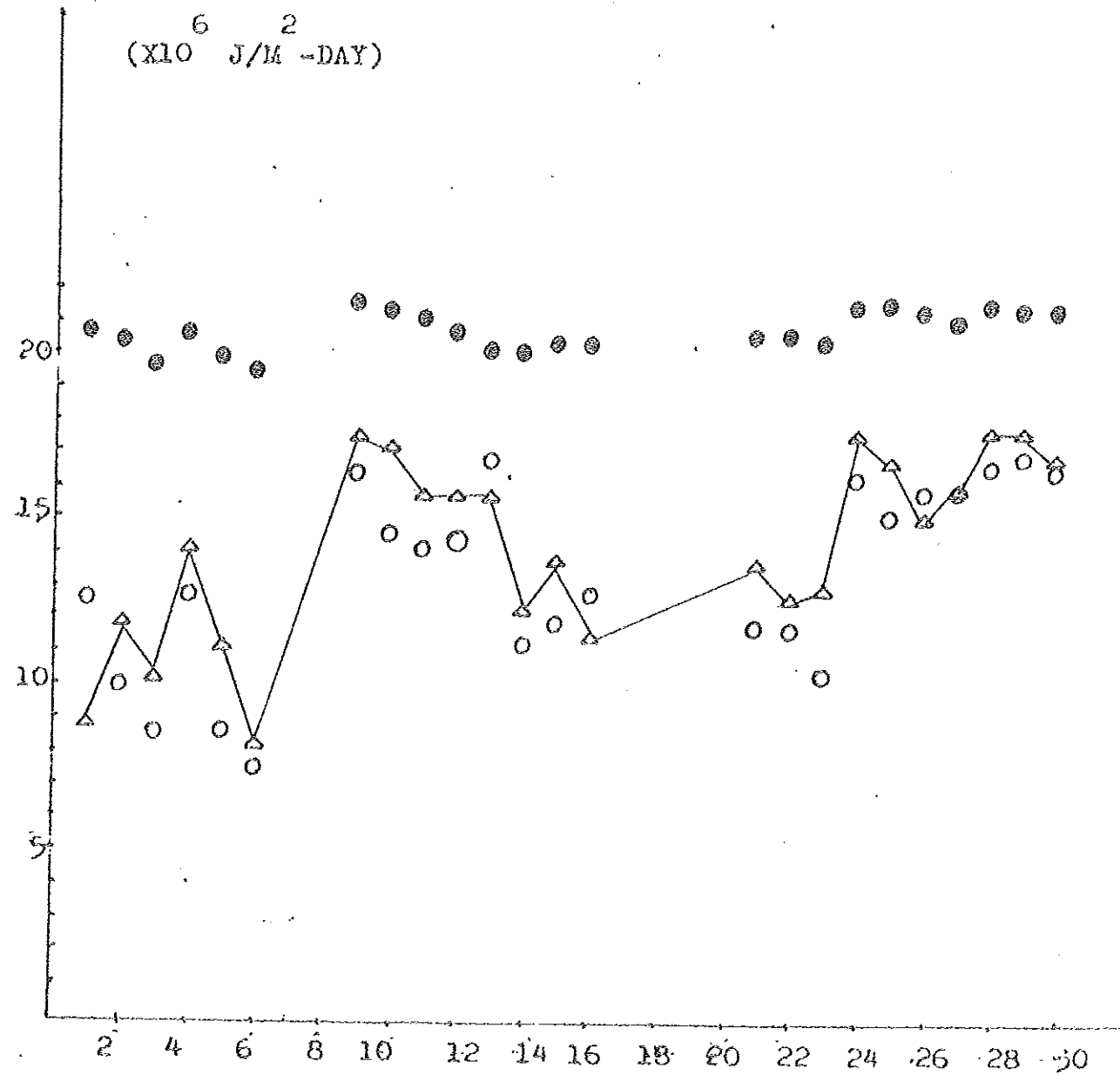


Fig. 10
 DAILY MEASURED AND ESTIMATED GLOBAL SOLAR RADIATION
 ON THE ADDIS ABABA REGION FOR THE MONTH NOVEMBER 1
 983.

△ MEASURED
 ● SAYIGH'S EQUATION
 ○ NEW CONSTATES

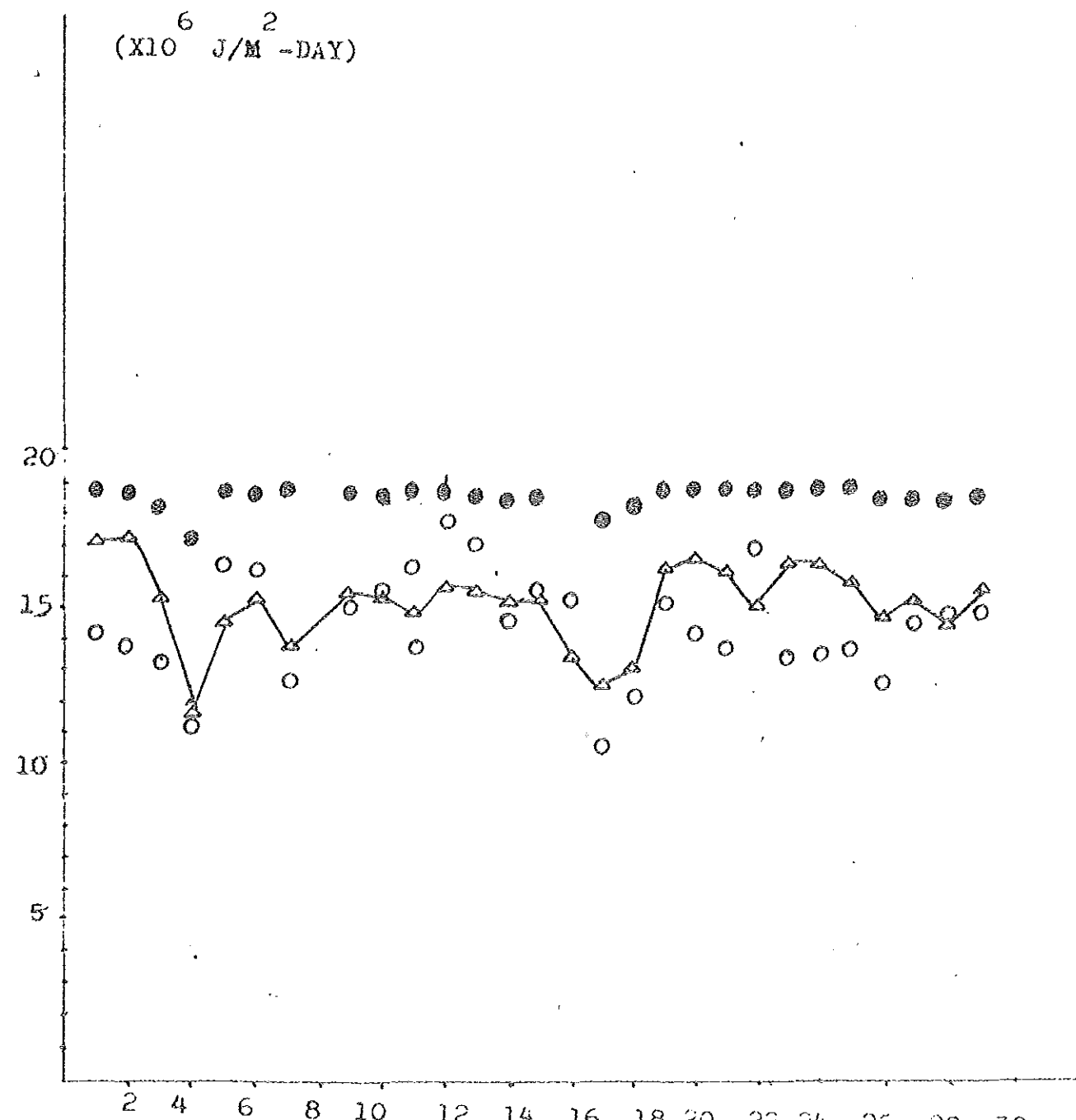
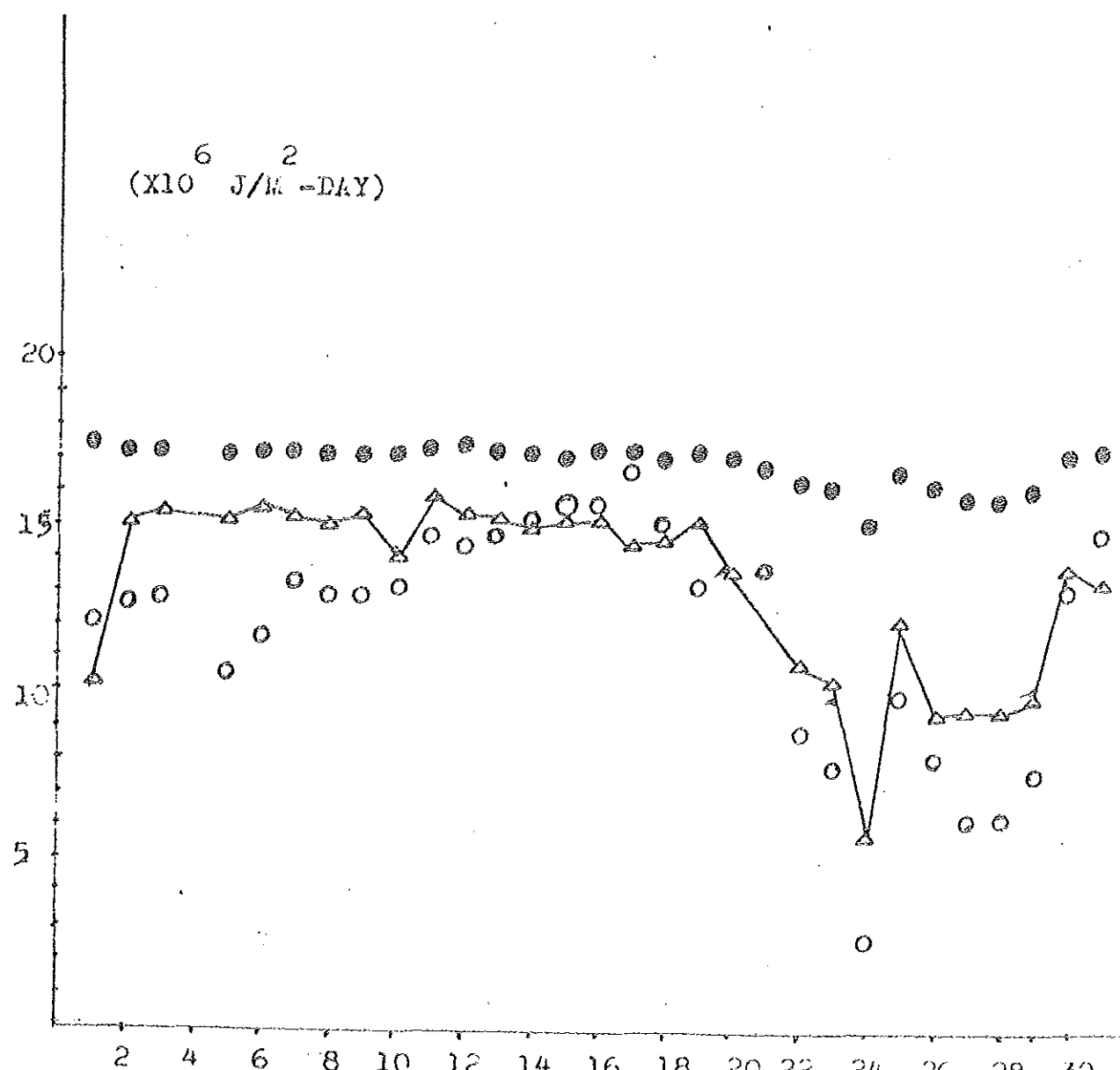


FIG. 11 DAILY MEASURED AND ESTIMATED GLOBAL RADIATION
OF ARDEN ABLEE PATTERN FOR THE MONTH DECEMBER 1983.

△ MEASURED
● SUTCH'S EQUATION
○ NEW CONSTANTS



the relative humidity. Since the constants were determined for one locality, it was not possible to see radiation dependence with latitude.

As it has been done by various authors in calculating, the global solar radiation via (3.1) and (3.2), needs determination of regression coefficients which have spatial variations.

Eventhough, the constants of (3.3) were said to be universal by the authors, the inadequency for estimating the radiation was shown by figures (3.1 - 3.11) and parangotopo et. al. [31] have also reported recently the inadequency of (3.3). And hence it is essential to determine a new set of constants for every region, however it could be possible to use the constants for extrapolating for other regions with the same geography and climate.

CHAPTER 44. CONSTRUCTION OF SIMPLE SOLAR
RADIATION MEASURING DEVICE4.1 Introduction

Solar radiation measurements were originally made in early nineteenth century for the purpose of basic research in radiation physics. This research centered largely on measurements of sunlight intensity and later on the spectral intensity of sunlight. The Smithsonian Institute of Washington D.C. was active with such research.

Radiometers are finding extensive application in the measurement of total and net exchange radiation and for evaluating the thermal efficiency of solar collecting devices. In general, the instruments for measuring solar radiation can be divided into three types, namely thermoelectric, photovoltaic and calorimetric type. In comparison with the thermoelectric and calorimetric type, the photovoltaic has disadvantage in that the spectral response is not linear, even though its response rate is very rapid.

4.2 Solar Radiation Instruments

The most commonly used instruments for determining the available solar energy are 1) Pyrheliometers - for measuring

the intensity at normal incidence of the direct radiation from the solar disk. The sensor of this instrument has a clock driven motor for tracking the sun's direction.

2) Pyranometers - are used for measuring the global radiation (direct plus diffuse) which falls on a horizontal surface. If the sensor is shielded from the direct sunlight it measures the diffuse solar radiation.

4.3 The Duration of Sunshine Instruments:

- Are instruments used for measuring the fraction of time during the day the solar disk is not hidden by clouds. Sunshine duration is important in characterizing the climate of a given location, its value is also used in roughly deducing the flux of solar radiation on a horizontal surface at locations for which no Pyranometric data are available.

These instruments Pyreheliometers, Pyranometers and duration of sunshine instruments, come in many forms by different manufacturing companies and are classified by the commission for instruments and methods of observation of the world meteorological organization (Geneva) as standard, first, second and third class according to a certain criteria. And these were summarized by Kinsell L. Culson [32].

A radiometer was constructed by making use of semi conductor type thermojunctions, but this type of radiometer has less sensitivity. Soldered Constantan Manganin strip was also used. Monteith [32]. But this has a disadvantage of high cost in production.

Simple and in-expensive sensitive radiometers can be obtained in two ways. One way is to use a thermopile material with high thermoelectric power, and another way is to place a thermocouple.

Using thermopile by plating Copper on Constantan has also been used and its design feature and its comparison with standard radiometers developed at central building research, Roorke, was also reported by K.N. Agaraval [35].

However, a simpler radiometer that operates on thermoelectric principle by measuring, heating and cooling rate of the sensor, constructed by S. X. CHENG [36] was adopted in our work. The experimental and theoretical analysis is as follows:

4.4 The Experimental Pyranometer

The Photo of the Experimental Pyranometer is shown in figure (4.1a). A Schematic diagram showing the construction of this Pyranometer is given in figure (4.1b).

PHOTO OF THE EXPERIMENTAL PYRANOMETER.

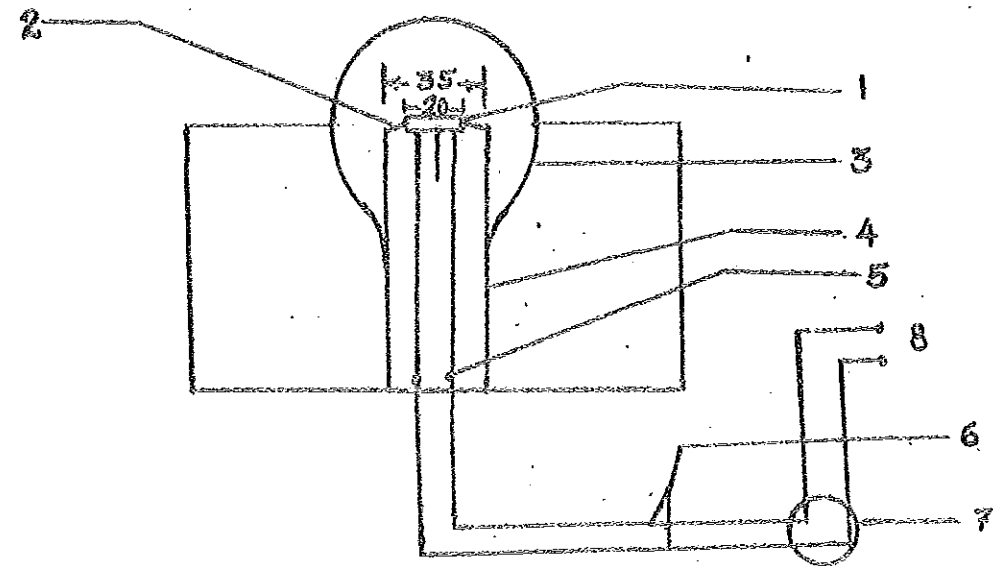
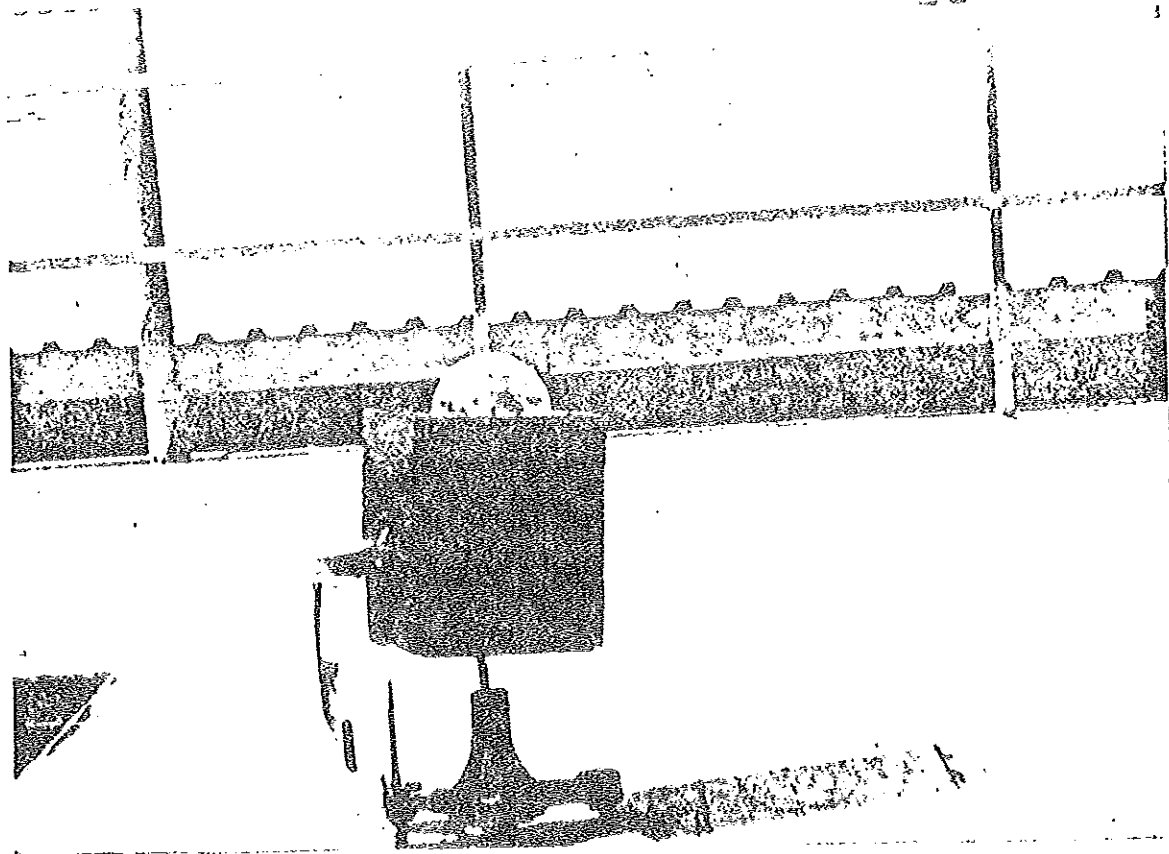


FIG. 4a. SKETCH OF THE CONSTRUCTION OF THE EXPERIMENTAL PYRANOMETER. 1 SENSOR 2 THREAD 3 GLASS COVER 4 SUPPORT TUBE 5 HOLES FOR THE WIRES 6 THERMO COUPLE, 7 ICE POINT, 8 TO CHART RECORDER

The sensor of the Pyranometer is a blackened red copper disk, 0.1 mm thick and 20 mm in diameter. A 0.1 mm constantan and 0.07 mm copper, thermocouple is soldered to the back of the disk for temperature measurements. The sensor is suspended by four threads inside a glass tubing to minimize conduction loss. The sensor is placed at the center of spherical glass cover 140 mm in diameter.

4.5 Theory of the Measuring Technique

The measuring technique presented in this work has made number of assumptions as follows:

1) Considering the heat capacity of the sensor is small and that of the spherical glass cover is large, in a cycle of heating and cooling processes the temperature of the glass cover can be considered to be constant.

2) In heating and cooling process, the temperature of the sensor is uniform at any given time.

Let G be the total solar radiation absorbed per unit time by the sensor when it is exposed to the sun, and hence G be formulated as,

$$G = A \int_0^{\pi} \int_0^{2\pi} \int_0^{\infty} \alpha_{\lambda}(\theta, \phi) I_{\lambda}(\theta, \phi) I_{\lambda}(\theta, \phi) \sin \theta \cos \theta d\theta d\phi d\lambda \quad (4.1)$$

Where I_λ is the spectral intensity of total solar radiation, T_λ is the spectral transmittance of the glass cover to total solar radiation, α_λ is the spectral absorptance of the sensor to total solar radiation, λ is the wavelength, A is area of the sensor, θ and ϕ are the angles of incidence in polar and azimuthal directions respectively.

Knowing that the global radiation is composed of beam and diffuse parts which have different spectral distributions, it is necessary to consider the contributions refined as follows:

$$G = \alpha_d T_d G_d A + \alpha_b (\theta_o, \phi_o) \tau_o (\theta_o, \phi_o) G_b C_o S \theta_o A. \quad (4.2)$$

Where the first term on the right of the equation denotes, the diffuse energy absorbed by the sensor, the second term is the beam radiation absorbed at θ_o, ϕ_o directions, G_d is diffuse irradiance. When the radiation is incident on the blackend sensor, the energy balance of the sensor can be written as:

$$G = C_p M \left(\frac{dT}{dt}\right)_h + Q_{rad} + Q_{conv} + Q_{cond} \quad (4.3)$$

Where M and C_p are specific heat and mass of the sensor respectively. $\left(\frac{dT}{dt}\right)_h$ is the temperature slope at any given

time during the heating process and Q_{rad} , Q_{conv} , and Q_{cond} represent the heat loss by radiation, convection, and conduction respectively.

When the sensor is shielded from the radiation the sensor undergoes a cooling process. The following energy balance equation can be written for this process.

$$-C_p M \left(\frac{dT}{dt}\right)_c = C_p M \left|\left(\frac{dT}{dt}\right)_c\right| = Q'_{\text{rad}} + Q'_{\text{conv}} + Q'_{\text{cond}} \quad (4.4)$$

Where $\left(\frac{dT}{dt}\right)_c$ is the temperature slope at any given time during cooling process, Q'_{rad} , Q'_{conv} and Q'_{cond} represent the heat loss for the cooling process.

The temperature-time history of the sensor can be plotted as shown in figure (4.2). Regardless of the heating and cooling process, the heat loss of the sensor at points A and B is identical, because the sensor is at the same temperature.

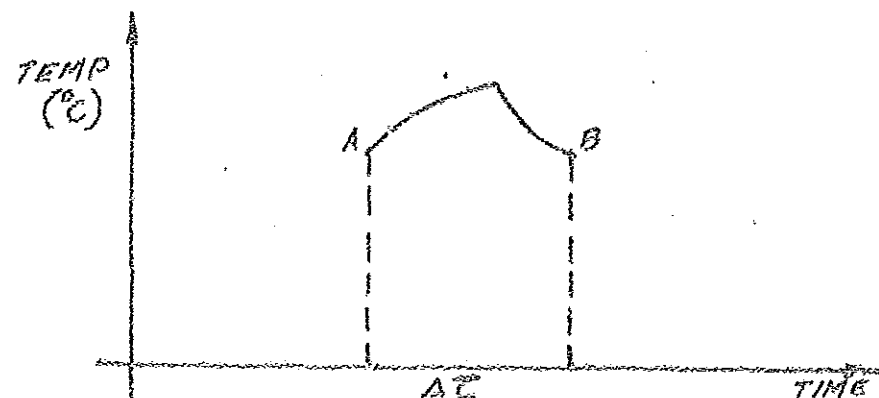


Fig. (4.2) Temperature - time history of the sensor in the heating and cooling process.

This permits writing:

$$Q_{\text{rad}} = Q'_{\text{rad}}$$

$$Q_{\text{conv}} = Q'_{\text{conv}} \quad (4.5)$$

$$Q_{\text{cond}} = Q'_{\text{cond}}$$

These equations enable the heat loss terms in (4.3) and (4.4) to be cancelled out there by resulting:

$$\alpha_d \tau_d G_d A + \alpha_b (\theta_o, \alpha_o) \tau_b (\tau_o, \phi_o) G_b C_o S \theta_o A =$$

$$C_p M \left(\left(\frac{dT}{dt} \right)_b + \left(\left| \frac{dT}{dt} \right| \right)_c \right) \quad (4.6)$$

Using the following relationships to account for the difference in materials spectral response to beam and diffuse radiation

$$\alpha_d - \alpha_b = \pm \Delta \alpha \text{ and } \tau_d - \tau_b = \pm \Delta \tau \quad (4.7)$$

Then (4.6) can be written as:

$$G_d + G_b \cos \theta_o = \left(\frac{C_p M}{A \alpha_d \tau_d} \right) \left(\left(\frac{dT}{dt} \right)_b + \left(\left| \frac{dT}{dt} \right| \right)_c \right) \pm$$

$$\left(\frac{\Delta \alpha}{\alpha_d} + \frac{\Delta \tau}{\tau_d} \right) G_b \cos \theta_o \quad (4.8)$$

or,

$$G_d + G_b \cos \theta_o = \left(\frac{C_p M}{A \alpha_b \tau_b} \right) \left(\left(\frac{dT}{dt} \right)_b + \left(\left| \frac{dT}{dt} \right| \right)_c \right) \pm$$

$$\left(\frac{\Delta \alpha}{\alpha_b} + \frac{\Delta \tau}{\tau_b} \right) G_d \quad (4.9)$$

The physical meaning of (4.8) and (4.9) can be explained as, if there is only diffuse radiation that is $G_d = 0$, it follows from (4.8) that:

$$G_d = \left(\frac{C_D M}{A \alpha_d \tau_d} \right) \left(\left(\frac{dT}{dt} \right)_h + \left(\left| \frac{dT}{dt} \right|_e \right) \right) \quad (4.10)$$

(4.10) Can be used to determine the diffuse radiation, if the sensor is shielded from the direct radiation.

If there is direct solar radiation, that is $G_d = 0$, (4.9) is simplified to be:

$$G_b = \left(\frac{C_D M}{A C_s \sin \theta \alpha_b \tau_b} \right) \left(\left(\frac{dT}{dt} \right)_h + \left(\left| \frac{dT}{dt} \right|_e \right) \right)$$

$$G_b = \left(\frac{C_D M}{A_B \alpha_b \tau_b} \right) \left(\left(\frac{dT}{dt} \right)_h + \left(\left| \frac{dT}{dt} \right|_e \right) \right) \quad (4.11)$$

Where A_B is the projection area of the sensor in the direct radiation direction. If the assumptions are further made, $\Delta \ll \tau_d$ and τ_b , and $\Delta \ll \alpha_d$ and α_b , the second terms of (4.8) and (4.9) are neglected. And hence,

$$G = \left(\frac{C_D M}{\tau_b \alpha_b A} \right) \left(\left(\frac{dT}{dt} \right)_h + \left(\left| \frac{dT}{dt} \right|_e \right) \right) \quad (4.12)$$

$$\text{Letting } K = \frac{C_D M}{\tau_b \alpha_b A} \quad (4.13)$$

For a given sensor and a glass cover, k is a constant. The calibration of k can be carried out by either of the following methods:

a) C_p , M , α_b and τ_b are accurately measured in advance. In this case the pyranometer becomes an absolute radiometer.

b) By means of standard pyranometer, the constant k can be calibrated.

Measured Result

After relatively calibrating the instrument by means of method b, several comparison measurements had been taken. The global solar radiation data measured by the newly constructed instrument and local meteorological station are shown on table (4.1)

1. LOC. P.	285	366	338	355	325	302	396
2. EXP. P.	328	420	350	364	270	245	318
% deviation	16	17	-0.003	-2	-16	-18	-19

Table (4.1) 1 Local station, 2 Experimental Pyranometer
Measured Radiation in $W \cdot M^{-2}$

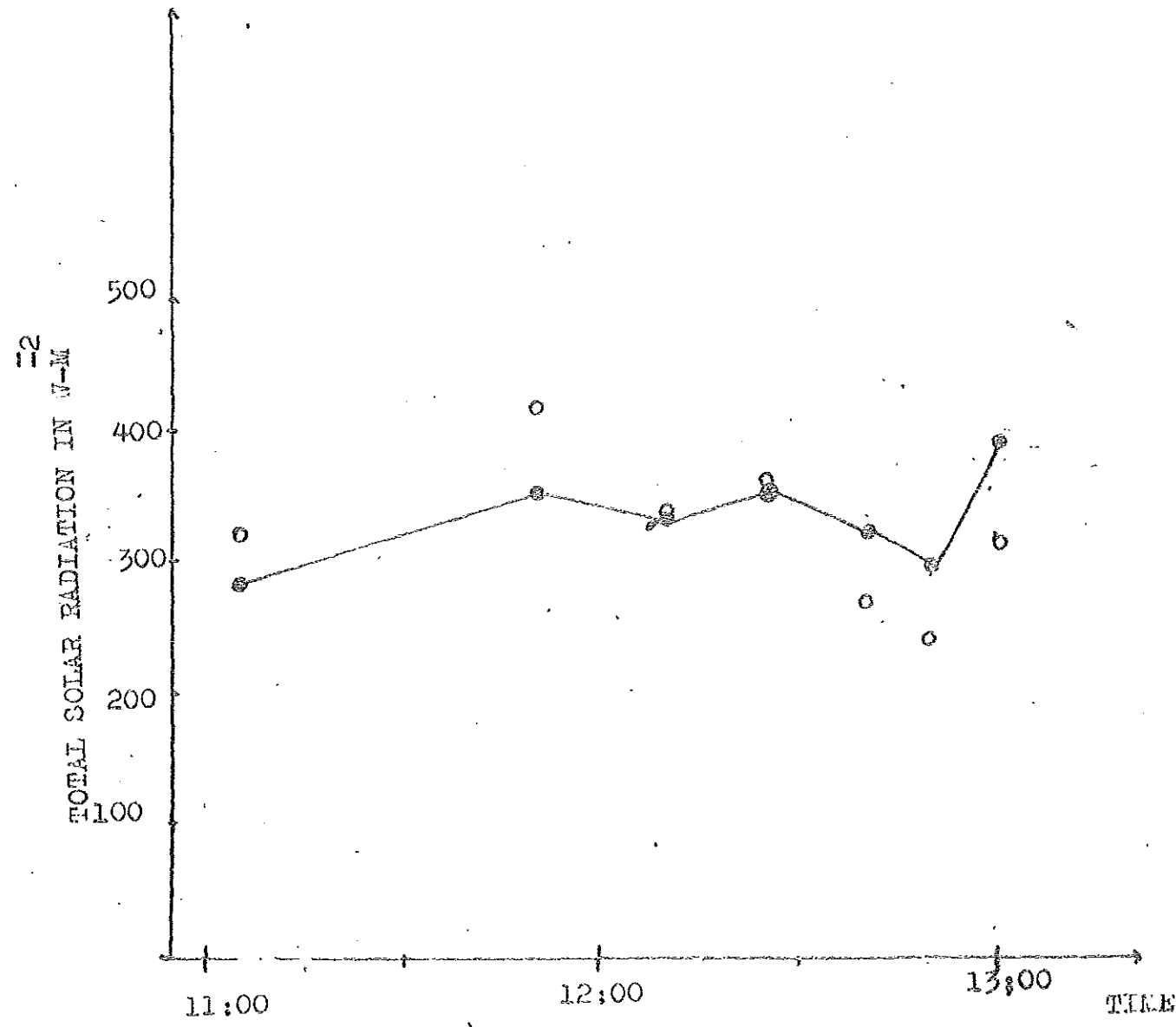


FIG.4.3 MEASURED RESULTS (MAY 23, 1985) BY THE EXPERIMENTAL PAVANOMETER 8; O MEASURED BY LOCAL METEOROLOGICAL STATION. (MAY 23, 1985)

The measurement was taken during (12.00 am. - 14.00 pm.) on may 23, 1985. These values are plotted on figure (4.3). The maximum relative deviation of the measured data from that of a local meteorological station is about 19 percent.

Conclusion

It is important to know the attenuation of solar radiation, for a particular place, and which attenuation mechanism affects the radiation more. Specially the integral turbidity factor, which predicts the solar radiation attenuation via (2.51) is very important. As it has been seen in Figure (2.4) the maximum monthly mean value is in August, which is by 50% less than that of great Britain; where we do have man made aerosol pollution. In addition, these molecularly attenuated radiation and integral turbidity factors are important in meteorological predictions.

In the estimation of global solar radiation, the newly determined set of constants would have betterly estimated the radiation for Addis Ababa, had we used several consecutive years of data. However the discrepancies between the measured and estimated, have reduced considerably, and these constants can be used for estimating the future years global solar radiation of Addis Ababa region.

For routine measurement of global solar radiation, the instrument in discussion of Chapter Four, can be produced cheaply and easily.

These types of instruments can be used to replace the expensive ones and can also be installed, in places where standard Pyranometers are not available. These instruments can also be made as Pyrheliometer, by mounting the sensor on clock driven motor for tracking the Sun's disc. The experimental instrument's output can also be directly connected to a micro-computer, having inner built solar radiation analysis program, to process the solar radiation measurement.

APPENDIX - A

INPUT DATA OF DIRECT SOLAR RADIATION, D(I), PRESSURE, P(I) AND DECLINATION ANGLE, DEL(I) OF A.A. RE-100
 FOR THE YEAR 1953.

DAY	D(I)	E(I)	P(I)	DEL(I)	DA	(I)	(I)	(I)	(I)	DAY	D(I)	(I)	P(I)
1	5373	0.9667	575.0	-23.0607	3	5107	0.9667	578.0	-22.8913	1	4371	0.9667	576.3
5	2380	0.9669	576.5	-22.6949	6	857	0.9669	577.3	22.5365	5	3330	0.9669	578.0
8	5571	0.9669	577.1	-22.3498	9	5407	0.9669	577.8	-22.2215	8	3411	0.9671	576.3
11	3686	0.9671	577.5	-21.9451	12	4326	0.9671	576.9	-1.7172	3	4932	0.9621	575.0
14	886	0.9670	575.0	-21.4820	19	304	0.9677	574.5	-20.5834	7	3859	0.9682	575.0
21	1481	0.9682	577.0	-20.1809	23	3192	0.9682	575.0	-19.7545	24	141	0.962	572.0
25	466	0.969	578.0	-19.3046	26	1056	0.9693	577.0	-19.0711	27	708	0.9693	578.0
28	1394	0.9693	577.	-18.5872	29	2381	0.9693	578.0	-18.3369	30	2546	0.9693	577.0
31	90	0.9693	577.0	-17.821	1	505	0.9708	577.0	-17.5538	2	21	0.9708	578.0
3	8	0.970	578.0	-17.0056	4	4100	0.9708	579.0	-16.7239	5	4392	0.9723	578.0
6	4952	0.9723	579.0	-16.1457	7	3937	0.9723	578.0	-15.8494	8	2883	0.9723	576.3
9	1711	0.9723	575.0	-15.247	10	4004	0.9738	574.0	-14.9326	11	1978	0.9738	576.0
12	2469	0.9738	577.1	-14.2991	13	1685	0.9738	576.0	-13.9760	14	570	0.9738	577.0
15	633	0.9757	576.0	-13.3174	16	500	0.9757	577.0	-12.9821	17	3208	0.9757	575.0
18	1389	0.9757	576.0	-12.3002	20	2342	0.9777	577.0	-11.6036	21	1830	0.9777	578.0
22	2450	0.9777	579.	-10.8933	23	5499	0.9777	579.0	-10.5333	24	3338	0.9777	578.0
25	5451	0.9798	578.0	-9.8019	27	301	0.9798	578.0	-9.0530	1	4595	0.9819	578.0
3	3060	0.9819	579.0	-7.5497	4	4210	0.9819	578.0	-7.1655	5	2925	0.9842	578.0
6	2827	0.9842	577.0	-6.3909	7	3033	0.9842	577.0	-6.0007	8	2753	0.9842	578.0
9	2050	0.9842	578.0	-5.2150	10	3441	0.9864	578.0	-4.8199	11	1864	0.9864	577.0
12	1234	0.9864	577.0	-4.0253	13	544	0.9864	578.0	-3.6262	14	7	0.9864	577.0
15	773	0.9892	576.0	-2.8248	16	174	0.9892	577.0	-2.4228	17	1323	0.9892	577.0
18	327	0.9892	578.0	-1.6167	19	375	0.9892	577.0	-1.2129	20	195	0.9920	578.0
21	4106	0.9920	577.0	-0.4044	22	1894	0.9920	576.0	0.0000	23	3431	0.9920	577.0
24	5090	0.9920	578.0	0.8089	25	1081	0.9950	578.0	1.2131	26	2023	0.9950	576.0
27	1144	0.9950	577.0	2.0202	28	1120	0.9950	577.0	2.4229	29	476	0.9950	576.0
30	2985	0.9950	575.0	3.2261	31	1780	0.9950	576.0	3.6263	1	2858	0.9986	576.0
2	3394	0.9986	577.0	4.4233	3	1076	0.9986	577.0	4.8200	4	838	0.9986	576.0
5	2438	1.0013	575.0	5.6088	6	655	1.0013	577.0	6.0008	7	490	1.0013	576.0
8	126	1.0013	577.0	6.7793	9	35	1.0013	576.0	7.1657	11	36	1.0040	576.0
12	3566	1.0040	578.0	8.3114	13	3390	1.0040	577.0	8.6385	14	1611	1.0040	578.0
15	2898	1.0067	577.0	9.4350	16	1612	1.0067	578.0	9.8041	17	23	1.0067	578.0
18	3955	1.0067	578.0	10.5334	19	4623	1.0067	578.0	10.8934	20	2184	1.0096	576.0
21	1075	1.009	578.0	11.2503	22	23	1.0096	577.0	11.6037	23	424	1.0096	575.0
24	4895	1.0096	577.0	12.3003	25	3387	1.0125	578.0	12.6431	26	570	1.0125	577.0

INPUT DATA OF DIRECT SOLAR RADIATION, D(I), PRESSURE, P(I), AND DECLINATION ANGLE, DEL(I) OF A.A. KOGI FOR THE YEAR 1983.

DAY	D(I)	E(I)	P(I)	DEL(I)	DAY	D(I)	E(I)	P(I)	DEL(I)	DAY	D(I)	E(I)	P(I)
27	1655	1.0125	578.0	13.3175	28	3156	1.0125	578.0	13.6488	29	6437	1.0125	577.0
30	5034	1.0125	577.0	14.2992	1	3522	1.0155	577.0	14.6181	2	3796	1.0155	576.0
3	1353	1.0155	575.0	15.2428	4	915	1.0155	578.0	15.5484	5	305	1.0177	577.0
6	4934	1.0177	577.0	16.1457	7	4604	1.0177	577.0	16.4373	8	2283	1.0177	578.0
9	1754	1.0177	577.0	17.0056	10	4646	1.0199	578.0	17.2823	11	4016	1.0199	576.0
12	3908	1.0199	576.0	17.8202	13	3482	1.0199	575.0	18.0813	14	2070	1.0199	576.0
15	2168	1.0220	575.0	18.5872	16	2599	1.0220	576.0	18.8320	17	1037	1.0220	576.0
18	1641	1.0284	576.0	21.9452	2	674	1.0284	577.0	22.0866	3	4626	1.0284	575.0
21	4710	1.0284	576.0	22.3499	5	3154	1.0296	575.0	22.4715	6	1471	1.0296	576.0
24	3875	1.0296	576.0	22.6949	8	4407	1.0296	578.0	22.7965	9	2765	1.0296	577.0
27	1424	1.0308	576.0	22.9794	11	3996	1.0308	577.0	23.0607	12	4000	1.0308	576.0
30	1465	1.0308	578.0	23.2027	14	5337	1.0308	578.0	23.2634	15	2946	1.0317	577.0
1	1601	1.0317	577.0	23.3641	17	1319	1.0317	579.0	23.4041	18	2370	1.0317	572.0
4	1351	1.0317	578.0	23.4632	20	1781	1.0327	578.0	23.4324	21	2188	1.0327	578.0
7	1748	1.0327	577.0	23.4998	23	1798	1.0327	577.0	23.4981	24	2746	1.0327	578.0
10	2721	1.0332	577.0	23.4737	26	1782	1.0332	578.0	23.4510	27	1165	1.0332	577.0
13	2120	1.0332	576.0	23.3850	29	305	1.0332	576.0	23.3416	30	350	1.0332	578.0
16	2578	1.0337	578.0	23.2339	2	2119	1.0337	578.0	23.1698	3	3366	1.0337	576.0
19	1797	1.0337	578.0	23.0209	5	3810	1.0337	577.0	22.9362	6	2602	1.0337	577.0
22	903	1.0337	577.0	22.7465	8	2059	1.0537	577.0	22.6416	9	3002	1.0337	578.0
25	2708	1.0337	577.0	22.4115	11	2709	1.0337	577.0	22.2865	12	1699	1.0337	578.0
28	2271	1.0337	577.0	22.0167	14	989	1.0337	576.0	21.8720	15	1250	1.0331	575.0
31	1137	1.0331	575.0	21.5632	17	1157	1.0331	577.0	21.3992	18	917	1.0331	576.0
1	1585	1.0331	577.0	21.0522	20	537	1.0324	577.0	20.8693	21	5	1.0324	578.0
4	8	1.0324	576.0	20.4851	23	343	1.0324	577.0	20.2838	24	478	1.0324	578.0
7	1032	1.0316	576.0	19.8633	26	608	1.0316	577.0	19.6442	27	477	1.0316	577.0
10	1411	1.0316	576.0	19.1886	29	1131	1.0316	577.0	18.9523	30	977	1.0316	578.0
13	69	1.0316	578.0	18.4620	1	294	1.0300	577.0	18.2098	2	1781	1.0300	576.0
16	347	1.0300	577.0	17.6870	4	3	1.0300	578.0	17.4187	5	681	1.0288	577.0
19	1381	1.0288	578.0	17.1446	7	742	1.0288	578.0	16.8654	8	4	1.0288	576.0
22	2603	1.0288	578.0	16.2921	10	2614	1.0272	576.0	15.9982	11	57	1.0272	577.0
25	190	1.0272	577.0	15.3961	13	13	1.0212	577.0	15.0883	14	50	1.0272	577.0
28	1	1.0257	577.0	14.4592	16	87	1.0257	576.0	14.1381	17	1064	1.0257	577.0

WAVELENGTH(LAM), RAYLEIGH MOLECULAR ATTENUATION COEFFICIENT THOU(LAM), EXTRATERRESTRIAL SOLAR RADIATION I(L) FOR THE SPECTRUM IN THE INTERVAL (290 - 4000)NM

WL(LAM)	THOU(LAM)	IO(LAM)	WL(LAM)	THOU(LAM)	IO(LAM)
0.290	0.141100E 01	5.240	0.300	0.122000E 01	6.180
0.310	0.106000E 01	6.350	0.320	0.962000E 00	7.810
0.330	0.812700E 00	9.000	0.340	0.716400E 00	8.940
0.350	0.634100E 00	9.490	0.360	0.563400E 00	10.510
0.370	0.502300E 00	10.400	0.380	0.449400E 00	9.450
0.390	0.403300E 00	11.340	0.400	0.363000E 00	16.310
0.410	0.327700E 00	17.000	0.420	0.296600E 00	16.590
0.430	0.269100E 00	16.720	0.440	0.244700E 00	19.280
0.450	0.223100E 00	20.060	0.460	0.203800E 00	19.860
0.470	0.186500E 00	19.890	0.480	0.171100E 00	18.880
0.490	0.157200E 00	19.560	0.500	0.144700E 00	19.020
0.510	0.133500E 00	18.310	0.520	0.123300E 00	18.590
0.530	0.114000E 00	19.170	0.540	0.105700E 00	18.560
0.550	0.980500E-01	18.410	0.560	0.911000E-01	18.280
0.570	0.847700E-01	18.340	0.580	0.789700E-01	18.080
0.590	0.736700E-01	17.630	0.600	0.688000E-01	17.410
0.610	0.643300E-01	17.050	0.620	0.602200E-01	16.580
0.630	0.564300E-01	16.330	0.640	0.529400E-01	15.990
0.650	0.497100E-01	15.200	0.660	0.467300E-01	15.550
0.670	0.439700E-01	15.160	0.680	0.414100E-01	14.890
0.690	0.390300E-01	14.500	0.700	0.368200E-01	14.160
0.710	0.347700E-01	13.850	0.720	0.328600E-01	13.560
0.730	0.310700E-01	13.160	0.740	0.294100E-01	12.840
0.750	0.278600E-01	12.650	0.760	0.264100E-01	12.360
0.770	0.250500E-01	12.070	0.780	0.237800E-01	11.830
0.790	0.225800E-01	11.610	0.800	0.214700E-01	101.170
0.900	0.133500E-01	81.560	1.000	0.873600E-02	6.820
1.100	0.595500E-02	5.560	1.200	0.419800E-02	4.640
1.300	0.304400E-02	3.850	1.400	0.226100E-02	2.870
1.500	0.171500E-02	3.230	1.600	0.132400E-02	2.140
1.700	0.103800E-02	1.750	1.800	0.825700E-03	1.440
1.900	0.664900E-03	1.200	2.000	0.541400E-03	0.950
2.500	0.221500E-03	0.150	3.000	0.106700E-03	0.030
3.500	0.576100E-04	0.530	4.000	0.000000	0.000

APPENDIX - B

AIR MASS AT THREE DIFFERENT HOUR ANGLES. $MH(H, THETA) = P(I)/P_0(SIN(THETA))$ FOR THE YEAR 1983 U.S. A.A. REGION. ($P_0 = 760MM$)

DAY	MH(TITA1)	MH(TITA2)	MH(TITA3)	DAY	MH(TITA1)	MH(TITA2)	MH(TITA3)
20	0.79930E 00	0.89170E 00	0.10928E 01	21	0.79693E 00	0.88893E 00	0.10891E 01
22	0.79461E 00	0.88622E 00	0.10855E 01	23	0.79508E 00	0.88663E 00	0.10857E 01
24	0.79560E 00	0.88709E 00	0.10859E 01	25	0.79477E 00	0.88605E 00	0.10843E 01
26	0.79125E 00	0.88200E 00	0.10791E 01	27	0.79188E 00	0.88259E 00	0.10795E 01
28	0.79119E 00	0.88170E 00	0.10781E 01	29	0.78916E 00	0.87933E 00	0.10749E 01
30	0.78718E 00	0.87701E 00	0.10717E 01	31	0.78798E 00	0.87773E 00	0.10724E 01
1	0.78745E 00	0.87708E 00	0.10712E 01	2	0.78832E 00	0.87794E 00	0.10720E 01
3	0.78787E 00	0.87733E 00	0.10709E 01	4	0.78610E 00	0.87524E 00	0.10661E 01
5	0.78436E 00	0.87319E 00	0.10653E 01	6	0.78676E 00	0.87575E 00	0.10681E 01
7	0.78510E 00	0.87379E 00	0.10655E 01	8	0.78621E 00	0.87491E 00	0.10665E 01
9	0.78462E 00	0.87305E 00	0.10640E 01	11	0.78429E 00	0.87246E 00	0.10627E 01
12	0.78691E 00	0.87526E 00	0.10658E 01	13	0.78547E 00	0.87356E 00	0.10635E 01
14	0.78679E 00	0.87492E 00	0.10648E 01	15	0.78542E 00	0.87329E 00	0.10626E 01
16	0.78680E 00	0.87473E 00	0.10641E 01	17	0.78686E 00	0.87469E 00	0.10637E 01
18	0.78695E 00	0.87468E 00	0.10635E 01	19	0.78706E 00	0.87471E 00	0.10632E 01
20	0.78434E 00	0.87168E 00	0.10596E 01	21	0.78721E 00	0.87477E 00	0.10631E 01
22	0.78603E 00	0.87336E 00	0.10611E 01	23	0.78350E 00	0.87046E 00	0.10573E 01
24	0.78646E 00	0.87364E 00	0.10609E 01	25	0.78808E 00	0.87535E 00	0.10627E 01
26	0.78700E 00	0.87405E 00	0.10609E 01	27	0.78867E 00	0.87581E 00	0.10628E 01
28	0.78900E 00	0.87608E 00	0.10629E 01	29	0.78799E 00	0.87487E 00	0.10612E 01
30	0.78836E 00	0.87519E 00	0.10613E 01	1	0.78875E 00	0.87553E 00	0.10615E 01
2	0.78779E 00	0.87438E 00	0.10599E 01	3	0.78685E 00	0.87325E 00	0.10583E 01
4	0.79141E 00	0.87822E 00	0.10641E 01	5	0.79050E 00	0.87713E 00	0.10625E 01
6	0.79098E 00	0.87757E 00	0.10628E 01	7	0.79147E 00	0.87803E 00	0.10632E 01
8	0.79334E 00	0.88003E 00	0.10654E 01	9	0.79248E 00	0.87900E 00	0.10639E 01
10	0.79438E 00	0.88102E 00	0.10662E 01	11	0.79217E 00	0.87849E 00	0.10629E 01
12	0.79271E 00	0.87901E 00	0.10633E 01	13	0.79188E 00	0.87801E 00	0.10619E 01
14	0.79380E 00	0.88008E 00	0.10642E 01	15	0.79298E 00	0.87909E 00	0.10629E 01
16	0.79492E 00	0.88117E 00	0.10652E 01	17	0.79548E 00	0.88172E 00	0.10657E 01
1	0.80340E 00	0.88964E 00	0.10731E 01	2	0.80524E 00	0.89164E 00	0.10754E 01
3	0.80288E 00	0.88899E 00	0.10721E 01	4	0.80469E 00	0.89095E 00	0.10743E 01
5	0.80368E 00	0.88981E 00	0.10729E 01	6	0.80546E 00	0.89174E 00	0.10751E 01
7	0.80582E 00	0.89210E 00	0.10755E 01	8	0.80896E 00	0.89555E 00	0.10795E 01

AIR MASS AT THICKLE DIFFERENT HOUR ANGLES. $MH(H, \theta) = P(I) / P_0 \sin(\theta)$ FOR THE YEAR 1963 OF A.A. REGION. ($P_0=760$ MM)

DAY	MH(TITA1)	MH(TITA2)	MH(TITA3)	DAY	MH(TITA1)	MH(TITA2)	MH(TITA3)
9	0.80788E 00	0.89433E 00	0.10780E 01	10	0.80673E 00	0.89308E 00	0.10754E 01
11	0.80846E 00	0.89492E 00	0.10766E 01	12	0.80732E 00	0.89363E 00	0.10770E 01
13	0.81036E 00	0.89697E 00	0.10809E 01	14	0.81057E 00	0.89719E 00	0.10816E 01
15	0.80935E 00	0.89583E 00	0.10795E 01	16	0.80952E 00	0.89599E 00	0.10795E 01
17	0.81247E 00	0.89924E 00	0.10835E 01	18	0.80276E 00	0.88849E 00	0.10750E 01
19	0.81127E 00	0.89791E 00	0.10819E 01	20	0.81134E 00	0.89798E 00	0.10812E 01
21	0.81139E 00	0.89802E 00	0.10820E 01	22	0.81000E 00	0.89649E 00	0.10801E 01
23	0.80999E 00	0.89648E 00	0.10801E 01	24	0.81137E 00	0.89700E 00	0.10820E 01
25	0.80991E 00	0.89639E 00	0.10800E 01	26	0.81123E 00	0.89735E 00	0.10813E 01
27	0.80972E 00	0.89620E 00	0.10798E 01	28	0.80819E 00	0.89453E 00	0.10773E 01
29	0.80804E 00	0.89436E 00	0.10777E 01	30	0.81067E 00	0.89729E 00	0.10812E 01
1	0.81046E 00	0.89708E 00	0.10810E 01	2	0.81024E 00	0.89632E 00	0.10807E 01
3	0.80719E 00	0.89350E 00	0.10768E 01	4	0.80973E 00	0.89633E 00	0.10803E 01
5	0.80804E 00	0.89448E 00	0.10781E 01	6	0.80772E 00	0.89417E 00	0.10770E 01
7	0.80739E 00	0.89383E 00	0.10775E 01	8	0.80704E 00	0.89347E 00	0.10772E 01
9	0.80807E 00	0.89464E 00	0.10787E 01	10	0.80628E 00	0.89270E 00	0.10754E 01
11	0.80588E 00	0.89229E 00	0.10760E 01	12	0.80685E 00	0.89341E 00	0.10770E 01
13	0.80502E 00	0.89142E 00	0.10752E 01	14	0.80317E 00	0.88941E 00	0.10722E 01
15	0.80131E 00	0.88740E 00	0.10705E 01	16	0.80083E 00	0.88691E 00	0.10701E 01
17	0.80312E 00	0.88950E 00	0.10733E 01	18	0.80122E 00	0.88745E 00	0.10710E 01
19	0.80209E 00	0.88847E 00	0.10723E 01	20	0.80157E 00	0.88794E 00	0.10713E 01
21	0.80242E 00	0.88894E 00	0.10732E 01	22	0.79910E 00	0.88532E 00	0.10691E 01
23	0.79994E 00	0.88631E 00	0.10703E 01	24	0.80076E 00	0.88730E 00	0.10717E 01
25	0.79745E 00	0.88367E 00	0.10674E 01	26	0.79827E 00	0.88465E 00	0.10684E 01
27	0.79771E 00	0.88409E 00	0.10683E 01	28	0.79576E 00	0.88209E 00	0.10659E 01
29	0.79658E 00	0.88298E 00	0.10673E 01	30	0.79740E 00	0.88395E 00	0.10667E 01
31	0.79684E 00	0.88340E 00	0.10682E 01	1	0.79491E 00	0.88133E 00	0.10659E 01
2	0.79298E 00	0.87927E 00	0.10636E 01	3	0.79381E 00	0.88027E 00	0.10655E 01
4	0.79465E 00	0.88128E 00	0.10664E 01	5	0.79327E 00	0.87975E 00	0.10645E 01
6	0.79412E 01	0.88077E 01	0.10660E 02	7	0.79360E 00	0.88027E 00	0.10644E 01
8	0.79309E 00	0.87979E 00	0.10652E 01	9	0.79254E 00	0.87932E 00	0.10642E 01
10	0.78937E 00	0.87583E 00	0.10608E 01	11	0.79027E 00	0.87691E 00	0.10624E 01
12	0.78981E 00	0.87649E 00	0.10621E 01	13	0.78938E 00	0.87609E 00	0.10613E 01
14	0.78895E 00	0.87572E 00	0.10616E 01	15	0.78855E 00	0.87535E 00	0.10614E 01
16	0.78680E 00	0.87351E 00	0.10594E 01	12	0.78644E 00	0.87320E 00	0.10590E 01

AIRMASS AT THREE DIFFERENT HOUR ANGLES. $MH(H, \theta) = P(I) / P_0(\sin(\theta))$ FOR THE YEAR 1953 IN A.A. REGION. ($P_0 = 760 \text{ mm}$).

DAY	MH(TITA1)	MH(TITA2)	MH(TITA3)	DAY	MH(TITA1)	MH(TITA2)	MH(TITA3)
18	0.78747E 00	0.87443E 00	0.10610E 01	23	0.78340E 00	0.87039E 00	0.10573E 01
24	0.78730E 00	0.87482E 00	0.10630E 01	25	0.78441E 00	0.87171E 00	0.10575E 01
26	0.78428E 00	0.87167E 00	0.10597E 01	27	0.78418E 00	0.87155E 00	0.10570E 01
28	0.78410E 00	0.87168E 00	0.10602E 01	29	0.78542E 00	0.87325E 00	0.10624E 01
30	0.78542E 00	0.87334E 00	0.10628E 01	31	0.78680E 00	0.87497E 00	0.10651E 01
1	0.78414E 00	0.87213E 00	0.10619E 01	2	0.78560E 00	0.87336E 00	0.10649E 01
4	0.78725E 00	0.87591E 00	0.10673E 01	5	0.78609E 00	0.87473E 00	0.10682E 01
6	0.78360E 00	0.87207E 00	0.10632E 01	7	0.78524E 00	0.87401E 00	0.10659E 01
9	0.78419E 00	0.87295E 00	0.10649E 01	9	0.78591E 00	0.87497E 00	0.10678E 01
10	0.78766E 00	0.87704E 00	0.10704E 01	11	0.78946E 00	0.87915E 00	0.10739E 01
12	0.78857E 00	0.87827E 00	0.10725E 01	13	0.78908E 00	0.87895E 00	0.10737E 01
14	0.79100E 00	0.88120E 00	0.10767E 01	15	0.79022E 00	0.88045E 00	0.10751E 01
16	0.78948E 00	0.87975E 00	0.10756E 01	17	0.78879E 00	0.87909E 00	0.10751E 01
18	0.79087E 00	0.88153E 00	0.10783E 01	19	0.79300E 00	0.88402E 00	0.10817E 01
20	0.79380E 00	0.88503E 00	0.10833E 01	21	0.79189E 00	0.88301E 00	0.10811E 01
22	0.79553E 00	0.88719E 00	0.10865E 01	23	0.79507E 00	0.88680E 00	0.10919E 01
24	0.79742E 00	0.88954E 00	0.10900E 01	25	0.79704E 00	0.88924E 00	0.10930E 01
26	0.79809E 00	0.89053E 00	0.10919E 01	27	0.80057E 00	0.89341E 00	0.10953E 01
28	0.80031E 00	0.89325E 00	0.10959E 01	29	0.80288E 00	0.89623E 00	0.10999E 01
30	0.80688E 00	0.90082E 00	0.11058E 01	1	0.80674E 00	0.90079E 00	0.11061E 01
2	0.80944E 00	0.90392E 00	0.11103E 01	3	0.80798E 00	0.90241E 00	0.11088E 01
4	0.80795E 00	0.90250E 00	0.11092E 01	5	0.81217E 00	0.90734E 00	0.11122E 01
6	0.81362E 00	0.90909E 00	0.11180E 01	7	0.81229E 00	0.90773E 00	0.11180E 01
8	0.81523E 00	0.91113E 00	0.11211E 01	9	0.81821E 00	0.91459E 00	0.11257E 01
10	0.81697E 00	0.91333E 00	0.11245E 01	11	0.81718E 00	0.91369E 00	0.11253E 01
12	0.81884E 00	0.91567E 00	0.11281E 01	13	0.82197E 00	0.91928E 00	0.11328E 01
14	0.82226E 00	0.91974E 00	0.11338E 01	15	0.82402E 00	0.92183E 00	0.11367E 01
16	0.82581E 00	0.92396E 00	0.11396E 01	17	0.82907E 00	0.92773E 00	0.11449E 01
18	0.83092E 00	0.92992E 00	0.11476E 01	19	0.83424E 00	0.93376E 00	0.11527E 01
20	0.83470E 00	0.93440E 00	0.11538E 01	21	0.83808E 00	0.93831E 00	0.11570E 01
22	0.84003E 00	0.94062E 00	0.11622E 01	23	0.84201E 00	0.94295E 00	0.11634E 01
24	0.84411E 00	0.94531E 00	0.11687E 01	25	0.84749E 00	0.94734E 00	0.11740E 01
26	0.85100E 00	0.95339E 00	0.11793E 01	27	0.85159E 00	0.95417E 00	0.11800E 01
28	0.85366E 00	0.95661E 00	0.11840E 01	29	0.85576E 00	0.95907E 00	0.11830E 01

30	0.85734E 00	0.96154E 00	0.11907E 01	31	0.85995E 00	0.96402E 00	0.11941E 01
1	0.86206E 00	0.96651E 00	0.11975E 01	2	0.86568E 00	0.97069E 00	0.12030E 01
3	0.86531E 00	0.97151E 00	0.12044E 01	4	0.86844E 00	0.97402E 00	0.12073E 01
5	0.87052E 00	0.97653E 00	0.12112E 01	6	0.87271E 00	0.97903E 00	0.12147E 01
7	0.87484E 00	0.98154E 00	0.12181E 01	8	0.87545E 00	0.98233E 00	0.12174E 01
9	0.87908E 00	0.98652E 00	0.12249E 01	10	0.87967E 00	0.98729E 00	0.12201E 01
11	0.88330E 00	0.99147E 00	0.12316E 01	12	0.88385E 00	0.99220E 00	0.12325E 01
13	0.88746E 00	0.99635E 00	0.12383E 01	14	0.88952E 00	0.99877E 00	0.12415E 01
15	0.89157E 00	0.10012E 01	0.12449E 01	16	0.89358E 00	0.10035E 01	0.12491E 01
17	0.89557E 00	0.10059E 01	0.12513E 01	18	0.89755E 00	0.10082E 01	0.12544E 01
19	0.90105E 00	0.10122E 01	0.12597E 01	20	0.90141E 00	0.10127E 01	0.12608E 01
21	0.90017E 00	0.10114E 01	0.12592E 01	22	0.90201E 00	0.10135E 01	0.12621E 01
23	0.90539E 00	0.10175E 01	0.12672E 01	24	0.90874E 00	0.10213E 01	0.12723E 01
25	0.90734E 00	0.10198E 01	0.12706E 01	26	0.90903E 00	0.10218E 01	0.12733E 01
27	0.91063E 00	0.10237E 01	0.12760E 01	28	0.91387E 00	0.10274E 01	0.12808E 01
29	0.91702E 00	0.10310E 01	0.12855E 01	30	0.91695E 00	0.10310E 01	0.12872E 01
1	0.92000E 00	0.10345E 01	0.12902E 01	2	0.92142E 00	0.10362E 01	0.12925E 01
3	0.92278E 00	0.10378E 01	0.12946E 01	4	0.92568E 00	0.10411E 01	0.12995E 01
5	0.92372E 00	0.10389E 01	0.12964E 01	6	0.92651E 00	0.10421E 01	0.13008E 01
7	0.92603E 00	0.10417E 01	0.13001E 01	8	0.92870E 00	0.10447E 01	0.13041E 01
9	0.92810E 00	0.10441E 01	0.13034E 01	10	0.92903E 00	0.10452E 01	0.13047E 01
12	0.93151E 00	0.10480E 01	0.13085E 01	13	0.93232E 00	0.10489E 01	0.13076E 01
14	0.93305E 00	0.10498E 01	0.13110E 01	15	0.93372E 00	0.10506E 01	0.13121E 01
16	0.93432E 00	0.10513E 01	0.13130E 01	17	0.93322E 00	0.10501E 01	0.13115E 01
18	0.93530E 00	0.10524E 01	0.13146E 01	19	0.93730E 00	0.10547E 01	0.13175E 01
20	0.93599E 00	0.10532E 01	0.13157E 01	21	0.93734E 00	0.10553E 01	0.13183E 01
22	0.93634E 00	0.10536E 01	0.13162E 01	23	0.93487E 00	0.10520E 01	0.13142E 01
24	0.93488E 00	0.10520E 01	0.13142E 01	26	0.93632E 00	0.10535E 01	0.13151E 01
27	0.93612E 00	0.10534E 01	0.13159E 01	28	0.93584E 00	0.10531E 01	0.13154E 01
29	0.93550E 00	0.10527E 01	0.13149E 01	30	0.93346E 00	0.10504E 01	0.13119E 01
31	0.93269E 00	0.10494E 01	0.13104E 01		0.00000	0.00000	0.00000

APPENDIX - C

CALCULATED VALUES OF DIRECT SOLAR RAD. DU TO AYLEIGH SCATTERING
 CHANNELS INITIALLY FORBIDDEN FACTORS FOR THE YEAR 1983 OF A.A. REGION

DAY	INITIALLY	FORBIDDEN	FACTORS	DAY	INITIALLY
1	0.9317124E-03	0.1775227E-01	3	0.9313916E-03	
5	0.9320720E-03	0.1958809E-01	6	0.9319731E-03	
8	0.9323645E-03	0.1767316E-01	9	0.9325623E-03	
11	0.9329453E-03	0.1860602E-01	12	0.9332134E-03	
14	0.9337903E-03	0.2182322E-01	19	0.9355142E-03	
21	0.9359907E-03	0.2067322E-01	23	0.9365967E-03	
25	0.9377734E-03	0.2328668E-01	26	0.9381736E-03	
28	0.9386431E-03	0.2082022E-01	29	0.9387146E-03	
31	0.9393667E-03	0.2700492E-01	1	0.9410659E-03	
3	0.9413953E-03	0.3247268E-01	4	0.9414780E-03	
6	0.9434365E-03	0.1797636E-01	7	0.9438552E-03	
9	0.9448572E-03	0.2038118E-01	10	0.9467329E-03	
12	0.9467231E-03	0.1955879E-01	13	0.9471560E-03	
15	0.9495032E-03	0.2264073E-01	16	0.9496352E-04	
18	0.9502417E-03	0.2087000E-01	20	0.9525105E-03	
22	0.9526606E-03	0.1959522E-01	23	0.9528955E-03	
25	0.9555667E-03	0.1778511E-01	27	0.9560166E-03	
3	0.9587661E-03	0.1911184E-01	4	0.9591335E-03	
6	0.9619438E-03	0.1930032E-01	7	0.9621384E-03	
9	0.9623557E-03	0.2002860E-01	10	0.9646899E-03	
12	0.9652021E-03	0.2118439E-01	13	0.9358564E-03	
15	0.9685981E-03	0.2225161E-01	16	0.9685947E-03	
18	0.9687327E-03	0.2419759E-01	19	0.9690339E-03	
21	0.9720471E-03	0.1848529E-01	22	0.9723337E-03	
24	0.9722563E-03	0.1759470E-01	25	0.9753101E-03	
27	0.9756819E-03	0.2138545E-01	28	0.9757805E-03	
30	0.9762815E-03	0.1921893E-01	31	0.9762065E-03	
2	0.9797341E-03	0.1893628E-01	3	0.9798040E-03	
5	0.9828989E-03	0.1969308E-01	6	0.9826360E-03	
8	0.9827349E-03	0.2639453E-01	9	0.9829368E-03	
12	0.9853708E-03	0.1883811E-01	13	0.9855559E-03	
15	0.9882446E-03	0.1931442E-01	16	0.9880986E-03	
18	0.9881143E-03	0.1861037E-01	19	0.9881160E-03	
21	0.9909614E-03	0.2156385E-01	22	0.9911157E-03	
24	0.9910957E-03	0.1813504E-01	25	0.9937668E-03	
27	0.9937275E-03	0.2059383E-01	28	0.9937036E-03	
30	0.9938083E-03	0.1807763E-01	1	0.9967197E-03	
3	0.9969717E-03	0.2105775E-01	4	0.9964485E-03	
6	0.9986846E-03	0.1813368E-01	7	0.9986404E-03	
9	0.9985459E-03	0.2047268E-01	10	0.1000493E-04	
12	0.1000715E-04	0.1866539E-01	13	0.1000825E-04	
15	0.1002778E-04	0.2000292E-01	16	0.1002559E-04	
1	0.1007976E-04	0.2064202E-01	2	0.1007765E-04	
4	0.1007843E-04	0.1825695E-01	5	0.1009143E-04	
7	0.1008900E-04	0.1870040E-01	8	0.1008533E-04	
10	0.1009977E-04	0.2096642E-01	11	0.1009781E-04	
13	0.1009564E-04	0.2090029E-01	14	0.1009541E-04	
16	0.1010552E-04	0.2070203E-01	17	0.1010204E-04	
19	0.1010348E-04	0.2108499E-01	20	0.1011321E-04	
22	0.1011480E-04	0.2050535E-01	23	0.1011481E-04	
25	0.1011980E-04	0.1950585E-01	26	0.1011822E-04	

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4	0.1417541	01	3	0.1011011	04
5	0.2097213	01	5	0.1011011	04
6	0.1871550	01	6	0.1011011	04
9	0.2013201E	01	9	0.1012644E	04
11	0.1951879E	01	12	0.1012771E	04
14	0.2179893E	01	15	0.1012821E	04
17	0.2144273E	01	18	0.1012851E	04
20	0.2317795E	01	21	0.1011993E	04
23	0.2419267E	01	24	0.1012121E	04
26	0.2259707E	01	27	0.1011672E	04
29	0.2147434E	01	30	0.1011872E	04
1	0.2453382E	01	2	0.1010599E	04
4	0.3490979E	01	5	0.1009302E	04
7	0.2244259E	01	8	0.1009345E	04
10	0.1959259E	01	11	0.1008072E	04
13	0.3157728E	01	14	0.1008187E	04
16	0.2723845E	01	17	0.1008971E	04
23	0.2017715E	01	24	0.1004799E	04
26	0.2270573E	01	27	0.1002864E	04
29	0.2163318E	01	30	0.1002867E	04
1	0.2534623E	01	2	0.9994543E	03
5	0.2011445E	01	6	0.9975605E	03
8	0.1959220E	01	9	0.9972344E	03
11	0.2054313E	01	12	0.9946091E	03
14	0.2974091E	01	15	0.9917109E	03
17	0.2033231E	01	18	0.9915788E	03
20	0.3121680E	01	21	0.9887559E	03
23	0.1870265E	01	24	0.9880500E	03
26	0.1935922E	01	27	0.9845074E	03
29	0.2263136E	01	30	0.9837935E	03
2	0.2026622E	01	3	0.9815743E	03
5	0.1963918E	01	6	0.9777109E	03
8	0.2247880E	01	9	0.9771221E	03
11	0.2106748E	01	12	0.9691743E	03
14	0.1943817E	01	15	0.9706719E	03
17	0.2136915E	01	18	0.9598245E	03
20	0.1893017E	01	21	0.9582202E	03
23	0.2266781E	01	24	0.9554895E	03
26	0.1852437E	01	27	0.9516543E	03
29	0.1827884E	01	30	0.9508943E	03
1	0.1755009E	01	2	0.9515538E	03
4	0.1926484E	01	5	0.9538230E	03
7	0.1910732E	01	8	0.9552292E	03
10	0.1827705E	01	11	0.9501715E	03
13	0.1858521E	01	14	0.9494897E	03
16	0.1850235E	01	17	0.9485907E	03
19	0.1774362E	01	20	0.9487008E	03
22	0.1734537E	01	23	0.9433130E	03
25	0.1774655E	01	26	0.9435007E	03
28	0.1895700E	01	29	0.9411559E	03
1	0.1792945E	01	2	0.9370959E	03

3	0.1774134E 01	7	0.1774134E 01	0.1774134E 01
6	0.1774134E 01	7	0.1774134E 01	0.1774134E 01
9	0.1774134E 01	7	0.1774134E 01	0.1774134E 01
12	0.1724643E 01	13	0.1724643E 01	0.1724643E 01
19	0.1426690E 01	20	0.1426690E 01	0.1426690E 01
23	0.1394467E 01	24	0.1394467E 01	0.1394467E 01
26	0.2144469E 01	27	0.2144469E 01	0.2144469E 01
29	0.1951315E 01	30	0.1951315E 01	0.1951315E 01
1	0.2311913E 01	2	0.2311913E 01	0.2311913E 01
4	0.1239767E 01	5	0.1239767E 01	0.1239767E 01
7	0.1644977E 01	8	0.1644977E 01	0.1644977E 01
10	0.1846744E 01	11	0.1846744E 01	0.1846744E 01
13	0.2042333E 01	14	0.2042333E 01	0.2042333E 01
16	0.2837294E 01	17	0.2837294E 01	0.2837294E 01
20	0.1969634E 01	21	0.1969634E 01	0.1969634E 01
23	0.1777023E 01	24	0.1777023E 01	0.1777023E 01
27	0.2424215E 01	28	0.2424215E 01	0.2424215E 01
4	0.1839252E 01	5	0.1839252E 01	0.1839252E 01
7	0.1914223E 01	8	0.1914223E 01	0.1914223E 01
10	0.1836378E 01	11	0.1836378E 01	0.1836378E 01
13	0.2296628E 01	14	0.2296628E 01	0.2296628E 01
16	0.2562330E 01	17	0.2562330E 01	0.2562330E 01
19	0.2334957E 01	20	0.2334957E 01	0.2334957E 01
22	0.2023642E 01	23	0.2023642E 01	0.2023642E 01
25	0.2151172E 01	26	0.2151172E 01	0.2151172E 01
28	0.2143397E 01	29	0.2143397E 01	0.2143397E 01
31	0.2033309E 01	32	0.2033309E 01	0.2033309E 01
3	0.2153517E 01	4	0.2153517E 01	0.2153517E 01
6	0.2266506E 01	7	0.2266506E 01	0.2266506E 01
9	0.2929363E 01	10	0.2929363E 01	0.2929363E 01
13	0.1395343E 01	14	0.1395343E 01	0.1395343E 01
16	0.2064070E 01	17	0.2064070E 01	0.2064070E 01
19	0.1525732E 01	20	0.1525732E 01	0.1525732E 01
22	0.3026238E 01	23	0.3026238E 01	0.3026238E 01
25	0.1897392E 01	26	0.1897392E 01	0.1897392E 01
28	0.1913345E 01	29	0.1913345E 01	0.1913345E 01
1	0.1339224E 01	2	0.1339224E 01	0.1339224E 01
4	0.2194009E 01	5	0.2194009E 01	0.2194009E 01
7	0.1322009E 01	8	0.1322009E 01	0.1322009E 01
10	0.1827317E 01	11	0.1827317E 01	0.1827317E 01
13	0.1392694E 01	14	0.1392694E 01	0.1392694E 01
16	0.1959179E 01	17	0.1959179E 01	0.1959179E 01
2	0.2265334E 01	3	0.2265334E 01	0.2265334E 01
5	0.1915701E 01	6	0.1915701E 01	0.1915701E 01
8	0.1540795E 01	9	0.1540795E 01	0.1540795E 01
11	0.1160227E 01	12	0.1160227E 01	0.1160227E 01
14	0.1737030E 01	15	0.1737030E 01	0.1737030E 01
17	0.2115351E 01	18	0.2115351E 01	0.2115351E 01
20	0.2045233E 01	21	0.2045233E 01	0.2045233E 01
23	0.2044158E 01	24	0.2044158E 01	0.2044158E 01
26	0.2046222E 01	27	0.2046222E 01	0.2046222E 01

CALCULATED VALUES OF DIRECT SOLAR RAD. DUE TO RAY SCATTER
 FROM THE SUN, BY THE EARTH'S SURFACE YEAR 1970
 DAY IN (TIME) DAY IN (TIME)

28	0.1012180E 04	0.2007111E 01	29	0.1012196E 04
1	0.1012392E 04	0.1962868E 01	2	0.1012410E 04
4	0.1012470E 04	0.2044507E 01	5	0.1012667E 04
7	0.1012734E 04	0.2206241E 01	8	0.1032365E 04
10	0.1012850E 04	0.1951943E 01	11	0.1012893E 04
13	0.1012982E 04	0.1991800E 01	14	0.1013196E 04
16	0.1012871E 04	0.2148352E 01	17	0.1012590E 04
19	0.1012696E 04	0.2073138E 01	20	0.1012063E 04
22	0.1012339E 04	0.3269361E 01	23	0.1012230E 04
25	0.1011724E 04	0.2170094E 01	26	0.1011615E 04
28	0.1011894E 04	0.2099422E 01	29	0.1011735E 04
31	0.1011732E 04	0.2781960E 01	1	0.1010361E 04
3	0.1010488E 04	0.2416448E 01	4	0.1010375E 04
6	0.1001660E 04	0.1896223E 01	7	0.1009297E 04
9	0.1009391E 04	0.1960402E 01	10	0.1008193E 04
12	0.1008113E 04	0.2552288E 01	13	0.1002262E 04
15	0.1006748E 04	0.3739030E 01	16	0.1006942E 04
18	0.1006834E 04	0.2794729E 01	23	0.1005280E 04
25	0.1002869E 04	0.2797448E 01	26	0.1002868E 04
28	0.1002856E 04	0.2272294E 01	29	0.1002683E 04
31	0.1002485E 04	0.2992448E 01	1	0.9996440E 03
4	0.9992231E 03	0.1985176E 01	5	0.9972832E 03
7	0.9973486E 03	0.1940306E 01	8	0.9974556E 03
10	0.9947517E 03	0.2089625E 01	11	0.9945205E 03
13	0.9945310E 03	0.3082137E 01	14	0.9942854E 03
16	0.9917798E 03	0.2108763E 01	17	0.9918435E 03
19	0.9913091E 03	0.3172833E 01	20	0.9885496E 03
22	0.9883096E 03	0.2378400E 01	23	0.9883447E 03
25	0.9850371E 03	0.1963362E 01	26	0.9848960E 03
28	0.9845981E 03	0.3327519E 01	29	0.9842800E 03
1	0.9807559E 03	0.2004913E 01	2	0.9804233E 03
4	0.9805583E 03	0.2003734E 01	5	0.9778982E 03
7	0.9778452E 03	0.2063915E 01	8	0.9774856E 03
10	0.9743142E 03	0.1841813E 01	11	0.9742708E 03
13	0.9736829E 03	0.2140621E 01	14	0.9736274E 03
16	0.9704480E 03	0.1810369E 01	17	0.9700554E 03
19	0.9694255E 03	0.2554647E 01	20	0.9666238E 03
22	0.9659785E 03	0.992888E 01	23	0.9657354E 03
25	0.9621548E 03	0.1786043E 01	26	0.9617395E 03
28	0.9614123E 03	0.1797642E 01	29	0.9611444E 03
31	0.9606389E 03	0.177961E 01	1	0.9569617E 03
3	0.9564702E 03	0.1872636E 01	4	0.9562146E 03
6	0.9535684E 03	0.2053928E 01	7	0.9533137E 03
9	0.9528033E 03	0.2795946E 01	1	0.953935E 03
12	0.9500974E 03	0.139213E 01	13	0.9496524E 03
15	0.9470671E 03	0.182561E 01	16	0.9469313E 03
18	0.9463655E 03	0.1886249E 01	19	0.9469643E 03
21	0.943980E 03	0.1762004E 01	22	0.9436921E 03
24	0.9429299E 03	0.180341E 01	25	0.9412361E 03
27	0.9408474E 03	0.184508E 01	28	0.944886E 03
30	0.9401331E 03	0.1849857E 01	31	0.937615E 03

APPENDIX - D

```

DOUBLE PRECISION SS(400,10),APD4(400),APD41(400),DS(3),CPD41(400),
F1(10),FM(10),YMIN(10),VMIN,CPD42(400),CPD43(400),VMN1,VMN2,VMN3,
(10),C(400),D(400),R(400,8),X,Z(400),SS5(400),F(10),T,G(400),FY(10),
),A,BB,CC,DD,E(400),Y(400)
REAL W(10),U(400,8),V(400),V,H
INTEGER NN,JT,IT,MTD
COMMON L(10),K(10),A(10,10),BB(10,10),F(10),CC(10),DD(10)
DATA H/9.033/
FORMAT(F5.0,F10.1,F6.2,F5.1,F8.2)
M10=0
MTD=313
X=0.00
DO 1115 I=1,MTD
D(I)=0.00
F(I)=0.00
C(I)=0.00
Z(I)=0.00
G(I)=0.00
115CONTINUE
DO 1116 I=1,8
DO 1117 J=1,MTD
R(J,I)=0.00
117CONTINUE
116CONTINUE
JT=10
IT=JT-1
READ(2,2020)(W(I),I=1,9)
DO 919 I=1,IT
F(I)=(DBLE(W(I)))
919CONTINUE
2020FORMAT(1X,9F7.4)
DO 3 J=1,MTD
READ(2,1)Q(J),(U(J,I),I=1,4)
3CONTINUE
DO 4 I=1,MTD
R(I,2)=(DBLE(U(I,2))/(100.0))
R(I,1)=(DBLE(U(I,1)))
R(I,3)=(DBLE(U(I,3)))
R(I,4)=(DBLE(U(I,4)))
4CONTINUE
DO 5 I=1,MTD
E(I)=DBLE(U(I,4))
C(I)=DBLE(Q(I))
Z(I)=R(I,1)/12.00
D(I)=0.00
5CONTINUE
CALL KROSE(MTD,D,E,Y)

```

```

      I=1
DO 9 LLL=1,5
DO 1112 I=1,MTD
  APOW(I)=0.00
  APOW1(I)=0.00
  QPOW1(I)=0.00
  QPOW2(I)=0.00
  QPOW3(I)=0.00
  SSS(I)=0.00
CONTINUE
DO 1113 I=1,IT
DO 1114 J=1,MTD
  SS(J,I)=0.00
CONTINUE
CONTINUE
DO 8 N=1,MTD
  APOW(N)=(((F(2)*(X**F(3)))*((F(4)*(Z(N)**F(5))+F(6)*(R(N,2)**F(7))+
  F(8)*(R(N,3)**F(9))))))
CONTINUE
DO 9 N=1,MTD
  APOW1(N)=((X**F(3))*((F(4)*(Z(N)**F(5))+F(6)*(R(N,2)**F(7))+F(8)*(
  R(N,3)**F(9))))))
CONTINUE
DO 10 N=1,MTD
  SSS(N)=((F(1)*D(N))*(DEXP(APOW(N))))
CONTINUE
DO 11 N=1,MTD
  SS(N,1)=(D(N)*(DEXP(APOW(N))))
CONTINUE
DO 12 N=1,MTD
  SS(N,2)=(((F(1)*D(N))*(APOW1(N)))*(DEXP(APOW(N))))
CONTINUE
DO 13 N=1,MTD
  SS(N,3)=((((F(1)*D(N))*((F(2)*(X**F(3)))*(Z(N)**F(5))))*(DEXP(APOW
  (N))))))
CONTINUE
DO 14 N=1,MTD
  SS(N,4)=((F(1)*D(N))*(((F(2)*(X**F(3)))*(R(N,2)**F(7))))*(DEXP(APO
  W(N))))))
CONTINUE
DO 15 N=1,MTD
  SS(N,5)=((F(1)*D(N))*((F(2)*(X**F(3)))*(R(N,3)**F(9))))*(DEXP(APOW(
  N))))))
CONTINUE
DO 16 N=1,MTD
  SS(N,6)=(((F(1)*D(N))*(DEXP(APOW(N))))*((X**F(3))*(DLOG(X))))*(F(
  2)*(((F(4)*(Z(N)**F(5)))+(F(6)*(R(N,2)**F(7)))+(F(8)*(R(N,3)**F(9)
  ))))))))
CONTINUE
DO 17 N=1,MTD
  SS(N,7)=(((F(1)*D(N))*(DEXP(APOW(N))))*(F(2)*(X**F(3))*F(4))*((DLOG
  (Z(N)))*(Z(N)**F(5))))))
CONTINUE
DO 18 N=1,MTD

```

```

1111 DO 1112 I=1,MT0
      APOW(I)=0.00
      APOW1(I)=0.00
      APOW2(I)=0.00
      APOW3(I)=0.00
      SSS(I)=0.00
1112 CONTINUE
      DO 1113 I=1,IT
      DO 1114 J=1,MT0
      SSS(J,I)=0.00
1114 CONTINUE
1115 CONTINUE
      DO 8 N=1,MT0
      APOW(N)=(((F(2)*(X**F(3)))*((F(4)*(Z(N)**F(5))+F(6)*(R(N,2)**F(7))+
4 F(8)*(R(N,3)**F(9))))))
      8 CONTINUE
      DO 9 N=1,MT0
      APOW1(N)=((X**F(3))*((F(4)*(Z(N)**F(5))+F(6)*(R(N,2)**F(7))+F(8)*(
5 R(N,3)**F(9))))))
      9 CONTINUE
      DO 10 N=1,MT0
      SSS(N)=((F(1)*D(N))*(DEXP(APOW(N))))
10 CONTINUE
      DO 11 N=1,MT0
      SSS(N,1)=(D(N)*(DEXP(APOW(N))))
11 CONTINUE
      DO 12 N=1,MT0
      SSS(N,2)=(((F(1)*D(N))*(APOW1(N))*(DEXP(APOW(N))))
12 CONTINUE
      DO 13 N=1,MT0
      SSS(N,3)=(((F(1)*D(N))*((F(2)*(X**F(3))*(Z(N)**F(5))))*(DEXP(APOW
6 (N))))))
13 CONTINUE
      DO 14 N=1,MT0
      SSS(N,4)=((F(1)*D(N))*((F(2)*(X**F(3))*(R(N,2)**F(7))))*(DEXP(APO
7 W(N))))
      CONTINUE
      DO 15 N=1,MT0
      SSS(N,5)=((F(1)*D(N))*((F(2)*(X**F(3))*(R(N,3)**F(9))))*(DEXP(APOW1
8 (N))))
15 CONTINUE
      DO 16 N=1,MT0
      SSS(N,6)=(((F(1)*D(N))*(DEXP(APOW(N)))*(((X**F(3))*(DLOG(X))))*(F(
9 2)*(((F(4)*(Z(N)**F(5)))+(F(6)*(R(N,2)**F(7)))+(F(8)*(R(N,3)**F(9)
))))))
16 CONTINUE
      DO 17 N=1,MT0
      SSS(N,7)=(((F(1)*D(N))*(DEXP(APOW(N)))*((F(2)*(X**F(3))*F(4))*((DLOG
10 (Z(N)))*(Z(N)**F(5))))))
17 CONTINUE
      DO 18 N=1,MT0

```

```

SS(N,9)=(((F(1)*D(N))*(DEXP(APOA(N))))*(F(2)*(X**F(3)))*(F(4)*(Z(N)**F(5)))+(F(6)*(R(N,2)**F(7))+F(8)*(R(N,3)**F(9))))))
18CONTINUE
DO 19 N=1,MT0
SS(N,9)=(((F(1)*D(N))*(DEXP(APOA(N))))*(F(2)*(X**F(3)))*(F(4)*(Z(N)**F(5)))+(F(6)*(R(N,2)**F(7))+F(8)*(R(N,3)**F(9))))))
19CONTINUE
DO 222 I=1,IT
B(I)=0.00
T(I)=0.00
CC(I)=0.00
DD(I)=0.00
DO 2221 J=1,JT
A(I,J)=0.00
BB(I,J)=0.00
2221CONTINUE
2222CONTINUE
DO 29 I=1,IT
DO 30 N=1,MT0
B(I)=B(I)+(C(N)-SSS(N))*SS(N,I)
30CONTINUE
31CONTINUE
DO 36 J=1,IT
DO 37 KK=1,IT
DO 35 K=1,MT0
A(J,KK)=A(J,KK)+(SS(N,J)*SS(N,KK))
BB(J,KK)=BB(J,KK)+(SS(N,J)*SS(N,KK))
35CONTINUE
37CONTINUE
36CONTINUE
DO 38 I=1,IT
A(I,10)=B(I)
BB(I,10)=B(I)
38CONTINUE
CALL GAUSS(JT,IT,MT0)
35FORMAT(1X,7T4,'Y (',I2,') = ',D12.5)
DO 41 I=1,IT
F1(I)=F(I)+T(I)*0.500
FM(I)=F(I)+T(I)
FY(I)=F(I)
41CONTINUE
WRITE(3,97)
97FORMAT(1X,7T5,15HY(I)+(0.5*X(I)),4X,9HY(I)+X(I),8X,4HY(I)/)
96FORMAT(1X,7I3,2X,D12.5,2X,D12.5,2X,D12.5)
DO 42 I=1,IT
WRITE(3,96)I,F1(I),FM(I),FY(I)
42CONTINUE
DS(1)=0.00
DS(2)=0.00
DS(3)=0.00
DO 44 N=1,MT0
QPQWI(N)=(((F(1)*D(N))*(DEXP((F(2)*(X**F(3)))*(F(4)*(Z(N)**F(5)))+(F(6)*(R(N,2)**F(7))+F(8)*(R(N,3)**F(9))))))
44CONTINUE

```

```

      DO 46 N=1, MTO
      QS(1)=QS(1)-(C(N)-QP0W1(N))**(2)
46 CONTINUE
      DO 1118 I=1, IT
      F(I)=0.00
1118 CONTINUE
      DO 47 I=1, IT
      F(I)=F1(I)
47 CONTINUE
      DO 48 N=1, MTO
      QP0W2(N)=((F(1)*D(N))*(DEXP((F(2)*(X**F(3))))*((F(4)*(Z(N)**F(5))+F
(5)*(R(N,2)**F(7))+F(8)*(R(N,3)**F(9))))))
48 CONTINUE
      DO 50 N=1, MTO
      QS(2)=QS(2)+(C(N)-QP0W2(N))**(2)
50 CONTINUE
      DO 1119 I=1, IT
      F(I)=0.00
1119 CONTINUE
      DO 51 I=1, IT
      F(I)=FM(I)
51 CONTINUE
      DO 52 N=1, MTO
      QP0W3(N)=((F(1)*D(N))*(DEXP((F(2)*(X**F(3))))*((F(4)*(Z(N)**F(5))+F
(5)*(R(N,2)**F(7))+F(8)*(R(N,3)**F(9))))))
52 CONTINUE
      DO 53 N=1, MTO
      QS(3)=QS(3)+(C(N)-QP0W3(N))**(2)
53 CONTINUE
      VMIN=0.00
      VMIN=((QS(1)-QS(3))/((QS(3)+QS(1)-QS(2)*2.00)*4.00)+0.500)
      WRITE(3,978)
978 FORMAT(1X,/T8,4H0(0),T21,6H0(1/2),T36,4H0(1)/)
      WRITE(3,210)(QS(I),I=1,3)
210 FORMAT(1X,3D12.5)
      WRITE(3,235)VMIN
235 FORMAT(1X,/T4,'VMINIMUM = ',D12.5)
      DO 1120 I=1, IT
      F(I)=0.00
1120 CONTINUE
      DO 54 I=1, IT
      F(I)=FY(I)+I(I)*VMIN
54 CONTINUE
      DO 55 I=1, IT
      WRITE(3,95)I,F(I)
55 CONTINUE
99 CONTINUE
100 STOP
      END

```

```

SUBROUTINE KABS (MTD, K, Z, ZI)
DOUBLE PRECISION AK(400), SI(400), A2, ZI(400), ZI1(400)
INTEGER M, NTO
DATA 0/0
DO 18 I=1, NTO
  SIG(I)=0.00
  ZZ(I)=(2.00/15.00)*EE(I)
18 CONTINUE
  A2=0.00
  M=0
  DO 4 I=1, NTO
    M=M+1
    IF((I.LE.27))A2=1.280
    IF((I.GE.28).AND.(I.LE.53))A2=1.3800
    IF((I.GE.54).AND.(I.LE.82))A2=1.5400
    IF((I.GE.83).AND.(I.LE.109))A2=1.7700
    IF((I.GE.110).AND.(I.LE.135))A2=2.3000
    IF((I.GE.136).AND.(I.LE.166))A2=2.4800
    IF((I.GE.167).AND.(I.LE.195))A2=2.4100
    IF((I.GE.196).AND.(I.LE.225))A2=2.3600
    IF((I.GE.226).AND.(I.LE.254))A2=1.7300
    IF((I.GE.255).AND.(I.LE.283))A2=1.3800
    IF((I.GE.284).AND.(I.LE.313))A2=1.1700
    SIG(M)=A2
  4 CONTINUE
  M=0
  LL=0
  KK=0
  DO 110 I=1, MTO
    KK=KK+1
    M=M+1
    LL=LL+1
    AK(KK)=((0.10500)*ZZ(LL)+((0.987500)*SIG(M)))*(4184.00)
110 CONTINUE
  K=TO-N
  END

```

```

SUBROUTINE GAUSS(JT,IT,IT)
CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC

```

THE PURPOSE OF THIS PROGRAM IS TO SOLVE UP TO 10 SIMULTANEOUS EQUATIONS BY GAUSS-JORDAN METHOD.

N = THE NUMBER OF DECIMAL PLACES DESIRED.

JT = THE NUMBER OF COLUMNS, NO MORE THAN 10 ENTRIES PER CARD.

```

CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC

```

```

DOUBLE PRECISION RR,T,A,PS,CC,DD

```

```

REAL ZZ

```

```

INTEGER N,JI,0,JI,IT

```

```

COMMON L(10),K(10),A(10,10),B(10,10),T(10),CC(10),DD(10)

```

```

DATA ZZ,0,RR/0.0,0,1.0-10/

```

```

DO 111 I=1,JI

```

```

K(I)=0

```

```

111 CONTINUE

```

IT(I TOTAL) IS THE NUMBER OF ROWS.

READ IN THE MATRIX BY COLUMNS, NO MORE THAN 10 ENTRIES PER CARD.

```

M=0

```

```

JI=0

```

SET ROW INDEXES, L(I).

```

DO 5 I=1,IT

```

```

5 L(I)=1

```

```

DO 66 J=1,IT

```

THE SUBROUTINE FINDS THE PIVOT ROW FOR EACH COLUMN.

```

CALL MAX(J,IT,M)

```

```

K(J)=M

```

~~THE PIVOT IS ZERO OR NEAR TO ZERO. THE METHOD IS INAPPLICABLE.~~

```

IF(DABS(A(M,J)).LE.RR) GO TO 60

```

REDUCE ROW M, THE PIVOT ROW, STARTING AT COLUMN J+1.

```

7 JI=J+1

```

```

DO 8 JP=JI,JT

```

```

8 A(M,JP)=A(M,JP)/A(M,J)

```

~~REDUCE OTHER ROWS, STARTING IN COLUMN JI.~~

```

DO 66 IP=1,IT

```

THE NEXT STEP SKIPS THE PIVOT ROW.

```

IF(IP.EQ.K)GO TO 66

```

```

DO 6 JP=JI,JT

```

```

A(IP,JP)=A(IP,JP)-A(M,JP)*A(IP,J)

```

```

CONTINUE

```

```

66 CONTINUE

```

STORAGE OF THE ANSWERS.

```

DO 10 I=1,IT

```

```

M=K(I)

```

```

10 T(I)=A(M,JT)

```

DISPLAY THE ANSWERS.

```

MI0=MI0+1

```

```

WRITE(2,103)MI0

```

```

103 FORMAT(1X,///T4,'CYCLE ',I3)

```

```

DO 51 I=1,IT

```

```

51 WRITE(3,104)I,T(I)

```

```

104 FORMAT(1X, /T4,'X(',I2,',') = ',D12.5)

```

APPENDIX - E

	1	2	3	4	5	6
15341	10.1	26.50	25.1	86.12		
7053	10.2	23.50	23.2	86.13		
14620	10.3	30.25	23.0	86.15		
13216	06.9	47.75	24.0	86.17		
10312	05.6	55.50	22.1	86.18		
9363	04.7	61.25	21.5	86.21		
14744	09.9	59.00	22.7	86.23		
15235	10.3	30.50	23.6	86.25		
15004	10.1	29.00	23.5	86.27		
11456	10.0	53.00	23.0	86.30		
14433	10.1	47.00	24.0	86.32		
14991	09.5	40.25	23.5	86.35		
15280	10.0	47.00	23.5	86.38		
13468	07.3	36.50	23.0	86.41		
13810	09.8	41.50	25.0	86.54		
13867	09.1	54.75	25.4	86.57		
13526	09.2	56.75	24.5	86.62		
10082	06.5	57.75	24.0	86.65		
13840	08.5	53.75	26.8	86.72		
9829	02.6	62.25	23.0	86.76		
10098	04.5	60.75	24.0	86.80		
11411	05.6	60.00	25.1	86.85		
8712	02.9	60.60	22.5	86.89		
12368	06.8	57.25	25.2	86.93		
10956	05.8	57.75	25.5	86.93		
11486	07.3	51.25	25.0	87.02		
5216	02.7	74.00	21.2	87.07		
7637	02.6	73.00	22.0	87.11		
7989	03.0	74.00	22.6	87.15		
10205	04.7	65.50	23.2	87.21		
8379	05.0	64.25	22.6	87.26		
12870	08.5	51.25	24.3	87.31		
13485	08.5	45.70	24.0	87.35		
13951	09.5	46.50	26.0	87.41		
13638	07.5	50.00	25.5	87.46		
11912	07.0	55.50	25.8	87.51		
10087	04.6	58.25	23.0	87.57		
13590	09.2	31.50	27.0	87.62		
10938	07.5	46.50	26.5	87.67		
10390	09.6	46.50	25.5	87.73		
13285	08.4	51.00	25.0	87.78		
11136	07.2	53.25	24.5	87.84		
13733	09.8	49.25	26.3	87.90		
13139	08.5	42.00	27.4	87.95		
13350	08.3	50.50	25.0	88.04		
10831	04.0	55.00	24.5	88.07		
12445	08.0	49.50	24.5	88.13		
11008	05.2	59.75	24.5	88.24		
9674	09.4	50.75	25.0	88.30		
13991	10.3	43.50	25.0	88.36		
17339	10.5	24.75	27.0	88.42		
11722	10.6	30.75	27.5	88.43		
13717	09.0	48.75	26.5	88.60		
16975	10.2	37.50	26.0	88.66		
14710	08.4	51.00	27.5	88.79		

.....1.....2.....3.....4.....5.....6.....
16358	09.2 5.00	28.5		88.85	
12005	08.4 34.00	28.7		88.91	
14324	08.7 37.50	27.0		88.98	
13816	07.1 43.75	27.2		89.04	
13314	07.5 33.75	27.0		89.10	
11664	04.5 48.00	27.5		89.16	
14007	06.5 62.00	27.0		89.23	
11637	05.4 47.00	25.8		89.29	
10638	04.1 48.25	25.6		89.35	
08637	01.8 67.50	24.0		89.42	
06250	00.3 65.00	21.5		89.48	
13204	05.2 52.25	26.0		89.55	
07922	01.1 63.00	24.7		89.61	
10638	04.0 54.00	24.5		89.67	
09490	04.0 49.50	25.7		89.74	
13096	09.5 58.50	26.4		89.80	
16843	10.2 45.00	27.7		89.87	
15352	09.0 36.00	28.5		89.93	
11784	03.2 52.00	26.2		90.00	
14612	03.2 58.00	22.4		90.06	
09708	03.3 59.25	24.5		90.13	
11801	06.8 63.00	24.3		90.25	
10690	04.6 63.50	24.5		90.33	
13798	05.2 49.25	25.6		90.38	
11859	06.4 54.00	24.6		90.44	
13245	08.8 46.50	26.5		90.51	
13510	05.2 50.75	26.5		90.57	
12651	06.9 51.25	27.0		90.64	
12699	06.0 80.25	26.0		90.70	
09104	02.6 55.75	21.3		90.76	
14114	07.5 53.75	25.4		90.83	
10554	04.8 49.25	25.1		90.95	
11095	05.4 71.50	25.5		91.01	
08131	02.4 50.50	21.6		91.08	
16232	08.4 40.50	25.8		91.20	
10867	03.4 65.00	24.5		91.26	
15896	08.8 72.25	25.4		91.33	
14472	09.8 51.00	24.8		91.39	
16087	10.1 59.50	26.2		91.45	
14813	09.3 60.25	27.0		91.51	
11516	03.6 43.25	24.5		91.57	
14984	06.2 65.75	25.1		91.69	
17267	09.0 75.25	24.5		91.75	
13250	04.6 45.25	24.5		91.81	
10788	01.5 49.00	21.2		91.86	
3563	00.1 53.25	18.5		91.92	
11707	03.8 67.25	23.5		91.98	
17764	09.9 53.50	25.3		92.04	
14470	07.4 60.50	26.1		92.09	
09085	01.7 65.50	22.8		92.15	
11827	03.4 61.00	24.0		92.21	
14250	06.1 54.00	25.2		92.26	
18587	11.0 47.00	24.9		92.32	
17817	10.0 66.25	26.1		92.37	
16908	09.9 56.25	25.0		93.75	

.....1.....2.....3.....4.....5.....6.....
13760	07.3	64.75	23.5	93.76	
14111	08.6	46.00	25.0	93.79	
16480	11.3	63.50	24.4	93.81	
17841	11.2	50.00	25.8	93.82	
13453	05.7	44.25	24.5	93.84	
14341	08.3	46.00	26.2	93.86	
16490	09.9	45.00	25.7	93.87	
16015	09.2	39.00	27.0	93.89	
13172	05.3	42.25	26.3	93.90	
17614	10.9	35.75	27.2	93.91	
13593	06.7	81.00	26.5	93.92	
12119	06.4	84.00	23.5	93.93	
09490	03.7	75.50	21.5	93.94	
12819	05.5	72.22	22.3	93.94	
10071	03.2	77.25	21.0	93.95	
12432	06.0	57.50	23.5	93.95	
10441	06.1	55.75	25.0	93.96	
11666	04.7	62.50	23.0	93.96	
11955	04.6	64.25	24.2	93.96	
13393	06.1	55.25	23.3	93.95	
12486	06.1	51.00	25.3	93.95	
11730	04.7	63.25	24.2	93.94	
07312	02.3	68.50	21.0	93.93	
10611	05.3	69.50	21.1	93.94	
06637	01.7	77.75	19.5	93.91	
09307	07.5	72.00	20.0	93.92	
13312	07.5	66.00	21.5	93.91	
12302	06.8	68.00	23.8	93.84	
14432	03.4	68.25	25.0	93.88	
12624	07.1	61.75	23.8	93.87	
15044	09.3	57.75	25.1	93.85	
15003	09.1	60.00	25.2	93.83	
10036	04.0	63.00	21.5	93.61	
13081	06.9	62.75	23.0	93.80	
13763	07.1	61.50	24.0	93.77	
14486	09.3	49.00	25.2	93.75	
14026	07.4	53.75	24.5	93.73	
11031	06.1	63.00	24.0	93.70	
16384	08.4	58.75	23.3	93.67	
07556	03.2	64.00	23.5	93.65	
09985	04.1	68.25	23.0	93.62	
10881	03.5	65.50	23.8	93.60	
08610	02.6	79.00	19.2	93.57	
08091	03.6	75.66	21.6	93.53	
10752	04.4	75.50	22.2	93.50	
06844	02.7	71.50	21.0	93.47	
08291	03.1	70.75	22.2	93.43	
09736	03.4	68.25	22.5	93.40	
09471	02.3	71.25	20.6	93.36	
06033	02.5	74.00	21.9	93.32	
07759	03.2	79.50	21.2	93.29	
08347	03.5	70.00	21.2	93.25	
08696	02.5	71.25	21.0	93.21	
10046	05.5	7.50	21.8	93.17	
09210	03.3	75.75	22.0	93.12	

	1	2	3	4	5	6
09931	05.1	73.75	21.0	93.18		
05466	00.6	77.50	20.0	93.14		
08298	02.3	74.00	20.4	92.99		
11060	05.6	71.00	23.3	92.95		
09008	03.7	62.25	23.1	92.90		
09629	02.8	58.75	22.9	92.85		
09902	05.6	63.75	22.2	92.80		
09224	03.4	73.50	22.0	92.76		
09070	02.3	69.25	22.5	92.71		
08946	04.4	73.75	20.2	92.66		
13720	06.9	69.25	23.5	92.61		
11095	04.8	67.75	23.0	92.55		
05906	00.6	69.75	17.6	92.50		
07638	00.6	82.75	18.6	92.45		
05980	01.4	79.00	19.2	92.40		
04049	00.1	89.25	17.5	92.34		
01441	00.1	97.50	13.5	92.29		
07282	00.7	74.75	19.0	92.23		
06940	02.2	86.00	20.4	92.19		
13269	01.1	80.50	19.6	92.12		
08145	01.3	68.00	20.7	92.07		
11760	05.1	69.25	22.9	91.84		
02537	00.1	84.00	16.8	91.78		
06262	00.5	80.25	19.3	91.72		
07163	02.3	74.75	19.5	91.66		
11183	04.7	71.50	21.1	91.60		
10326	05.4	76.25	21.5	91.54		
10165	03.7	65.00	21.6	91.48		
08985	02.3	76.75	21.0	91.42		
05796	00.7	81.25	18.9	91.36		
09248	03.5	74.00	20.5	91.29		
06546	00.5	75.50	20.2	91.23		
10857	04.4	71.50	20.7	91.17		
12110	06.4	63.50	22.8	91.11		
11098	08.8	57.75	23.8	91.05		
08415	03.9	68.75	20.9	90.98		
15414	08.9	61.50	23.9	90.92		
13249	07.7	62.75	23.9	90.86		
06995	02.2	78.75	20.4	90.79		
06790	03.2	73.00	19.9	90.73		
10520	03.8	77.25	22.0	90.67		
13858	09.0	59.50	23.5	90.60		
08292	02.8	70.50	20.5	90.54		
12731	06.6	71.50	21.8	90.48		
07705	01.4	71.75	19.7	90.41		
08009	03.8	70.50	20.7	90.35		
10420	05.2	76.50	21.1	90.28		
05443	01.7	73.00	21.0	90.22		
06604	01.5	75.75	19.5	90.16		
08232	02.6	63.75	21.0	90.09		
11640	04.5	48.50	22.7	90.03		
14505	08.6	44.25	23.8	89.96		
15954	10.0	61.00	23.6	89.90		
11438	04.4	69.75	22.7	89.83		
09888	04.8	54.50	22.4	89.77		

	1	2	3	4	5	6
12494	05.1	71.00	20.5		89.71	
07507	03.1	74.75	20.8		89.64	
07940	01.3	62.25	22.3		89.58	
11171	04.8	64.25	23.7		89.51	
13388	05.8	74.50	19.6		89.54	
08889	07.3	52.25	23.5		89.59	
11978	05.9	57.00	22.5		89.32	
10184	03.3	58.75	23.5		89.26	
14134	07.2	46.75	23.3		89.20	
11018	04.7	67.75	22.5		89.13	
08194	03.9	75.25	22.5		89.07	
12563	04.5	59.75	22.5		89.01	
17570	10.6	30.75	22.9		88.88	
17248	10.4	29.50	22.1		88.82	
15710	08.8	40.00	23.0		88.76	
15732	07.9	39.75	23.6		88.70	
15746	09.3	41.00	24.3		88.63	
12306	05.7	43.50	23.0		88.57	
13945	06.2	50.50	23.5		88.51	
11350	06.3	49.00	24.1		88.45	
09036	03.5	74.75	20.7		88.39	
07941	01.4	76.00	20.6		88.33	
08447	01.8	75.50	20.2		88.27	
13779	07.1	50.50	22.8		88.15	
12852	07.8	47.75	22.3		88.10	
12903	07.0	46.50	21.8		88.04	
17375	10.8	41.50	23.3		87.98	
16828	10.8	36.75	22.5		87.92	
15130	10.3	43.00	23.4		87.87	
15881	09.0	48.00	24.2		87.81	
17598	10.8	41.50	23.5		87.76	
17637	10.8	36.00	23.4		87.70	
16882	10.8	40.00	23.4		87.65	
17576	10.8	19.25	23.0		87.59	
17070	10.8	34.00	22.9		87.54	
17122	10.4	31.25	22.8		87.49	
15264	10.1	42.25	23.0		87.44	
11780	06.5	42.25	23.6		87.38	
14475	10.2	41.00	24.7		87.33	
15299	10.6	38.00	24.2		87.28	
13722	07.9	47.75	24.0		87.23	
11234	04.1	49.00	23.0		87.19	
15572	10.6	44.50	24.0		87.14	
15226	09.6	40.00	24.5		87.09	
14865	10.6	44.25	24.6		87.04	
15630	10.1	41.25	25.5		87.00	
15579	10.1	41.75	25.1		86.95	
14216	10.0	47.75	24.2		86.91	
15167	08.9	45.50	25.1		86.87	
13425	09.8	43.00	24.3		86.82	
12562	07.0	50.00	23.2		86.78	
13042	09.3	47.75	23.0		86.74	
16203	10.8	38.25	23.2		86.70	
16498	10.7	48.50	23.6		86.67	
16018	10.9	40.00	23.0		86.63	

L. COPRNT RELEASE

EXTENDED A SET CARDPRNTOL

	1	2	3	4	5	6
15051	10.6	31.00	24.4	86.59		
16400	10.8	40.75	22.9	86.52		
15949	10.6	32.50	22.8	86.49		
15996	10.8	36.75	23.0	86.46		
14777	08.9	37.00	23.0	86.42		
15159	10.6	36.25	23.4	86.39		
14294	10.3	41.25	23.8	86.37		
15448	10.6	32.25	23.4	86.34		
10304	10.9	38.25	22.6	86.31		
15155	10.6	45.00	23.3	86.29		
15556	10.3	46.25	23.5	86.26		
15191	09.4	49.25	22.1	86.24		
15188	10.1	47.75	22.0	86.22		
15461	10.7	48.50	22.8	86.19		
15237	10.5	41.75	23.6	86.18		
15039	10.6	40.75	23.3	86.16		
15364	10.6	33.25	22.9	86.14		
14108	10.5	42.00	23.5	86.12		
15900	10.8	30.25	23.7	86.11		
15366	10.5	26.75	23.5	86.10		
15128	10.7	24.25	23.5	86.08		
14952	10.6	24.50	23.2	86.07		
15150	10.0	22.25	22.7	86.06		
15111	10.7	25.00	24.2	86.05		
14453	10.7	22.50	24.5	86.05		
14686	09.9	28.75	24.3	86.04		
15099	10.6	36.00	23.3	86.04		
13636	10.3	37.00	23.8	86.04		
13598	08.8	37.75	24.5	86.03		
10847	06.4	42.75	22.5	86.03		
10302	05.4	57.75	22.6	86.03		
05688	01.1	71.00	19.7	86.04		
12282	08.0	53.00	22.9	86.05		
09308	06.0	57.25	22.6	86.05		
09457	04.5	63.75	21.9	86.06		
09985	05.6	58.50	22.5	86.07		
13952	09.9	42.25	23.8	86.08		
13251	10.3	38.75	24.5	86.10		

APPENDIX - F

MEASURED	SAYIGH S	NEW CONST
15341.000E 03	17921.070E 03	17823.120E 03
70530.000E 02	17937.250E 03	14570.700E 03
14620.000E 03	17960.140E 03	13609.400E 03
13216.000E 03	17179.660E 03	11272.260E 03
10312.000E 03	16879.200E 03	77223.870E 02
93630.000E 02	16679.660E 03	64138.470E 02
14744.000E 03	17867.810E 03	11109.790E 03
15235.000E 03	17973.390E 03	14650.580E 03
15004.000E 03	17928.080E 03	14446.070E 03
11456.000E 03	17900.830E 03	11907.730E 03
14433.000E 03	17932.400E 03	13932.270E 03
14991.000E 03	17793.280E 03	13086.450E 03
15280.000E 03	17912.450E 03	13030.600E 03
13468.000E 03	17296.540E 03	10874.610E 03
13810.000E 03	17889.820E 03	15941.020E 03
13867.000E 03	17728.770E 03	15037.700E 03
13526.000E 03	17752.770E 03	13472.420E 03
10082.000E 03	17134.770E 03	10487.610E 03
13840.000E 03	17610.380E 03	17142.000E 03
99290.000E 02	16284.150E 03	59886.100E 02
10098.000E 03	16704.330E 03	86633.140E 02
11411.000E 03	16957.150E 03	11025.480E 03
87120.000E 02	16358.160E 03	59245.320E 02
12368.000E 03	17235.090E 03	12452.610E 03
10956.000E 03	17016.210E 03	11882.670E 03
11486.000E 03	17357.390E 03	12956.170E 03
52160.000E 02	16323.300E 03	46264.280E 02
76370.000E 02	17003.180E 03	54390.050E 02
79880.000E 02	17055.660E 03	60237.440E 02
10205.000E 03	17449.600E 03	81921.220E 02
83790.000E 02	17520.240E 03	78635.710E 02
12870.000E 03	18360.110E 03	13455.800E 03
13485.000E 03	18364.450E 03	13314.440E 03
13951.000E 03	18622.580E 03	18468.110E 03
13638.000E 03	18140.990E 03	14619.440E 03
11912.000E 03	18027.630E 03	14260.620E 03
10087.000E 03	17463.620E 03	87900.360E 02
13590.000E 03	18576.340E 03	21440.190E 03
10938.000E 03	18166.800E 03	16916.940E 03
10390.000E 03	18675.780E 03	16872.000E 03
13285.000E 03	18386.060E 03	14599.950E 03
11136.000E 03	18101.780E 03	12537.940E 03
13733.000E 03	18745.650E 03	18787.200E 03
13139.000E 03	18438.770E 03	20305.630E 03
13350.000E 03	18387.790E 03	14558.190E 03
10831.000E 03	17376.820E 03	89795.930E 02
12445.000E 03	18327.100E 03	13509.440E 03
11008.000E 03	17669.700E 03	10376.260E 03
96740.000E 02	18681.490E 03	15597.990E 03
13991.000E 03	18910.910E 03	17006.500E 03
17338.990E 03	18979.520E 03	24252.910E 03
11722.000E 03	19011.200E 03	24824.990E 03
13717.000E 03	18620.580E 03	18358.140E 03
16974.990E 03	20070.420E 03	20862.610E 03
14710.000E 03	19618.500E 03	20826.910E 03
16358.000E 03	19838.670E 03	26556.820E 03
12005.000E 03	19638.260E 03	25979.710E 03
14324.000E 03	19715.230E 03	21456.510E 03
13816.000E 03	19310.800E 03	19064.110E 03
13314.000E 03	19418.500E 03	19683.280E 03

Cont'd

MEASURED	SAYIGH'S	NEW CONSTANTS
12119.000E 02	2417.190E 02	12119.000E 02
12490.000E 02	24391.290E 03	12490.000E 02
12519.000E 02	24401.730E 03	12519.000E 02
12571.000E 03	24731.420E 03	12571.000E 03
12432.000E 03	24140.380E 03	12432.000E 03
10441.000E 03	24100.780E 03	10441.000E 03
11666.000E 03	24222.700E 03	11666.000E 03
11955.000E 03	24198.330E 03	11955.000E 03
13393.000E 03	24674.750E 03	13393.000E 03
12486.000E 03	24100.300E 03	12486.000E 03
11730.000E 03	24225.710E 03	11730.000E 03
73120.000E 02	23450.100E 03	73120.000E 02
10611.000E 03	24390.190E 03	10611.000E 03
65370.000E 02	23251.360E 03	65370.000E 02
93070.000E 02	21342.700E 03	93070.000E 02
13312.000E 03	24307.280E 03	13312.000E 03
12302.000E 03	26132.290E 03	12302.000E 03
14432.000E 03	26694.360E 03	14432.000E 03
12624.000E 03	26234.620E 03	12624.000E 03
15044.000E 03	27011.760E 03	15044.000E 03
15003.000E 03	26939.100E 03	15003.000E 03
10036.000E 03	25104.320E 03	10036.000E 03
13081.000E 03	26152.300E 03	13081.000E 03
13763.000E 03	26226.360E 03	13763.000E 03
14486.000E 03	27004.270E 03	14486.000E 03
14026.000E 03	26331.040E 03	14026.000E 03
11031.000E 03	25377.130E 03	11031.000E 03
16384.000E 03	26064.320E 03	16384.000E 03
75560.000E 02	24901.340E 03	75560.000E 02
99850.000E 02	25190.690E 03	99850.000E 02
10881.000E 03	24997.330E 03	10881.000E 03
86100.000E 02	24659.530E 03	86100.000E 02
80910.000E 02	25005.100E 03	80910.000E 02
10752.000E 03	25271.380E 03	10752.000E 03
68440.000E 02	24701.260E 03	68440.000E 02
82910.000E 02	24837.370E 03	82910.000E 02
97360.000E 02	24935.900E 03	97360.000E 02
94710.000E 02	24558.240E 03	94710.000E 02
60330.000E 02	24629.300E 03	60330.000E 02
77590.000E 02	24847.340E 03	77590.000E 02
83470.000E 02	24943.730E 03	83470.000E 02
86960.000E 02	24812.660E 03	86960.000E 02
10046.000E 03	25005.100E 03	10046.000E 03
92100.000E 02	24872.030E 03	92100.000E 02
99310.000E 02	25454.330E 03	99310.000E 02
54660.000E 02	23539.040E 03	54660.000E 02
82980.000E 02	24070.160E 03	82980.000E 02
11060.000E 03	25157.140E 03	11060.000E 03
90030.000E 02	24532.540E 03	90030.000E 02
96290.000E 02	24233.050E 03	96290.000E 02
99020.000E 02	25135.030E 03	99020.000E 02
92240.000E 02	24412.360E 03	92240.000E 02
90700.000E 02	24062.020E 03	90700.000E 02
89450.000E 02	24710.350E 03	89450.000E 02
13720.000E 03	25559.040E 03	13720.000E 03
11095.000E 03	24654.730E 03	11095.000E 03
59060.000E 02	23455.010E 03	59060.000E 02
70380.000E 02	23471.050E 03	70380.000E 02
53800.000E 02	23721.790E 03	53800.000E 02
40490.000E 02	23294.360E 03	40490.000E 02
14410.000E 02	23227.310E 03	14410.000E 02

Measured	CONT'd	SAYIGH'S	NEW CONSTANTS
11554.000E	03	10774.340E	10737.340E
14767.000E	03	12777.300E	12727.300E
11557.000E	03	10771.300E	10727.300E
11638.000E	03	10770.300E	10727.300E
95370.000E	02	17779.100E	17752.270E
52500.000E	02	17579.100E	17752.270E
13204.000E	02	18677.070E	18672.980E
79220.000E	02	17885.730E	69139.230E
10638.000E	03	18535.000E	18045.870E
94900.000E	02	18598.320E	11231.800E
13896.000E	03	19499.000E	19127.380E
15842.990E	03	20199.570E	20057.240E
15352.000E	03	19894.520E	20221.570E
11784.000E	03	18430.350E	11438.000E
14612.000E	02	18418.210E	69139.810E
27080.000E	02	18443.460E	69443.840E
11801.000E	03	19336.250E	12251.100E
10590.000E	03	18794.850E	18336.970E
13798.000E	03	19716.530E	16768.170E
11859.000E	03	19256.140E	12794.380E
13245.000E	03	19590.130E	19524.130E
13510.000E	03	19739.230E	18514.430E
12651.000E	03	20994.560E	19319.650E
12699.000E	03	20744.750E	14326.180E
91040.000E	02	19821.140E	60237.980E
14114.000E	03	21176.990E	16587.030E
10554.000E	03	20444.110E	13059.800E
11095.000E	03	20610.220E	13241.600E
31310.000E	02	19801.340E	62151.600E
16232.000E	03	21464.130E	19706.620E
10867.000E	03	20094.960E	97380.180E
15896.000E	03	21582.720E	16841.550E
14472.000E	03	21874.210E	18139.030E
16487.000E	03	21972.560E	21004.210E
14818.000E	03	21752.300E	21925.970E
11516.000E	03	20180.240E	10794.690E
14984.000E	03	20892.270E	13651.620E
17266.990E	03	21676.180E	15175.950E
13250.000E	03	20469.330E	12185.100E
19788.000E	03	19430.380E	49954.850E
35630.000E	02	19262.900E	16835.870E
11707.000E	03	20262.240E	90012.480E
17764.000E	03	21971.120E	19211.740E
14470.000E	03	21269.280E	17480.580E
90850.000E	02	19722.690E	59987.470E
11927.000E	03	20431.200E	93497.490E
14250.000E	03	20922.110E	14634.530E
18586.990E	02	22316.680E	20074.430E
17816.990E	03	22335.530E	20155.250E
16904.000E	03	25926.300E	21621.150E
13760.000E	03	25047.740E	14733.900E
14111.000E	03	25495.100E	20435.400E
16480.000E	03	26410.030E	20625.590E
17840.990E	03	26367.510E	21731.020E
13453.000E	03	24345.330E	16047.140E
14341.000E	03	25410.110E	23661.730E
15490.000E	03	25947.920E	14732.820E
15815.000E	03	25711.520E	20454.430E
13172.000E	03	24433.550E	19364.710E
17614.000E	03	26366.300E	22640.600E
13593.000E	03	24060.300E	18130.000E

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MEASURED	SAVIGN'S	NEW CONSTANTS
73320.000E 02	25412.70E 03	51557.390E 02
59400.000E 02	25262.70E 03	50317.710E 02
13259.000E 03	25112.70E 03	49078.030E 02
31450.000E 02	25072.70E 03	47838.350E 02
11760.000E 03	24932.70E 03	46598.670E 02
25370.000E 02	24855.70E 03	45358.990E 02
62620.000E 02	24803.70E 03	44119.310E 02
71630.000E 02	24739.70E 03	42879.630E 02
11183.000E 03	24712.370E 03	41639.950E 02
13326.000E 03	24641.30E 03	40399.510E 02
10165.000E 03	24577.30E 03	39159.070E 02
89850.000E 02	24531.30E 03	37918.630E 02
57960.000E 02	24488.10E 03	36678.190E 02
92480.000E 02	24397.42E 03	35437.750E 02
65460.000E 02	23633.25E 03	33257.620E 02
10357.000E 03	24243.230E 03	78559.420E 02
12110.000E 03	24901.540E 03	12749.650E 03
11098.000E 03	25701.520E 03	17494.180E 03
84150.000E 02	24069.060E 03	77070.230E 02
15414.000E 03	25722.270E 03	17541.630E 03
13249.000E 03	25313.240E 03	16130.160E 03
69950.000E 02	23514.000E 03	54575.800E 02
67900.000E 02	25815.250E 03	60767.870E 02
10520.000E 03	24016.260E 03	35135.830E 02
13858.000E 03	25721.570E 03	16916.560E 03
82920.000E 02	23679.540E 03	62310.870E 02
12731.000E 03	24896.740E 03	11032.780E 03
77050.000E 02	23228.930E 03	43079.350E 02
80090.000E 02	23976.520E 03	73520.340E 02
10420.000E 03	24417.470E 03	87764.260E 02
54430.000E 02	23314.740E 03	54704.900E 02
66040.000E 02	23234.300E 03	42288.490E 02
82320.000E 02	23581.470E 03	66691.470E 02
11640.000E 03	24188.100E 03	11254.890E 03
14505.000E 03	25528.420E 03	16281.040E 03
15954.000E 03	25936.560E 03	16716.190E 03
11438.000E 03	24133.680E 03	10194.950E 03
98880.000E 02	24255.540E 03	10861.580E 03
12494.000E 03	24326.830E 03	81699.180E 02
75070.000E 02	23691.720E 03	60952.510E 02
79490.000E 02	23146.560E 03	61190.800E 02
11171.000E 03	24238.500E 03	12269.980E 03
13388.000E 03	24526.000E 03	75937.790E 02
88890.000E 02	20658.220E 03	12813.440E 03
11978.000E 03	20297.760E 03	99253.140E 02
10184.000E 03	19616.330E 03	85254.770E 02
14134.000E 03	20642.260E 03	12714.010E 03
11018.000E 03	19960.380E 03	85091.300E 02
31940.000E 02	19745.550E 03	75953.120E 02
12563.000E 03	19076.320E 03	65979.640E 02
17576.000E 03	21553.190E 03	10262.620E 03
17248.000E 03	21455.390E 03	14358.320E 03
15716.000E 03	21036.140E 03	14078.990E 03
15732.000E 03	20788.720E 03	14310.550E 03
15746.000E 03	21169.230E 03	10940.770E 03
12306.000E 03	20178.020E 03	11011.770E 03
13945.000E 03	20005.910E 03	11021.430E 03
11350.000E 03	20331.420E 03	12918.970E 03
90360.000E 02	19907.190E 03	57199.970E 02
79410.000E 02	19023.510E 03	39244.270E 02
34470.000E 02	19121.430E 03	40763.590E 02

MEASURED	Cont'd	SAYRE'S	NEW CONSTANTS
13779.000E 03	20509.010E 03	11532.840E 03	
12352.000E 03	20290.770E 03	1157 .000E 02	
12903.000E 03	20465.570E 03	12312.500E 03	
17374.990E 03	21517.700E 03	15225.920E 02	
15828.000E 03	21507.360E 03	1542.950E 03	
15130.000E 03	21365.970E 03	15556.650E 03	
15881.000E 03	21602.020E 03	15822.820E 03	
17598.000E 03	21496.380E 03	16616.750E 03	
17636.990E 03	21490.780E 03	16895.710E 03	
15382.000E 03	21434.300E 03	16530.800E 03	
17576.000E 03	21481.540E 03	17966.960E 03	
17070.000E 03	18942.270E 03	14151.680E 03	
17122.000E 03	18837.950E 03	13394.330E 03	
15264.000E 03	18757.790E 03	13120.750E 03	
11780.000E 03	17858.860E 03	11189.420E 03	
14475.000E 03	18780.350E 03	16429.350E 03	
15299.000E 03	18872.210E 03	16057.290E 03	
13722.000E 03	18207.170E 03	12663.800E 03	
11234.000E 03	17311.920E 03	80367.960E 02	
15572.000E 03	18855.710E 03	15144.040E 03	
15226.000E 03	18607.940E 03	15554.910E 03	
14865.000E 03	18848.560E 03	16301.100E 03	
15630.000E 03	18725.870E 03	17918.510E 03	
15579.000E 03	18718.860E 03	17038.800E 03	
14216.000E 03	18684.860E 03	14758.190E 03	
15167.000E 03	18417.550E 03	15570.760E 03	
13425.000E 03	18628.050E 03	15104.350E 03	
12562.000E 03	17944.940E 03	10623.510E 03	
13042.000E 03	18490.510E 03	12201.870E 03	
16203.000E 03	18857.060E 03	14284.680E 03	
16498.000E 03	18829.650E 03	14185.640E 03	
16018.000E 03	18873.220E 03	13871.910E 03	
15051.000E 03	18804.100E 03	17062.930E 03	
16400.000E 03	18836.480E 03	13554.400E 03	
15949.000E 03	18785.020E 03	13866.990E 03	
15996.000E 03	18831.650E 03	14027.480E 03	
14777.000E 03	18363.500E 03	12550.590E 03	
15159.000E 03	18777.010E 03	14462.720E 03	
14294.000E 03	18702.700E 03	14711.480E 03	
15448.000E 03	18773.090E 03	14958.250E 03	
10304.000E 03	17318.980E 03	12217.610E 03	
15155.000E 03	17251.340E 03	12679.900E 03	
15556.000E 03	17181.260E 03	12706.790E 03	
15191.000E 03	16969.710E 03	99701.710E 02	
15188.000E 03	17124.080E 03	10321.500E 03	
15461.000E 03	17260.700E 03	11770.630E 03	
15237.000E 03	17219.470E 03	13293.530E 03	
15039.000E 03	17238.750E 03	12942.800E 03	
15364.000E 03	17235.980E 03	12827.560E 03	
14108.000E 03	17212.320E 03	13108.740E 03	
15900.000E 03	17282.350E 03	14534.500E 02	
15366.000E 03	17213.520E 03	14317.330E 03	
15128.000E 03	17257.420E 03	14724.920E 03	
14952.000E 03	17235.140E 03	15159.510E 03	
15150.000E 03	17093.890E 03	13001.920E 03	
15111.000E 03	17257.520E 03	15959.720E 03	
14453.000E 03	17259.490E 03	16353.420E 03	
14686.000E 03	17075.790E 03	15056.930E 03	
15099.000E 03	17227.310E 03	13267.530E 03	
13636.000E 03	17161.320E 03	15306.460E 03	
13598.000E 03	16829.100E 03	13706.040E 03	

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